



Guide to the Science of Climate Change in the 21st Century

Chapter 4 Earth's Energy Budget

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https://www.ipcc.ch/report/ar6/wg1/downloads/report/IPCC_AR6_WGI_Chapter07.pdf

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4.0 Earth's Energy Budget

4.1 Introduction

Earth's energy budget forms the framework with which it is possible to discuss the science of climate change. All of the energy that 'fuels' planet Earth comes from the Sun – solar radiation. All energy is measured in Watts per square metre or Wm^{-2} . (Recall, the brightness of light bulbs is measured in Watts. Visible light is short wave radiation.) Short wave radiation that is not immediately reflected back into space is absorbed by the atmosphere and the surface of the Earth. The exchange of energy between the surface of the Earth and the atmosphere results in all of the absorbed energy being emitted into space in the form of long wave radiation. In the process both the surface of the Earth and the atmosphere warm. The amount of warmth will depend on the ability of the atmosphere and the Earth's surface to reflect short wave radiation and the ability of the atmosphere to absorb long wave radiation. Human activities can affect the amount of short-wave energy absorbed by changing land use and the ability of the atmosphere to absorb long wave radiation by changing the composition of the gases and aerosols that form the atmosphere. In so doing human activities influence how warm the Earth's surface and the atmosphere must become to return the absorbed solar energy to space and preserve the energy balance.

4.2 Concept of energy

It is very important to understand the concept of energy before discussing the Earth's energy budget.

Definitions and descriptions can be found in the following web sites, United States Energy Administration, Britannica, and Wikipedia.

<https://www.eia.gov/energyexplained/what-is-energy/>

<https://www.britannica.com/science/energy>

https://en.wikipedia.org/wiki/Outline_of_energy

The table below, taken from the US Energy Administration web site, identifies the different forms of energy.

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Forms of energy

Many forms of energy exist, but they all fall into two basic categories:

- Potential energy
- Kinetic energy

Potential energy

Potential energy is stored energy and the energy of position.

Chemical energy is energy stored in the bonds of atoms and molecules. Batteries, biomass, petroleum, natural gas, and coal are examples of chemical energy. Chemical energy is converted to thermal energy when people burn wood in a fireplace or burn gasoline in a car's engine.

Mechanical energy is energy stored in objects by tension. Compressed springs and stretched rubber bands are examples of stored mechanical energy.

Nuclear energy is energy stored in the nucleus of an atom—the energy that holds the nucleus together. Large amounts of energy can be released when the nuclei are combined or split apart.

Gravitational energy is energy stored in an object's height. The higher and heavier the object, the more gravitational energy is stored. When a person rides a bicycle down a steep hill and picks up speed, the gravitational energy is converting to motion energy. Hydropower is another example of gravitational energy, where gravity forces water down through a hydroelectric turbine to produce electricity.

Kinetic energy

Kinetic energy is the motion of waves, electrons, atoms, molecules, substances, and objects.

Radiant energy is electromagnetic energy that travels in transverse waves. Radiant energy includes visible light, x-rays, gamma rays, and radio waves. Light is one type of radiant energy. Sunshine is radiant energy, which provides the fuel and warmth that make life on earth possible.

Thermal energy, or heat, is the energy that comes from the movement of atoms and molecules in a substance. Heat increases when these particles move faster. Geothermal energy is the thermal energy in the earth.

Motion energy is energy stored in the movement of objects. The faster they move, the more energy is stored. It takes energy to get an object moving, and energy is released when an object slows down. Wind is an example of motion energy. A dramatic example of motion energy is a car crash—a car comes to a total stop and releases all of its motion energy at once in an uncontrolled instant.

Sound is the movement of energy through substances in longitudinal (compression/rarefaction) waves. Sound is produced when a force causes an object or substance to vibrate. The energy is transferred through the substance in a wave. Typically, the energy in sound is smaller than in other forms of energy.

Electrical energy is delivered by tiny charged particles called electrons, typically moving through a wire. Lightning is an example of electrical energy in nature.

Table 4.1 Forms of energy taken from <https://www.eia.gov/energyexplained/what-is-energy/>

Perhaps the best way to start to understand the concept of energy is how we use it or experience it. Consider the following examples?

1. Using a fuel burning stove to boil water. Natural gas, propane, kerosene, wood, coal, paraffin or some other fuel is burned to create heat which is transferred to the bottom of the kettle that then heats the water. The water temperature increases and steam is produced. The energy in the fuel is a 'chemical form' of potential energy. When burned the chemical energy produces 'heat or thermal energy' and 'radiant energy', forms of kinetic energy, which are transferred to the bottom of the pot which conducts it through the walls of the pot (a form of thermal energy transfer) to the water which warms ('thermal energy increasing') until it starts to steam (thermal energy leaving the water) by the process of evaporation.

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2. An incandescent light bulb uses 'electrical energy' to produce 'radiant energy'.
3. A fuel burning car uses the 'chemical energy' in the fuel transformed to 'mechanical energy'.
4. The sun uses 'nuclear energy' to produce 'radiant energy' which the earth and our skin absorbs and turns it into 'thermal energy'.

Energy is measured in terms of the 'work' it can do (potential energy) or the work it is doing (kinetic energy).

Work is defined in mechanical terms as equal to force times the distance over which is applied. Force is measured in Newtons (SI units) and distance is measured in metres. One unit of work is equal to one Newton times one metre. One unit of work is known as one Joule.

Potential energy is measured in Joules (J). Potential energy is the amount of work the energy can produce when transformed into kinetic energy.

Kinetic energy is measured in terms of rate of energy transfer or Watts (W). One Watt equals one Joule per second (J/s or $J s^{-1}$).

Energy, any form, passing through a unit area, e. g. one square metre, is known as the energy flux. One unit of energy flux equals one Watt per square metre (These units, W/m^2 or Wm^{-2} mean Watts per square meter). Energy flux is a measure of the movement of energy. If all the light from a 1000-Watt light bulb is shone through an area 1 m by 1 m, the energy flux is 1000 W per square meter. If the light bulb only produced 10 Watts, the energy flux would be 10 W per square meter. This is the most important term to understand.

4.3 Solar energy

The flux of radiant energy from the Sun to the outer surface of the Earth's atmosphere is $342 Wm^{-2}$. The layer of gases covering the Earth (the air we breath) is called the atmosphere. The 'top' or 'outside' of the atmosphere is 100 kilometers from the surface of the Earth. (Passenger jets fly about 10 kilometers above the surface of the Earth.) This is the average energy per square metre over the whole planet. It is very important to realize that the solar radiation is the sum of radiation of energy from the invisible ultraviolet, light that is visible to the human eye, to the thermal infrared as shown in Figure 4.1. The radiation moves in waves and a particular radiation is identified by its wavelength (See Figure 4.2). The energy of radiation with a specific wavelength is called spectral irradiance measured in Watts per square metre per

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wavelength (measured in micro metres, one millionth of a metre or microns which has the symbol, μm).

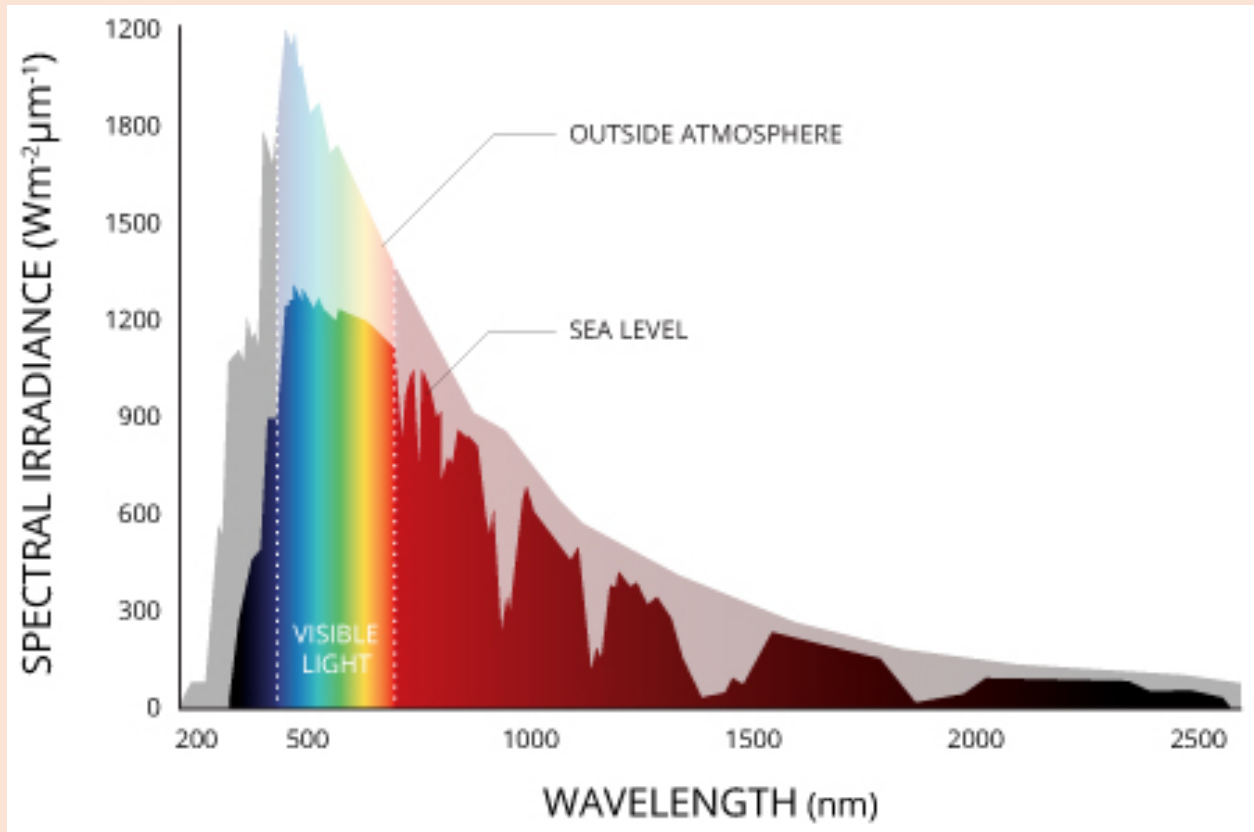


Figure 4.1 Sun's radiant energy that reaches Earth's atmosphere and surface.

<https://www.fondriest.com/environmental-measurements/parameters/weather/photosynthetically-active-radiation/>

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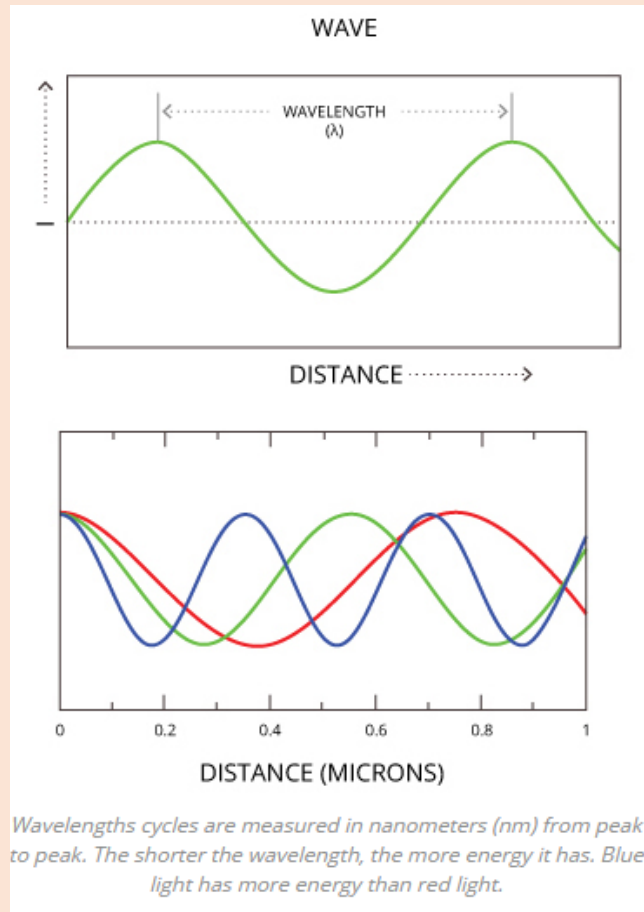


Figure 4.2 Wavelength description of radiant energy.

<https://www.fondriest.com/environmental-measurements/parameters/weather/photosynthetically-active-radiation/>

As shown in Figure 4.1 not all of the radiation reaches the earth's surface. As the radiation moves through the Earth's atmosphere certain wavelengths of radiation are absorbed by the different gases and particles (aerosols) that comprise the atmosphere. This includes nitrogen, oxygen, hydrogen, the inert gases such as helium (water, carbon dioxide, gases produced by volcanoes, industrial pollutants and particles from a variety of natural and industrial sources. So called greenhouse gases (GHG), like carbon dioxide, water vapour and gases and aerosols emitted by volcanoes are a natural part of the atmosphere. Different substances absorb specific wavelengths of radiation. Human activity resulting in changes in land use, burning of fossil fuels and introduction of a variety of industrial processes over the past two hundred years have greatly contributed to the variety and increase in greenhouse gases.

The distribution of radiation with wavelength is known as the solar radiation spectrum. As shown in Figures 4.1 and 4.3 this is typically divided into ultraviolet or UV, visible light and near

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infrared (where almost all of the Sun's radiant energy occurs). The spectrum of solar radiation outside or above the atmosphere, without atmospheric absorption (coloured yellow) is significantly different than that at sea level (coloured red). The visible radiation is a narrow band of the radiation received from the Sun. Infrared radiation refers to radiation with wavelengths greater than visible light. It is felt as warmth or heat because water (human body) is particularly effective in absorbing infrared radiation. Ultraviolet radiation, UV, is high intensity invisible radiation that is potentially damaging to living organisms (sunburns and melanoma type cancer).

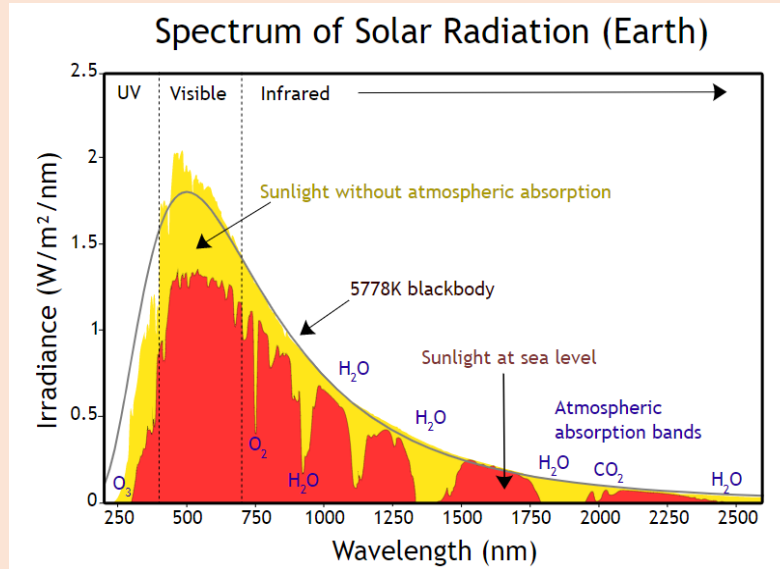


Figure 4.3 Absorption of Sun's radiant energy that reaches Earth's atmosphere and surface. [https://energyeducation.ca/encyclopedia/Solar energy to the Earth](https://energyeducation.ca/encyclopedia/Solar_energy_to_the_Earth)

Note that ozone, O₃, a molecule of oxygen made up of three atoms of oxygen, absorbs UV. (Ozone existing in a layer of the atmosphere high above the surface of the Earth protects Earth from harmful UV). Water and carbon dioxide absorb much of the infrared radiation. The absorption of all wavelengths of solar radiation is described in greater detail in Figure 4.4.

When radiation energy is absorbed, the effect is to increase the thermal energy of the substance absorbing the radiation – typically observed by increases in temperature of the substance absorbing the radiation.

The thermal energy (heat energy) of a substance is lost by emission of radiant energy in the form of long wave or infrared radiation (similar to the radiant energy one feels from the surface of hot stove). The rate at which the energy is emitted varies with the temperature of the substance. The radiation is sent in all directions. The warmer the substance, the greater the

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rate of emitted radiant energy (the greater the energy flux). The Earth is warm and so emits long wave or infrared radiation. This latter characteristic is very important to the understanding of global warming because water, carbon dioxide, sulfur dioxide and many industrially generated gases are very effective absorbers and emitters of longwave radiation.

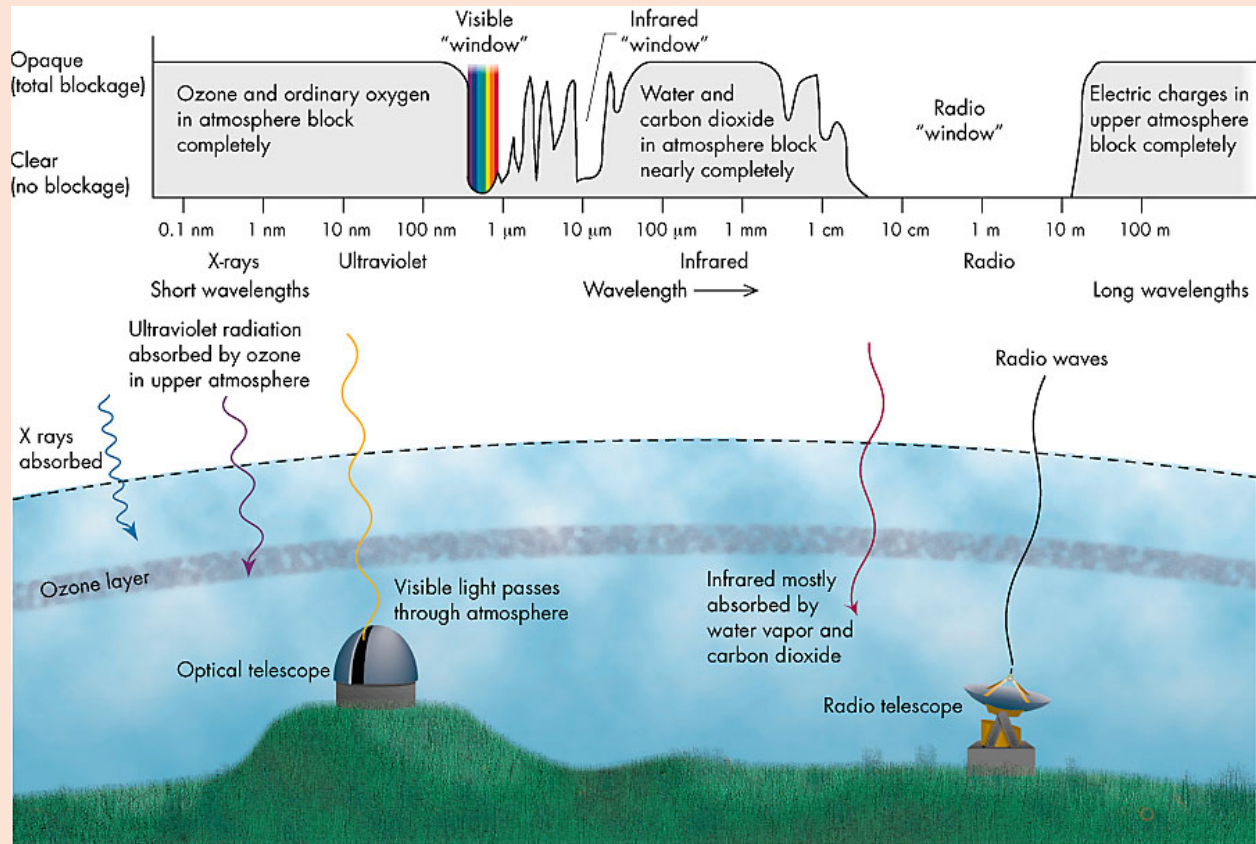


Figure 4.4 Absorption of solar radiation by Earth's atmosphere.

4.4 Energy budget

The study of the science of climate change must start with developing a thorough understanding of the 'global energy budget'. This might be the most complicated part of the science. A simplified but complete version of the global energy budget taken from Kiehl and Trenberth 1997 is shown in Figure 4.5. It shows the 'long term average' values taken over the whole surface of the planet.

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Units: Wm^{-2}

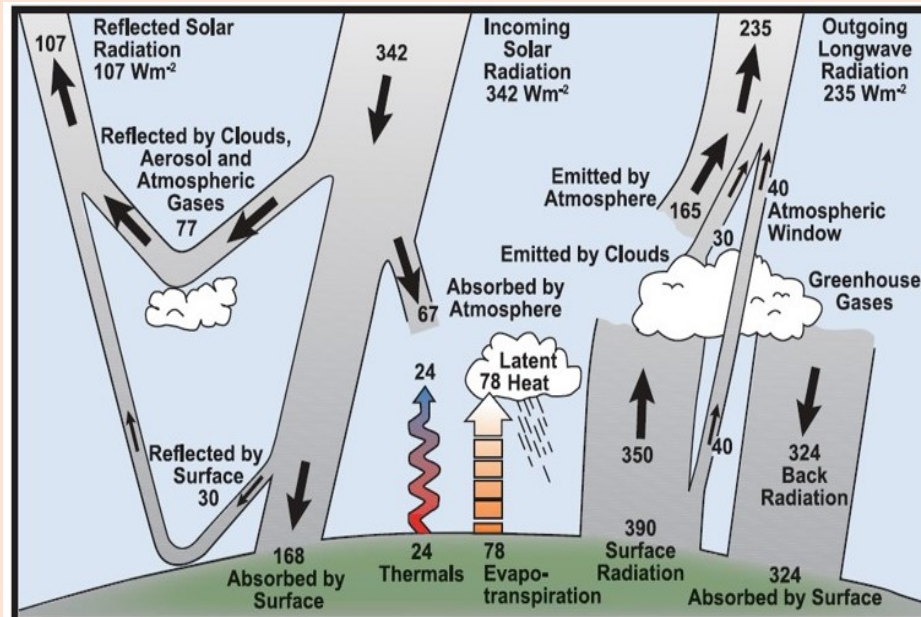


Figure 4.5 Global energy budget 1997 taken from https://journals.ametsoc.org/view/journals/bams/78/2/1520-0477_1997_078_0197_eagmeb_2_0_co_2.xml?tab=body=pdf

Note: The energy budget shown in Figure 4.5 is an early but informative version of the global energy budget. This has been updated several times with the most recent version published in IPCC AR6 Chapter 7 The Earth's Energy Budget, Climate Feedbacks and Climate Sensitivity, Figure 7.2. The IPCC version shows the global energy budgets for 'All Sky' (with clouds) and 'Clear Sky' (without clouds). https://www.ipcc.ch/report/ar6/wg1/downloads/report/IPCC_AR6_WGI_Chapter07.pdf. The 'All Sky' graphic is shown in Figure 4.6. The 'All Sky' figure does not identify the amount of energy which passes through the atmospheric window, the longwave, the heat radiation emitted by the atmosphere to space, or the longwave or heat radiation emitted by clouds. As well, an energy imbalance is shown that is associated with energy being retained by the Earth as a result of the greenhouse effect when clouds and greenhouse gases are present – this is not shown in Figure 4.5.

Units: Wm^{-2}

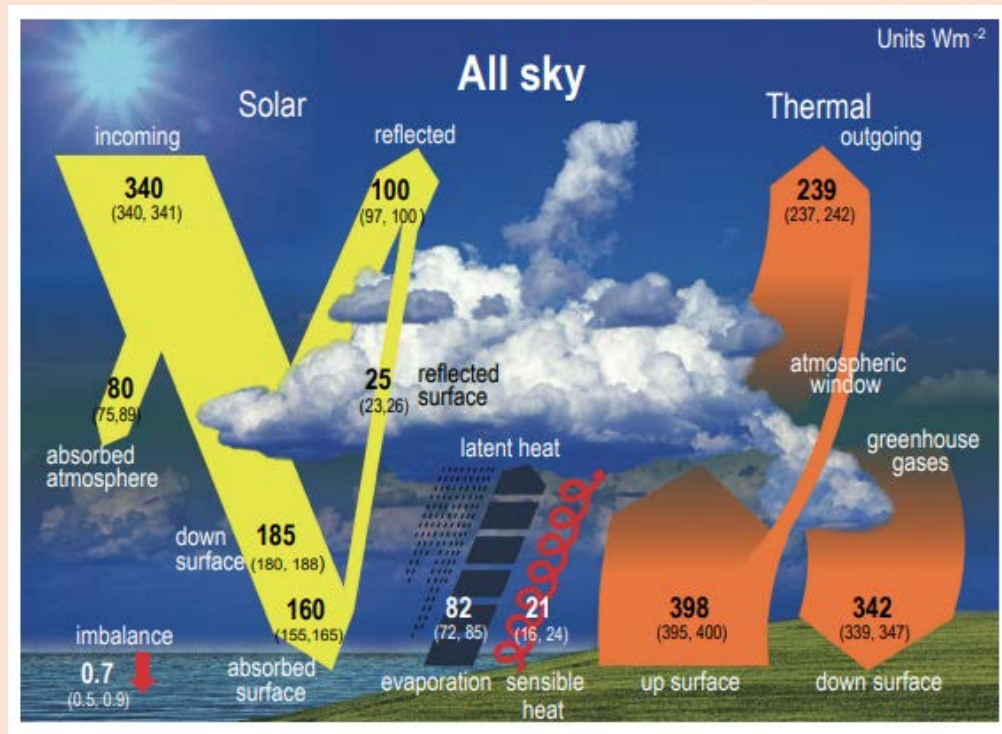


Figure 4.6 Schematic representation of the global mean energy budget of the Earth. Numbers indicate best estimates for magnitudes of the globally averaged energy balance components in W/m^2 together with their uncertainty ranges in parentheses (5-95% confidence range).

https://www.ipcc.ch/report/ar6/wg1/downloads/report/IPCC_AR6_WGI_Chapter07.pdf

Figure 4.7 shows an example of the Earth's energy budget, with the various fluxes indicated as percentage of the incoming solar radiation. Note that the energy imbalance is not shown indicating that the energy balance is stable – Earth is neither warming or cooling. The magnitude of each of the various energy fluxes shown in Figure 4.7 is within the range of those indicated in Figure 4.6.

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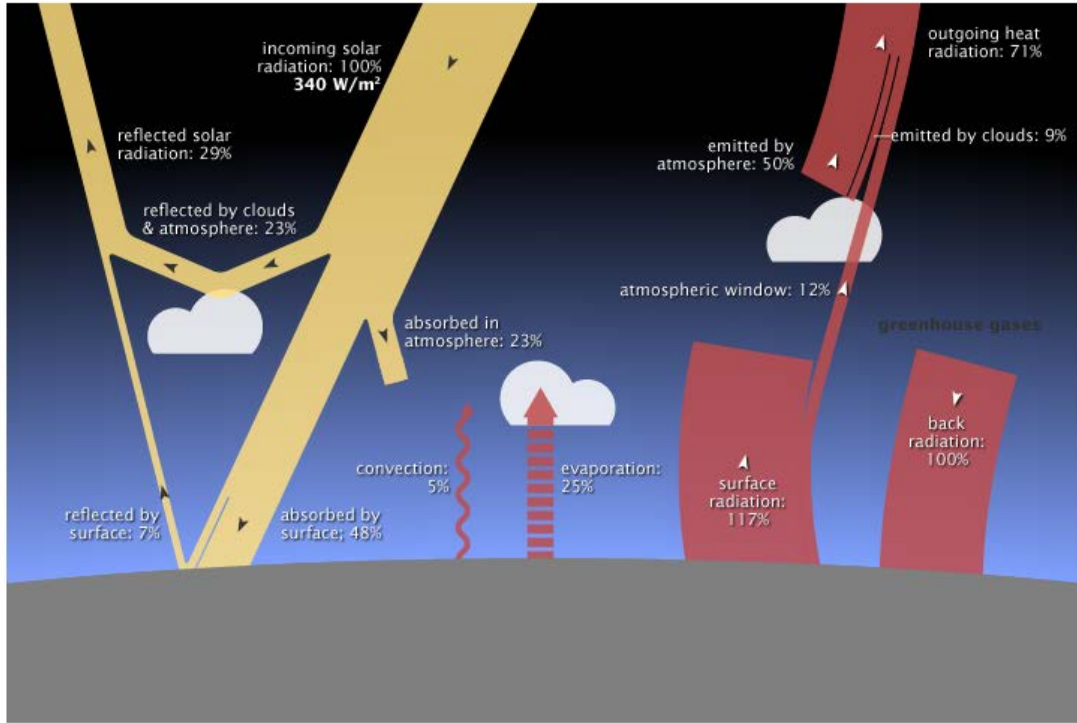


Figure 4.7 Earth's energy budget. <https://earthobservatory.nasa.gov/features/EnergyBalance>

Much of the radiant energy received from the Sun is reflected by the atmosphere (including clouds, aerosols and greenhouse gases) or the surface of the Earth. Some of the radiant energy will be absorbed by the atmosphere and by clouds, aerosols and greenhouse gases. The amount of radiation reflected is unique to the nature of the Earth's surface, presence of clouds, presence of aerosols and greenhouse gases.

When the air immediately above the surface of the Earth is warmed, it may be less dense than the surrounding air and it moves upward (like a hot-air balloon) – thermals or convection. The effect is a transfer of energy from the Earth's surface to the cooler air higher in the atmosphere where the warm air mixes with the cool air and warms it, a process known as sensible heat transfer. Energy transfer from the surface of the Earth by convection cools the Earth and warms the atmosphere.

'Evapotranspiration' refers to the use of energy to convert liquid water to water vapour by evaporation and transpiration by plants. Water vapour and the energy it contains leaves the surface of the Earth as a vapour and becomes part of the atmosphere. The water vapour will warm the atmosphere until it cools sufficiently to condense to form liquid droplets at which time and the remaining energy it contains as a vapour is released equal to the latent heat of condensation (or latent heat of sublimation if turns to ice).

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The amount of energy warming the atmosphere is equal to the sum of the absorption of incoming solar radiation, absorption of surface radiation (up-radiation) that did not pass through the atmospheric window, thermals and the cooling and latent heat of condensation of water vapour. The warmed atmosphere will emit longwave radiation to outer space and to the surface of the Earth as back radiation.

The Earth surface will warm until the energy being absorbed by the surface (absorbed solar radiation plus back radiation) equals the energy lost due to surface radiation, thermals and evapotranspiration.

4.5 Conservation of Energy

The **principle of conservation of energy in a closed system** states that energy cannot be created or destroyed but may change from one form to another. The closed system is Earth, like the inside of the box shown in Figure 4.7. Energy reaching Earth and its atmosphere from the Sun (on top of the atmosphere) **MUST** equal the amount of energy leaving Earth and its atmosphere **plus** the amount of energy stored by Earth and its atmosphere. The energy balance of a closed system is said to be stable if the energy going into the system equals the amount of energy leaving the system which implies that the energy stored within the system is constant. However, much may be happening in terms of how the energy is moving around or being transformed from one type to another **within** the system.

The sides of Figure 4.7 represent the closed system considered when assessing conservation of energy and system stability. **Energy will only move vertically in or out of the system as shown. No energy moves horizontally (sideways).**

The top of Figure 4.7 (one side of our system) shows incoming solar radiation (energy from the sun) is equal to 340 Wm^{-2} . 78.2 Wm^{-2} is reflected by the atmosphere (atmospheric gases, clouds, and aerosols) back into space. 78.2 Wm^{-2} is absorbed by the atmosphere. 28.3 Wm^{-2} is reflected by the surface of the Earth back into space. 163.2 Wm^{-2} is absorbed by the Earth. Note that the water vapour in the atmosphere is not part of the clouds where the water particles, liquid or ice, behave like other aerosols.

Energy that is absorbed by the Earth, 163.2 Wm^{-2} , warms the Earth; and, the energy that is absorbed by the atmosphere, 78.2 Wm^{-2} , warms the atmosphere. The energy from the sun that warms the Earth and the atmosphere is transformed into thermal energy (warm Earth and warm air).

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The total radiant energy absorbed by the Earth equals the sum of absorbed solar radiation and back radiation, 503.2 Wm^{-2} .

The warm Earth loses some of its energy by convection (thermals), 17 Wm^{-2} and by evapotranspiration, 85 Wm^{-2} . The remaining energy that was absorbed by the surface, (solar energy absorbed by surface and back radiation), 401.2 Wm^{-2} , will be emitted to the atmosphere. The atmosphere and clouds will allow some of the long wave radiation to pass through to space without being absorbed, 40.8 Wm^{-2} .

The warm atmosphere and clouds will radiate 340 Wm^{-2} of long wave radiation back to the Earth and 200.6 Wm^{-2} (170 Wm^{-2} plus 30.6 Wm^{-2}) into space.

Consider the energy budget of the Earth's surface. The sum of the energy being absorbed and leaving the surface of the Earth must be equal or the Earth would be steadily warming. The amount of energy being absorbed by the Earth equals portion of solar energy absorbed, 163.2 Wm^{-2} , plus back radiation from the atmosphere, 340 Wm^{-2} , equals 503.2 Wm^{-2} . The amount of energy leaving the surface of the Earth equals the sum of thermals, 17 Wm^{-2} , evapotranspiration, 85 Wm^{-2} , and surface radiation, 397.8 Wm^{-2} , equals 499.8 Wm^{-2} . The surface of the Earth is warming very slightly 3.2 Wm^{-2} – *nearly balanced*.

Consider the energy budget for the atmosphere as an internal system. Energy into atmosphere equals the sum of solar radiation absorbed by atmosphere plus thermals plus evapotranspiration plus long wave radiation from the surface minus energy passing through atmospheric window, $78.2 \text{ Wm}^{-2} + 17 \text{ Wm}^{-2} + 85 \text{ Wm}^{-2} + 397.8 \text{ Wm}^{-2} - 40.8 \text{ Wm}^{-2} = 537.2 \text{ Wm}^{-2}$. Energy leaving the atmosphere equals the sum of that emitted by clouds plus that emitted by atmosphere plus back radiation equal to $30.6 \text{ Wm}^{-2} + 170 \text{ Wm}^{-2} + 340 \text{ Wm}^{-2} = 540 \text{ Wm}^{-2}$. The atmosphere is cooling very slightly 2.8 Wm^{-2} - *nearly balanced*.

Consider the global energy budget for the Earth. The sum of the solar radiation, 98.6 Wm^{-2} , plus the amount of out going long wave radiation, 241.4 Wm^{-2} , equals the amount of incoming solar radiation, 340 Wm^{-2} . The energy budget of the Earth is stable; that is, it is neither warming or cooling.

4.6 Greenhouse effect

The 'greenhouse effect' refers to the increased warming of the Earth's surface when an atmosphere (including water vapour, carbon dioxide and aerosols) and clouds are present. It is very difficult to predict the temperature of the Earth without an atmosphere because this condition also implies that there isn't any water or vegetation on Earth. If it is assumed that the reflectivity of Earth under these conditions is the same as our moon, the amount of the solar energy reflected is 12%, (geometric albedo <https://earthobservatory.nasa.gov/features/EnergyBalance>). This means that 88% of the energy reaching the Earth will be absorbed and remitted to space; that is, $0.88 \times 340 \text{ Wm}^{-2} = 299.2 \text{ Wm}^{-2}$.

The amount of energy emitted by the surface of the Earth when there is an atmosphere (including water vapour, carbon dioxide and aerosols) and clouds are present is 397.8 W/ m^2 .

The Stephan-Boltzmann Equation or S-B Equation, (<http://hyperphysics.phy-astr.gsu.edu/hbase/thermo/stefan.html>) can be used to estimate the amount of radiation leaving a solid, in this case the Earth. The equation is:

$$P/A = eT^4 / \sigma \quad \text{or} \quad T = (P/A \times \sigma / e)^{-1/4}$$

Where: P = energy in Watts (W)

A = area in m^2 from which energy is emitted

e = emissivity of mass (1 for a black body)

T = temperature of mass in degrees Kelvin or $^{\circ}\text{K}$ (-273 degrees Centigrade)

σ = Stephan-Boltzmann constant equal to $5.6703 \times 10^{-8} \text{ W/m}^2/^{\circ}\text{K}^4$

Using the S-B Equation and assuming it is emitting like a black body, the temperature of the surface of the Earth with an atmosphere is given by:

$$T_{\text{with atmosphere}} = (397.8 \text{ W/ m}^2 / 1 / 5.6703 \times 10^{-8} \text{ W/m}^2/^{\circ}\text{K}^4)^{-1/4} = 295.5 \text{ }^{\circ}\text{K} \text{ or}$$

$$T_{\text{with atmosphere}} = \mathbf{16.4^{\circ}\text{C}}$$

The temperature of the surface of the Earth without an atmosphere is given by:

$$T_{\text{without atmosphere}} = (299.2 \text{ W/ m}^2 / 1 / 5.6703 \times 10^{-8} \text{ W/m}^2/^{\circ}\text{K}^4)^{-1/4} = 269.5 \text{ }^{\circ}\text{K} \text{ or}$$

$$T_{\text{without atmosphere}} = \mathbf{-2.7^{\circ}\text{C}}$$

Estimates by others are somewhat similar.

Earth would not be habitable without an atmosphere and the greenhouse effect.

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4.7 Impact of human activities on the energy budget.

It is very important to remember that the global energy budget as shown in Figures 4.6 and 4.7 are the 'long term average' for the whole planet. The energy budget does in fact vary considerably locally and regionally and with time.

The most significant human activities are identified in Figure 4.8. Human activities include surface modification and changes in concentration of greenhouse gases (including aerosols) in the atmosphere.

The effects of modifying the Earth surface might be to alter the ability to absorb solar and longwave radiation, reflect solar radiation, emit radiation to the atmosphere, and altering the manner with which the processes of evapotranspiration and advective (thermal) transfer of energy occur.

The addition of greenhouse gases will affect the reflection of solar radiation (aerosols and clouds), and the absorption and emission of radiation (all greenhouse gases including water vapour and aerosols). While the amounts of reflected solar radiation and the amount of longwave radiation leaving the Earth's atmosphere may vary, their sum will always equal the amount of incoming solar radiation.

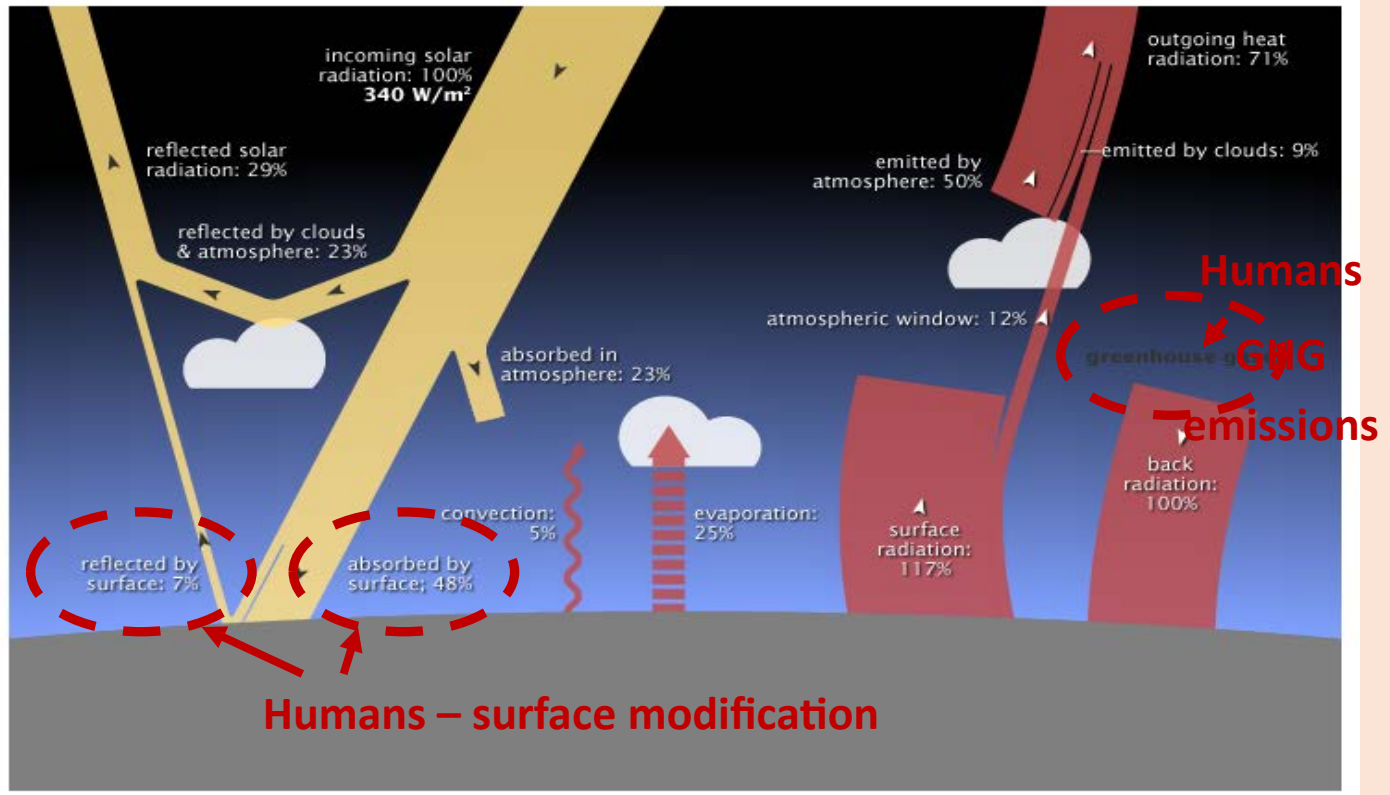


Figure 4.8. Components of energy budget circled in dashed red are directly affected by human activities.

NASA, in their November, 2021 newsletter, discusses climate change <https://climate.nasa.gov/causes/>. It includes a discussion of the greenhouse effect, its cause, and some of the impacts.

Syukuro Manabe and Klaus Hasselmann who jointly win one-half of the Nobel Prize in Physics in 2021 for their pioneering work on the “physical modelling of Earth’s climate, quantifying variability and reliably predicting global warming”. It is well worth reviewing. See <https://www.nobelprize.org/prizes/physics/2021/manabe/facts/>

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4.6 Information support

Key web sites:

1. What is energy. <https://www.eia.gov/energyexplained/what-is-energy/>
2. Energy defined. <https://www.britannica.com/science/energy>
3. Forms of energy. https://en.wikipedia.org/wiki/Outline_of_energy
4. Overview of climate system.
<https://www.ipcc.ch/site/assets/uploads/2018/03/TAR-01.pdf>
5. Earth's energy budget. https://en.wikipedia.org/wiki/Earth%27s_energy_budget
6. Earth's annual global mean energy budget.
https://journals.ametsoc.org/view/journals/bams/78/2/1520-0477_1997_078_0197_eagmeb_2_0_co_2.xml?tab_body=pdf
7. Earth's annual global mean energy budget. IPCC AR6 Chapter 7.
https://www.ipcc.ch/report/ar6/wg1/downloads/report/IPCC_AR6_WGI_Chapter07.pdf
8. Sunlight. <https://en.wikipedia.org/wiki/Sunlight>
9. Solar radiation and photosynthetically active radiation.
<https://www.fondriest.com/environmental-measurements/parameters/weather/photosynthetically-active-radiation/>
10. Energy education. <https://energyeducation.ca/encyclopedia/Insolation>
11. Solar energy to the Earth.
https://energyeducation.ca/encyclopedia/Solar_energy_to_the_Earth
12. Schematic representation of the global mean energy budget of the Earth.
https://www.ipcc.ch/report/ar6/wg1/downloads/report/IPCC_AR6_WGI_Chapter07.pdf
13. Schematic representation of the global mean energy budget of the Earth.
<https://earthobservatory.nasa.gov/features/EnergyBalance>
14. Black-body radiation. https://en.wikipedia.org/wiki/Black-body_radiation

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15. Greenhouse gases. https://en.wikipedia.org/wiki/Greenhouse_gas#Atmospheric_lifetime
16. Greenhouse effect. https://en.wikipedia.org/wiki/Greenhouse_effect
17. Greenhouse effect. <https://www.nrdc.org/stories/greenhouse-effect-101>
18. Moon fact sheet, NASA. <https://nssdc.gsfc.nasa.gov/planetary/factsheet/moonfact.html>
19. Stephen-Boltzmann Equation. <http://hyperphysics.phy-astr.gsu.edu/hbase/thermo/stefan.html>
20. The causes of climate change, NASA, November 2021. <https://climate.nasa.gov/causes/>.
21. Syukuro Manabe, Nobel Prize in Physics 2021. <https://www.nobelprize.org/prizes/physics/2021/manabe/facts/>

Videos:

1. Greenhouse effect. <https://www.youtube.com/watch?v=SN5-DnOHQmE>

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