See discussions, stats, and author profiles for this publication at: https://www.researchgate.net/publication/322800632

# Intraspecific variation in the response of the scleractinian coral Acropora digitifera to ocean acidification

Article *in* Marine Biology · January 2018 DOI: 10.1007/s00227-018-3295-1

citations 14			READS 332			
4 authors, including:						
	Haruko Kurihara Jniversity of the Ryukyus O PUBLICATIONS 3,317 CITATIONS SEE PROFILE	•	Alejandro Reyes-Bermudez Universidad de la Amazonia 56 PUBLICATIONS 319 CITATIONS SEE PROFILE			

All content following this page was uploaded by Alejandro Reyes-Bermudez on 01 February 2018.

#### **ORIGINAL PAPER**



# Intraspecific variation in the response of the scleractinian coral *Acropora digitifera* to ocean acidification

Haruko Kurihara<sup>1</sup> · Asami Takahashi<sup>1</sup> · Alejandro Reyes-Bermudez<sup>1,2</sup> · Michio Hidaka<sup>1</sup>

Received: 13 August 2017 / Accepted: 16 January 2018 © Springer-Verlag GmbH Germany, part of Springer Nature 2018

#### Abstract

To examine the possible variation in responses of corals to ocean acidification (OA) among populations, we compared the sensitivity of two Okinawan populations (Sesoko and Bise) of the scleractinian coral *Acropora digitifera* to high  $pCO_2$ . We found that both light and dark calcification rates of Sesoko corals did not change with an increase in seawater  $pCO_2$ , while the calcification rates of Bise corals significantly decreased. Additionally, calcification rate of Sesoko corals was significantly lower than Bise corals at control conditions. Expressions of two putative calcification-related genes (BAT: bicarbonate transporter and galaxin) were up-regulated at high  $CO_2$  compared to the control and expression of the BAT gene was significantly higher in Sesoko compared to Bise corals. Consequently, differences in the calcification rate between populations and differences in the expression of genes related to inorganic carbon transport regulation could be reasons that explain the difference in the response to OA between the two populations. Furthermore, taking into account that Sesoko corals were located in relatively nearshore areas where the environmental conditions are more variable, while Bise corals were located in the forereef which shows more stable conditions, plasticity for coral calcification in response to different environmental conditions and/or acclimation response to changes such as seawater  $pCO_2$  may lead to differences in sensitivity between the two populations to high seawater  $pCO_2$ . Studies considering the potential variability in corals sensitivity to OA among local populations from different habitats are important to predict the potential effects of climate change on reef ecosystems.

# Introduction

Ocean acidification (OA) caused by the increasing atmospheric  $CO_2$  is predicted to pose a significant threat to scleractinian corals and reef ecosystems (Kleypas et al. 2006; Hoegh-Guldberg et al. 2007). When seawater pH decreases, the concentration of carbonate ions also declines, resulting

Responsible Editor: S. Uthicke.	
Reviewed by undisclosed experts.	

**Electronic supplementary material** The online version of this article (https://doi.org/10.1007/s00227-018-3295-1) contains supplementary material, which is available to authorized users.

Haruko Kurihara harukoku@sci.u-ryukyu.ac.jp; harukoku@e-mail.jp

<sup>1</sup> Department of Chemistry, Biology, and Marine Science, Faculty of Science, University of the Ryukyus, 1 Senbaru, Nishihara, Okinawa 903-0213, Japan

<sup>2</sup> School of Biology, Faculty of Natural Sciences, Universidad de la Amazonia, Florencia, Caquetá, Colombia

Published online: 30 January 2018

in a reduction in seawater aragonite saturation ( $\Omega_{arag}$ ). From field observations, it has been reported that net reef calcification shows a strong correlation with  $\Omega_{arag}$ , and a number of experimental studies have demonstrated that net calcification rates of most corals decline with  $\Omega_{arag}$  (Langdon and Atkinson 2005; Chan and Connolly 2013). Presently, the negative effect on calcification is one of the most robust predictions regarding the potential impact of OA on marine organisms. However, recent studies have revealed that some coral species can be relatively more tolerant than other species to OA (Jokiel et al. 2008; Jury et al. 2010; De Putron et al. 2011), suggesting interspecific variations in the response of corals to OA (Pandolfi et al. 2011; Comeau et al. 2014a).

Identification of the interspecific variations in the response to OA is essential information for evaluating how OA will affect organisms at community levels. If the tolerance capacity differs among species, changes in community structure can be predicted, tending to become dominated by species that show higher tolerance to OA. A tank experiment study evaluating the sensitivity of eight different functional coral species to OA demonstrated that the responses significantly differed among species (Comeau et al. 2014a). Additionally, studies of  $CO_2$ -vent associated assemblages in reef ecosystems have demonstrated that these coral communities can dramatically change to potentially more OA-tolerant coral species such as massive *Porites* or soft-corals, resulting in the loss of biodiversity (Fabricius et al. 2013; Inoue et al. 2013).

Meanwhile, identification of intraspecific variations in response to OA could give useful information for evaluating the adaptation or acclimation potential of individual species to OA (Sunday et al. 2011). Although studies evaluating intraspecific variations are still limited, some studies have demonstrated that sensitivity to OA could differ among populations within the same species. For example, Parker et al. (2011) demonstrated that selectively bred populations of the oyster Crassostrea gigas are more resilient to elevated  $pCO_2$  than wild populations. Additionally, a blue mussel Mytilus edulis population living in a Fjord where seawater  $pCO_2$  highly fluctuates in response to upwelling can maintain similar shell growth rate to controls when reared at high-CO<sub>2</sub> (1400 µatm) seawater (Thomsen et al. 2010). Differences in the response to OA among different local populations were also documented in the mussel Mytilus galloprovincialis on the Mediterranean Sea coast (Range et al. 2014). These results suggest that the OA sensitivity of the organisms can differ among local populations experiencing different environmental conditions. For corals, one study has evaluated the possible differences in the resistance of the corals *Pocillopora* damicornis and massive Porites resistance to OA between populations from different regions (Okinawa, Moorea and Hawaii); however, both coral species from all locations were insensitive to OA (Comeau et al. 2014b). Another study using a common garden experiment in the coral Acropora pulchra in Moorea, French Polynesia, revealed intraspecific variation on calcification rate to elevated temperature but not to OA (Shaw et al. 2016). Meanwhile, studies that have investigated the effect of OA among colonies of the coral Acropora digitifera, Montipora digitata and *Porites cylindrica* indicated that the response of corals to OA could vary among individuals within the same population (Kavousi et al. 2015, 2016). Additionally, in our previous study, we identified that the coral Acropora digitifera in Okinawa did not show a significant correlation between calcification rate and seawater  $\Omega_{\rm arag}$  (Takahashi and Kurihara 2013), suggesting that this species is potentially resilient to OA, which differed from other studies suggesting high vulnerability of fast-growing corals such as Acropora spp. (e.g. Comeau et al. 2014a, Shaw et al. 2016). However, when the same experiment was repeated using A. digitifera collected from another location, we found a strong correlation between calcification rate and seawater  $\Omega_{arag}$ . Consequently, we hypothesized that the OA

sensitivity of the coral A. *digitifera* may show variations among different sites.

Here, to examine the possible variation in responses of corals to OA among populations, we collected the coral *A*. *digitifera* from two locations around Okinawa, one collection site being the same location where *A*. *digitifera* were collected for the previous study (Takahashi and Kurihara 2013), and evaluated the effect of high  $pCO_2$  on coral net calcification. Effect of high  $pCO_2$  on the rate of light and dark calcification rate, photosynthesis rate, respiration rate, photosynthesis efficiency, zooxanthellae density, and protein content was compared between the corals from the two sites. In addition, we examined the gene expression of four genes with putative roles in the calcification process of corals to evaluate the potential mechanism for OA sensitivity variation between sites.

# **Materials and methods**

## **Coral collection**

Five colonies of the coral Acropora digitifera were collected in June 2011 from each of two locations around Okinawa, Japan: inshore patch reefs (ca. 0.5-1.0 m depth) at Sesoko (26°38'11.64"N, 127°51'57.08"E) and at edge of reef crest towards to the forereef zone (ca. 0.5-1.0 m depth) at Bise (26°42'40.0"N, 127°52'53.3"E). The distance between the two Bise and Sesoko sites was less than 8 km. Light sensor (DEFI2-L, JFE-Advantech, Japan) and multi-parameter water quality sonde (6820 V2, YSI, USA) were placed at the two sites for 3 days to measure light intensity, seawater temperature, dissolved oxygen, salinity, and pH (calibrated with NBS scale buffer). Several seawater samples were also collected for total alkalinity (TA) and dissolved inorganic carbon (DIC) measurement during the 3 days. Three pieces with three branches from each of the five colonies were collected 4 weeks prior to the experiment. After sampling, all coral pieces were immediately transferred to the Sesoko Marine Laboratory Station of University of the Ryukyus and placed for 2 weeks in an outdoor tank  $(1.2 \text{ m} \times 0.7 \text{ m} \times 0.2 \text{ m}, \text{ vol-}$ ume = 187 L) under natural light condition and supplied continuously with seawater  $(1.5 \text{ Lmin}^{-1})$  pumped from 4-5 m depth in front of the station. Thereafter, all coral pieces were placed individually in incubation chambers (volume = 450 mL), which were kept in an indoor water bath supplied continuously with seawater pumped from the front of the station to control temperature. The corals were set under artificial light provided by a 400 W metal halide lamp (light intensity: 200  $\mu$ mol photons m<sup>-2</sup> s<sup>-1</sup> L:D = 14 h:10 h). The light intensity was slightly lower but within the range of the mean June daytime light intensity measured at Sesoko (mean light intensity: 280 µmol photons m<sup>-2</sup> s<sup>-1</sup>, range 70–430  $\mu$ mol photons m<sup>-2</sup> s<sup>-1</sup>) and cultured for another 2 weeks for acclimation.

#### **Experimental setup**

One piece with three branches from each of the five colonies and two locations (Sesoko and Bise) was incubated individually under control (420  $\pm$  40 µatm; mean  $\pm$  SD), mid-CO<sub>2</sub>  $(1030 \pm 48 \ \mu atm)$  and high-CO<sub>2</sub>  $(2381 \pm 156 \ \mu atm)$  conditions for 5 weeks. Seawater was bubbled in a bubbling tank (18 L) using air, or air mixed with CO<sub>2</sub> controlled by mass flow controllers (SEC-E40, Horiba Stec, Japan), and supplied to each incubator chamber at a flow rate of ca. 150 mL min<sup>-1</sup>. Seawater pH and temperature of the bubbling tank were measured daily using a temperature-compensated pH electrode (MP125, Mettler Toledo, USA) calibrated using NBS scale buffer solutions. Temperature and pH of each bubbling tanks were verified as being same to the downstream incubation chambers. Salinity  $(34 \pm 0)$  and total alkalinity (TA) were measured every week using a refractometer (100-S, Atago, Japan) and autoburette titrator (ATT-05, Kimoto, Japan), respectively. Seawater  $pCO_2$ ,  $HCO_3^-$ ,  $\text{CO}_3^{2-}$  and  $\Omega_{\text{arag}}$  were calculated based on pH, temperature, TA and salinity using the CO2SYS program of Lewis and Wallace (1998), with dissociation constants  $K_1$  and  $K_2$  from Mehrbach et al. (1973) and an aragonite solubility of Mucci (1983) (Table 1).

#### Coral physiology

The net calcification rates (*G*) of all coral pieces were measured by the buoyant weight technique (Davis 1989) using an electronic balance (0.1 mg precision HR-200, A&D, Japan) before and after the 5 weeks of experimental culture. Skeletal dry weight was calculated based on aragonite density (2.94 g cm<sup>-3</sup>) and the increase in the skeletal dry weight was normalized to the surface area measured by covering the surface of the coral pieces with aluminum foil and the area of the foil was analyzed using image J (Image J, U.S.).

Light calcification (light *G*), dark calcification (dark *G*), net photosynthesis (nP) and dark respiration (*R*) rates of all coral pieces were measured by the dissolved inorganic carbon (DIC)–total alkalinity (TA) technique (Smith 1973; Smith and Key 1975) at the end of the 5-week incubations.

Light G and nP were measured during the daytime under light conditions (light intensity: 200  $\mu$ mol photons m<sup>-2</sup> s<sup>-1</sup>) and dark G and R during the nighttime under dark conditions. Initial seawater samples (ca. 100 mL) were collected from each airtight glass chamber (450 mL), and immediately after the chambers were re-filled with new flowing seawater, all chambers were sealed without a head space. All experimental chambers were well mixed with a magnetic stirrer (450 rpm) during culture, and seawater was sampled at the end of 2 h culture. The seawater was poisoned by adding 100 µl saturated mercuric chloride immediately after sampling. DIC and TA were measured using closed cell titration (ATT-05, Kimoto, Japan) using 0.1 N HCl. Accuracy and precision of the determinations were evaluated by analyzing reference material Batch AG (TA =  $2295 \pm 0.55 \,\mu\text{mol kg}^{-1}$ , DIC =  $2032.8 \pm 0.72 \ \mu mol \ kg^{-1}$ , from Kanso Technos, Japan) which was verified against certified reference material (CRM, Batch 95, from A. Dickson Laboratory). Light and dark G were calculated using the following equation:

Light, dark G (µmol CaCO<sub>3</sub> cm<sup>-2</sup> h<sup>-1</sup>)

 $= \Delta TA \times V \times D_{sw} (2 \times SA \times t)^{-1},$ 

where  $\Delta$ TA is the change in TA measured at the beginning and end of the day and night time incubations, *V* is the seawater volume of each chamber minus the displacement volume of the coral piece,  $D_{sw}$  is seawater density, SA is the surface area of each coral piece, and t is the duration of culture. nP and *R* were calculated by the organic production rate during day and night time incubations and calculated by the following equation:

nP,  $R \ (\mu \text{mol C cm}^{-2}\text{h}^{-1}) = (\Delta \text{DIC} - 0.5\Delta \text{TA}) \times V \times D_{\text{sw}}(\text{SA} \times t)^{-1}$ , where  $\Delta \text{DIC}$  and  $\Delta \text{TA}$  for nP and R are the DIC and TA change during the light and dark incubations, respectively. Gross photosynthesis (gP) was calculated by summing nP and R.

# Photosynthetic efficiency, zooxanthellae density, and lipid content

Photochemical efficiency of photosynthesis II  $(F_v F_m^{-1})$  of zooxanthellae was measured using a pulse amplitude

Table 1Seawater carbonchemistry during the 5 weeks ofcoral culture

	$pCO_2$ (µatm)	pH (NBS scale)	TA (µmol/kg)	DIC (µmol/kg)	$arOmega_{ m arag}$	Temperature
Control	$420 \pm 40$	$8.16 \pm 0.03$	2236 ± 22	1931 ± 22	$3.51 \pm 0.25$	28.9 ± 0.9
Mid-CO <sub>2</sub>	$1027 \pm 48$	$7.83 \pm 0.02$	2233 ± 33	$2091 \pm 8$	$1.87\pm0.08$	$28.9\pm0.9$
High-CO <sub>2</sub>	$2374 \pm 155$	$7.50 \pm 0.03$	2227 ± 27	$2205 \pm 8$	$0.93 \pm 0.05$	$28.9\pm0.9$

 $pCO_2$ , dissolved inorganic carbon (DIC) and  $\Omega_{arag}$  were calculated with CO2SYS program by daily measured pH, temperature and weekly measured total alkalinity (TA) and salinity. Mean  $\pm$  SD

fluorescence system (MINI-PAM, Walz Effeltrich, Germany) according to Schreiber et al. (1986). Maximum fluorescence of each coral piece was measured by a 0.8 s saturation light pulse (8000  $\mu$ mol photon m<sup>-2</sup> s<sup>-1</sup>). Measurements were conducted at night under dark conditions after the 5 weeks of incubation.

After all physiological measurements, one branch from each coral piece was cut and the coral tissue was removed with a waterpik and homogenized with a hand-held homogenizer in filtered seawater. The sample was centrifuged (4 min at 2400 rpm), and after removing the supernatant, the zooxanthellae pellet was re-suspended with filtered seawater and the number of zooxanthellae was determined by counting four replicates using a hematocytometer and standardized against surface area.

For analysis of the total lipid amount, extraction was performed according to the method by Yamashiro et al. (1999). A second branch was cut from each coral piece, freeze-dried and then transferred into a chloroform–methanol solution (CM, 2:1 volume) and kept at 40 °C for 30 min to extract lipids. Extraction was conducted three times for each branch, the CM extracts were centrifuged, and the supernatant was dried up and weighed using an electronic balance. After the lipid extraction, the coral piece was decalcified in 10% formaldehyde containing 20% acetic acid and then dried. The dry tissue was weighed and lipid content (%) was calculated as lipid weight divided by tissue dry weight.

#### **Gene expression**

The third branch from each coral piece was frozen with liquid nitrogen and kept frozen in the deep freezer at - 80 °C until molecular analysis. Total RNA from frozen coral tissues was isolated using Qiazol lysis reagent (Qiagen) following product specifications. Coral branches were homogenized, and frozen coral powder was transferred to Qiazol lysis reagent (Qiagen) for RNA extraction. RNA pellets were re-dissolved in nuclease-free water and cleaned further with RNeasy Mini columns (Qiagen). RNA quantity and integrity were assessed with a NanoDrop ND-1000 spectrophotometer and an Agilent 2100 Bioanalyzer, respectively. cDNA was synthesized from 300 ng of DNase-treated RNA samples using the SuperScript III system for RT-PCR (Invitrogen). For the gene expression analysis, only samples of control and high-CO<sub>2</sub> condition were used.

Here, we selected four *Acropora digitifera* genes with putative roles in calcification as molecular markers: a plasma membrane calcium-ATPase (PMCA)(aug\_v2a.22365.t1), a bicarbonate transporter (BAT)(aug\_v2a.09009.t1), a carbonic anhydrase II (CAII)(aug\_v2a.24568.t1) and the structural organic matrix protein galaxin (aug\_v2a.18631. t1) (Grasso et al. 2008; Reyes-Bermudez et al. 2009; Hayward et al. 2011; Zoccola et al. 2004). The specific forward

and reverse primers for the genes were designed (100 bp amplicon length, 55% G–C content and  $T_{\rm m}$  of 65 °C) using primer-express®v3 software (Applied Biosciences). The GAPDH primers used in this experiment were previously used and designed for A. millepora by Souter et al. (2010) and were able to amplify the corresponding A. digitifera GAPDH ortholog. Amplicons were checked by conventional gel electrophoresis and primer efficiencies were determined using standard curve analysis with a twofold dilution series from pooled control and treated cDNA samples. Primer efficiencies ranged between 1.9 and 2.2. Primer sequences can be found in Table S1 (Supplementary information). Quantitative real-time PCR was performed on an ABI-7500 platform (Applied Biosciences) using SYBR green chemistry. Briefly, samples were run in triplicate in 20 µl reactions consisting of 10 µl SYBR green mix, 15 ng of template and 500 nM of each primer. Thermo-cycling conditions included a 3 min initial activation stage at 95 °C followed by 40 cycles of 15 s at 95 °C, 1 min at 60 °C and a final disassociation step (60-95 °C) to confirm the absence of either non-specific products or primer dimers. Threshold cycle differences ( $\Delta C_t$ ) were obtained by manually locating the threshold immediately above the background in the exponential amplification phase.  $C_t$  values were analyzed by relative quantification using the Pfaffl (2001) method.

#### **Statistical analysis**

Effect of CO<sub>2</sub> on net calcification, light and dark calcification, photosynthesis, respiration,  $F_v F_m^{-1}$ , zooxanthella density, lipid content and expression of the four studied genes between sites were evaluated using mixed-effect-model ANOVA with CO<sub>2</sub> and sites as fixed effects and colony as a random effect based on the restricted maximum likelihood (REML) method that automatically corrected variances across interaction effects. Model assumptions of normality and homoscedasticity were checked by examining model residuals. Differences between CO<sub>2</sub> and sites were assessed using Tukey's honestly significant difference (HSD) multiple comparison tests. All analyses were performed using JMP (JMP 7; SAS Inc.)

#### Results

#### **Coral physiology**

The net calcification rate (*G*) of *Acropora digitifera* over the 5-week incubation was significantly different among  $pCO_2$  conditions but not between Sesoko and Bise corals (REML,  $CO_2$ :  $F_{(2,16)} = 15.9$ , p = 0.0002, locations:  $F_{(1,8)} = 2.03$ , p = 0.19) without interaction ( $CO_2 \times$  location:  $F_{(2,16)} = 1.93$ , p = 0.17, Fig. 1). The calcification rate of control Bise corals



**Fig. 1** Percentage change in skeletal dry weight of the corals from Sesoko and Bise during 5 weeks of culture under three different  $CO_2$  conditions. Mean  $\pm$  SE of five colonies. Significant differences among  $CO_2$  conditions and corals from the two sites are indicated by lower case letters (Tukey HSD test p < 0.05)

(16.2 ± 1.9 µmol CaCO<sub>3</sub> cm<sup>-2</sup> day<sup>-1</sup>) was significantly higher compared to mid- (10.7 ± 0.6 µmol CaCO<sub>3</sub> cm<sup>-2</sup> d ay<sup>-1</sup>), and high-CO<sub>2</sub> corals (9.6 ± 0.9 µmol CaCO<sub>3</sub> cm<sup>-2</sup> day<sup>-1</sup>, p < 0.05, Tukey HSD), while the calcification rate of Sesoko corals did not differ among all three CO<sub>2</sub> conditions (control = 11.4 ± 1.0 µmol CaCO<sub>3</sub> cm<sup>-2</sup> d<sup>-1</sup>, mid-CO<sub>2</sub> = 8.9 ± 0.9 µmol CaCO<sub>3</sub> cm<sup>-2</sup> day<sup>-1</sup> and high CO<sub>2</sub> = 9.0 ± 0.9 µmol CaCO<sub>3</sub> cm<sup>-2</sup> day<sup>-1</sup>, Tukey HSD, Fig. 1).

The light calcification rate (light G) was significantly different among  $pCO_2$  conditions, but not between locations (REML, CO<sub>2</sub>:  $F_{(2,16)} = 19.7$ , p < 0.0001, locations:  $F_{(1.8)} = 0.66, p = 0.44$ ) and there was a significant interaction (CO<sub>2</sub> × location:  $F_{(2,16)} = 75.25$ , p = 0.01). Light G of control Bise corals  $(1.32 \pm 0.42 \,\mu\text{mol CaCO}_3 \,\text{cm}^{-2} \,\text{h}^{-1})$  was significantly higher than Sesoko corals  $(0.65 \pm 0.03 \mu mol C)$  $aCO_3 \text{ cm}^{-2} \text{ h}^{-1}$ , p < 0.05, Tukey HSD, Fig. 2), but there was no significant difference at high-CO<sub>2</sub> between Bise (0.54  $\pm$ 0.21  $\mu$ mol CaCO<sub>3</sub> cm<sup>-2</sup> h<sup>-1</sup>) and Sesoko corals (0.64  $\pm$  0. 15  $\mu$ mol CaCO<sub>3</sub> cm<sup>-2</sup> h<sup>-1</sup>). Additionally, light G of control Bise corals was significantly higher than mid-CO<sub>2</sub> (0.90  $\pm$ 0.19  $\mu$ mol CaCO<sub>3</sub> cm<sup>-2</sup> h<sup>-1</sup>) and high-CO<sub>2</sub> (0.54  $\pm$  0.21  $\mu$ mol CaCO<sub>3</sub> cm<sup>-2</sup> h<sup>-1</sup>, p < 0.05, Tukey HSD, Fig. 2), while there was no significant difference of light G of Sesoko corals among all CO<sub>2</sub> conditions.

Dark calcification rate (dark *G*) was significantly different among CO<sub>2</sub> conditions and between Sesoko and Bise corals (REML, CO<sub>2</sub>:  $F_{(2,16)} = 40.3$ , p < 0.0001, locations:  $F_{(1,8)} = 11.8$ , p = 0.0088) with interaction (CO<sub>2</sub> × location:  $F_{(2,16)} = 37.8$ , p < 0.0001). Dark *G* of control (0.79 ± 0.1 5 µmol CaCO<sub>3</sub> cm<sup>-2</sup> h<sup>-1</sup>) and mid-CO<sub>2</sub> (0.72 ± 0.10 µmol CaCO<sub>3</sub> cm<sup>-2</sup> h<sup>-1</sup>) Bise corals were significantly higher than Sesoko control (0.31 ± 0.08 µmol CaCO<sub>3</sub> cm<sup>-2</sup> h<sup>-1</sup>) corals and mid-CO<sub>2</sub> (0.22 ± 0.06 µmol CaCO<sub>3</sub> cm<sup>-2</sup> h<sup>-1</sup>) corals



**Fig. 2** Light and dark calcification of the corals from Sesoko and Bise after 5 weeks of culture under three different CO<sub>2</sub> conditions. Mean  $\pm$  SE of five colonies. Significant differences among CO<sub>2</sub> conditions and corals from the two sites, are indicated by lower case letters (Tukey HSD test *p* < 0.05)

(p < 0.05, Tukey HSD), respectively, while there was no significant difference between high-CO<sub>2</sub> Bise  $(0.20 \pm 0.15 \mu mol CaCO_3 cm^{-2} h^{-1})$  and Sesoko  $(0.40 \pm 0.08 \mu mol CaCO_3 cm^{-2} h^{-1})$  corals (Fig. 2). Additionally, dark *G* of control Bise corals was significantly higher than both mid-CO<sub>2</sub> and high-CO<sub>2</sub> corals, while there was no significant effect of CO<sub>2</sub> on dark *G* of Sesoko corals.

The gross photosynthesis rate (gP) was significantly different among CO<sub>2</sub> conditions but not between Sesoko and Bise corals (REML, CO<sub>2</sub>:  $F_{(2,16)} = 34.3$ , p < 0.0001, locations:  $F_{(1,16)} = 0.08$ , p = 0.78) without interaction (CO<sub>2</sub> × location:  $F_{(2,16)} = 0.16$ , p = 0.85). gP of control corals (Bise:  $3.24 \pm 0.46 \mu$ mol C cm<sup>-2</sup> h<sup>-1</sup>, Sesoko:  $3.19 \pm 0.26 \mu$ mol C cm<sup>-2</sup> h<sup>-1</sup>) was significantly higher compared to mid- and high-CO<sub>2</sub> corals in both Bise and Sesoko corals (p < 0.05, Tukey HSD, Fig. 3).

The respiration rate (*R*) was significantly different among CO<sub>2</sub> conditions (REML, CO<sub>2</sub>:  $F_{(2,16)} = 21.0$ , p < 0.0001) but not between Sesoko and Bise corals (locations:  $F_{(1,8)} = 5.04$ , p = 0.05) with interaction (CO<sub>2</sub> × location:  $F_{(2,16)} = 8.63$ , p = 0.0029, Fig. 3). Average  $F_v F_m^{-1}$  of Sesoko and Bise corals for all CO<sub>2</sub> conditions ranged from 0.67 to 0.71 and there was no significant difference among CO<sub>2</sub> conditions and between Sesoko and Bise corals. Average zooxanthella density of Sesoko and Bise corals for all



**Fig. 3** Gross photosynthesis and respiration of the corals from Sesoko and Bise sites after 5 weeks of culture under three different CO<sub>2</sub> conditions. Mean  $\pm$  SE of five colonies. Significant differences among CO<sub>2</sub> conditions and corals from the two sites are indicated by lower case letters (Tukey HSD test *p* < 0.05)

 $CO_2$  conditions ranged from 8.6 to  $12.3 \times 10^6$  cell cm<sup>-2</sup>. No significant differences were found in  $F_v F_m^{-1}$ , zooxanthella density or lipid content among  $pCO_2$  conditions and between Sesoko and Bise corals (Fig. 4).

## **Gene expression**

Bicarbonate transporter (BAT) gene expression was significantly up-regulated at high-CO<sub>2</sub> compared to control conditions (REML, CO<sub>2</sub>:  $F_{(1,12)} = 6.02$ , p = 0.03) and was significantly different between Sesoko and Bise corals (locations:  $F_{(1,12)} = 7.52$ , p = 0.01) without interaction  $(CO_2 \times \text{location}; F_{(1,12)} = 0.15, p = 0.70, \text{Fig. 5})$ . Expression of galaxin was also up-regulated at high-CO<sub>2</sub> compared to control conditions (REML,  $CO_2$ :  $F_{(1,12)} = 26.5$ , p = 0.0002), and there was a significant difference between Sesoko and Bise corals (REML, CO<sub>2</sub>:  $F_{(1,12)} = 5.85$ , p = 0.03) and with interaction among CO<sub>2</sub> and sites  $(CO_2 \times \text{location}; F_{(1,12)} = 7.09, p = 0.02, \text{Fig. 5})$ . Plasma membrane calcium-ATPase (PMCA) and carbonic anhydrase II (CAII) expression did not change significantly between CO<sub>2</sub> conditions and Sesoko and Bise corals (Fig. 5).



**Fig. 4** Photosynthesis efficiency, zooxanthella density and lipid content of corals from Sesoko and Bise after 5 weeks of culture under three different  $CO_2$  conditions. Mean  $\pm$  SE of five colonies

# Discussion

# Physiology

In the present study, the first finding of note is that the coral Acropora digitifera collected from two different sites shows significantly different responses to seawater  $pCO_2$  change. Similar to our previous study (Takahashi and Kurihara 2013), all net calcification (G), light calcification (light G) and dark calcification (dark G) rates of the A. digitifera colonies collected from Sesoko site did not differ when incubated at control ( $pCO_2 = 420 \ \mu atm$ ) or high-CO<sub>2</sub> conditions ( $pCO_2 = 2374 \mu atm$ ), in which seawater  $\Omega_{arag}$  was under-saturated ( $\Omega_{arag} = 0.93$ ). Meanwhile, A. digitifera collected from the Bise site showed high sensitivity to elevated  $pCO_2$  and both light and dark G significantly declined with the increasing seawater  $pCO_2$ . Light and dark G of Bise corals was estimated to decrease at a rate of 2.24 and 1.87  $\mu$ mol CaCO<sub>3</sub> per  $\Omega_{arao}$ , respectively. These results suggest that the coral Acropora



**Fig. 5** Gene expression of the four putative calcification-related genes (*BAT* bicarbonate transporter, *PMCA* plasma membrane calcium ATPase, *galaxin and CAII* carbonic anhydrase II) in the control corals

from Sesoko and Bise after 5 weeks culture under three different  $CO_2$  conditions. Mean  $\pm$  SE of five colonies

*digitifera* shows intraspecific variability of calcification to high  $pCO_2$  among populations.

Coral calcification is known to be a process with a high energy demand to control the pH and  $\Omega_{arag}$  in the calicoblastic fluid where calcification occurs, and the required energy is suggested to be mainly supplied by photosynthesis (Al-Horani et al. 2003; Holcomb et al. 2014). Hence, if high  $pCO_2$  increases the photosynthesis rate of the coral, the negative effect of high  $pCO_2$  on calcification may be compensated by the additional energy. Indeed, the corals Acropora millepora collected from a CO<sub>2</sub> vent site was found to show increased net photosynthesis rate at high $pCO_2$  condition and the light G was not affected by the high  $CO_2$ , while dark G decreased in high  $pCO_2$  (Strahl et al. 2015). However, in the present study, although light G of Sesoko corals was not affected by high  $pCO_2$ , the gross photosynthesis rate (gP) of Sesoko corals was found to decrease with high-CO<sub>2</sub> conditions. Additionally, dark G was also not affected by high  $CO_2$ , suggesting that the insensitivity of net G in Sesoko corals is not related to the effect of  $CO_2$  on their photosynthesis rate. Meanwhile, Bise corals showed a strong correlation among light G and gP. Decrease in gP and suppression of metabolism gene at high  $pCO_2$  was found for A. millepora (Kaniewska et al. 2012), while photosynthesis rates were reported not to be effected by CO<sub>2</sub> on other corals such as Galaxea fascicularis, Acropora eurystoma and Cladocora caespitosa (Goiran et al. 1996; Schneider and Erez 2006; Rodolfo-Metalpa et al. 2010). These results suggest that high-seawater  $pCO_2$  condition can affect both the calcification and photosynthesis rates, although this effect could differ among coral species and populations. Additionally, another study reported a positive correlation between calcification and photochemical efficiency of photosynthesis II ( $F_v F_m^{-1}$  for the massive coral *Porites australiensis*, and authors suggested that differences in coral calcification sensitivity to high  $pCO_2$  among individuals could be related with the difference in the  $F_v F_m^{-1}$  among colonies (Iguchi et al. 2012). However, in the present study, we did not find differences in  $F_v F_m^{-1}$  between the corals from the two sites and also there was no interaction between  $F_v F_m^{-1}$  and calcification rate. Since lipid content was suggested to be a proxy of energy storage and physical conditions (Stimson 1987; Anthony 2006), here we also measured the lipid content of the corals to evaluate the possibility that energy storage and physical condition can be attributed to the difference in the sensitivity of corals to high  $CO_2$ . However, lipid content did not differ between Sesoko and Bise corals.

One notable result, however, was that Sesoko corals demonstrated significantly lower calcification rates than Bise corals at control  $pCO_2$  conditions. While the calcification rate of Bise corals was negatively affected by an increase in seawater  $pCO_2$ , Sesoko corals showed constant low calcification rate at all  $pCO_2$  concentrations in both light and dark calcification rates. Movilla et al. (2012) studied the effect of high pCO<sub>2</sub> on two temperate corals, *Cladocora caespi*tosa and Oculina patagonica, and found that although both species show negative responses in calcification rate, faster growing colonies of both species were more affected by high  $pCO_2$ . Similar findings were reported by Shaw et al. (2016) using coral Acropora pulchra and Kavousi et al. (2016) using Montipora digitata. Additionally, Comeau et al. (2014a) compared the sensitivity of eight coral species to OA and reported that tolerance of the corals to OA was correlated with their growth rate. Fast-growing corals are suggested to need higher amounts of energy to extract protons from the calicoblastic fluid compared to slow-growing corals (Tambutté et al. 2011; Comeau et al. 2014a). Consequently, it can be speculated that slow-growing Sesoko corals are more efficient to maintain the pH in the calicoblastic fluid at high-CO<sub>2</sub> condition than Bise corals, and hence they are more resistant to high CO<sub>2</sub>. Indeed when the calcifying fluid pH of Sesoko A. digitifera corals was measured using boron isotopes  $(^{11}\delta B)$ , it was found that the pH was maintained at 8.3, and  $\varOmega_{\rm arag}$  was higher than 5, even when the corals were reared at seawater, pH 7.4 (Tanaka et al. 2015).

Reasons that could explain the differences in calcification rate of the corals between the two sites are unclear. One possible reason could be due to the difference in the local scale habitat that the corals were collected; colonies collected from the Bise site were at edge of reef crest towards to the forereef, while corals collected from Sesoko site were located in a relatively more inner reef area and close to nearshore. Although there were no significant differences in the mean values of pH, temperature, salinity, oxygen concentration and light intensity measured over a 3-day period in those two sites, Sesoko site showed slightly higher seawater  $pCO_2$ , lower pH and lower  $\Omega_{arag}$  and higher variability in environmental conditions compared to the Bise site (Table 2). The study that has evaluated the skeletal extension change of the coral Siderastrea sidera colonies from forereef, backreef and nearshore in Belize since the early 1900s found that colonies from forereef and backreef sites show significantly faster growth rate than nearshore; however, the annual skeletal extension has significantly decreased in the forereef corals over the last century, while no change was found in the backreef and nearshore colonies (Castillo et al. 2011). Castillo et al. (2011) suggested that these differences could be attributed to historical differences of the environmental stress; less anthropogenic impact at forereef allowed higher calcification rates compared to nearshore colonies in the past, while forereef coral are more susceptible to environmental change than nearshore colonies which are more acclimatized and/or adapted to environmental change (e.g. more variable environmental conditions, higher sedimentation, etc.). Phenotypic plasticity for calcification among habitat was also reported in the coral Porites lobata, and reciprocal transplant experiment demonstrated that the calcification of colonies transplanted to the backreef always shows higher calcification than forereef colonies (Smith et al. 2007). Therefore, one possibility could be that the corals from the two sites show different sensitivities to OA due to the difference in calcification rate caused by the difference of local environment found in forereef (Bise site) and nearshore (Sesoko site). Alternatively, Sesoko coral could be acclimatized to more variable  $pCO_2$  and pH environmental conditions found in nearshore, and hence show higher tolerance to high-CO<sub>2</sub> conditions.

In the present study, calcification rate of the corals from the two sites was significantly different even after they were acclimated for 1 month under the same condition. However, because the distance between the Bise and Sesoko sites was less than 8 km, and previous work, evaluating the genetic connectivity of these broadcast spawning *A. digitifera* populations around Okinawa, reported that this species shows high connectivity among populations within the 1000 km long Ryukyu Archipelago, including Okinawa (Nakajima et al. 2010), genetic variations seem unlikely to be the mechanisms that could explain the variation in tolerance

Table 2 Seawater field environment in Sesoko and Bise sites

	pH (NBS scale)	Temperature (°C)	Salinity	$DO (mg L^{-1})$	TA (μmol kg <sup>-1</sup> )	DIC (µmol kg <sup>-1</sup> )	$pCO_2 (\mu atm)$	$arOmega_{ m arag}$	Light ( $\mu$ mol pho- ton m <sup>-2</sup> s <sup>-1</sup> )
Sesoko	$8.15 \pm 0.10$	$27.3 \pm 0.69$	$34.2 \pm 0.09$	$6.37 \pm 1.66$	$2265 \pm 20$	$1975 \pm 61$	$456 \pm 121$	$3.35 \pm 0.64$	$864 \pm 570$
Bise	$8.17 \pm 0.06$	$26.9 \pm 0.43$	$34.6 \pm 0.02$	$6.61 \pm 0.66$	$2227 \pm 5$	$1926 \pm 39$	$403 \pm 64$	$3.41 \pm 0.41$	$771 \pm 488$

Seawater pH, temperature, salinity, dissolved oxygen (DO) and light intensity was measured continuously for 3 days. Total alkalinity (TA) and dissolved inorganic carbon (DIC) were measured by sampling seawater, and  $pCO_2$  and  $\Omega_{arag}$  were calculated with CO2SYS program. Means  $\pm$  SD

capacity between the Sesoko and Bise corals. Some studies have found that genetic differentiation among populations even within small-scale habitats (Magalon et al. 2005); however, data that have shown local adaptation for broadcast spawning organisms are highly restricted (Sanford and Kelly 2011). Interestingly, a study that evaluated the effect of OA on different colonies of *Acropora digitifera* sampled in the inner reef area of Bise has reported that there was no significant effect of high  $CO_2$  (1000 atm) at all three studied colonies (Kavousi et al. 2015).

#### **Gene expression**

Expression of the bicarbonate transporter (BAT) gene was found to be significantly higher in the Sesoko population compared to the Bise population and being up-regulated under high-CO<sub>2</sub> conditions compared to the control in both Sesoko and Bise corals. Although the main source of dissolved inorganic carbon (DIC) used for coral calcification and photosynthesis is still under debate, using isotopes and BAT inhibitors Furla et al. (2000) showed that bicarbonate plays an important role in both mechanisms. Later, Bertucci et al. (2013) reviewed the movement of DIC within coral tissues and suggested the existence of BAT that transports bicarbonate from the oral ectoderm to endoderm cells, where bicarbonate is converted to CO<sub>2</sub> by carbonic anhydrase (CA) and supplied to the zooxanthellae. Additionally, BAT is suggested to transport bicarbonate from the aboral endoderm to calicoblastic cells and the calcifying medium, which is converted to carbonate and used for calcification. By immunostaining BAT in the coral Stylophora pistillata, it was determined that these transporters are indeed specifically localized in the endodermal and ectodermal cells of oral and aboral tissues (Zoccola et al. 2015). Similar to our study, the expression of the BAT gene was found to be up-regulated in the coral Pocillopora damicornis exposed to pH 7.4 compared to control (pH 8.1, Vidal-Dupiol et al. 2013). Although the calcification rate of *P. damicornis* was not studied by Vidal-Dupiol et al. (2013), other studies have demonstrated that this species is relatively robust to high CO<sub>2</sub> (Comeau et al. 2014b). From all these studies, up-regulation of the BAT gene can increase the availability of bicarbonate ions in the calcifying fluid, and to counteract the effect of low pH. Since we found higher expression of this gene in Sesoko corals, this could be one of the reasons for the insensitivity of Sesoko coral to the high-CO<sub>2</sub> condition. In the previous study using boron isotopes  $(^{11}\delta B)$ , the calcifying fluid pH of Sesoko A. digitifera corals was calculated to be maintained at 8.3, and  $\Omega_{\text{arag}}$  to be maintained at higher than 5, even when the corals were reared in seawater, pH 7.4 (Tanaka et al. 2015), which is in line with the present findings. Meanwhile, the reason for the higher expression of this gene in Sesoko corals, which show lower calcification rates compared to Bise corals under the control condition, is still open to question. Galaxin, which is the skeletal organic matrix protein identified in the coral (Fukuda et al. 2003), was found to be up-regulated under high CO<sub>2</sub> conditions only in Bise corals. Moya et al. (2012) evaluated the change in expression of genes encoding several organic matrices under high-CO<sub>2</sub> conditions in the juvenile coral Acropora millepora, and revealed that the expression of many of these genes was altered. For galaxins, 40% of the genes were up-regulated while about 20% were downregulated at 1000 µatm CO<sub>2</sub>. Similarly, up-regulation of this gene was also found in P. damicornis. Hence, these results suggest that the expression of galaxin could be up-regulated by an increase in CO<sub>2</sub>; however, it is not clear if there is a direct relationship between the expression of this gene and sensitivity of coral calcification to OA. Calcium-ATPase (PMCA) is hypothesized to be another key transporter that removes protons from the calcifying medium to calicoblastic cells and transports calcium ions into the calcifying medium (McConnaughey and Falk 1991; Cohen and McConnaughey 2003; Allemand et al. 2011). However, there was no significant difference in the expression of PMCA between Sesoko and Bise corals, and among CO<sub>2</sub> conditions. Meanwhile, the expression of PMCA in high-CO<sub>2</sub> Sesoko corals showed extremely high variations, suggesting PMCA expression variations among colonies. Finally, carbonic anhydrase (CA) is the enzyme that regulates the inter-conversion of HCO<sub>3</sub><sup>-</sup> to CO<sub>2</sub> and is suggested to convert CO<sub>2</sub> derived from metabolic respiration into bicarbonate, which is transported by the bicarbonate transporter genes to the calcifying fluid and used for calcification (Bertucci et al. 2013). Because CA inhibitor was found to reduce the calcification rate of corals by about 50–70%, CA is suggested to be another factor in coral calcification (Goreau 1959; Isa and Yamazato 1984). However, the expression of CA II did not differ between Sesoko and Bise corals, or among CO<sub>2</sub> conditions. These results contradict previous results showing down-regulation in A. millepora (Moya et al. 2012), and up-regulation in P. damicornis at low-pH conditions (Vidal-Dupiol et al. 2013), which suggest different response of CA to high CO<sub>2</sub> among coral species.

#### Conclusion

This study provides the first evidence that the calcification of the two populations of the coral *Acropora digitifera* shows significantly different responses to seawater  $pCO_2$  change. The population that was highly resistant to OA exhibited a slower calcification rate and higher expression of the BAT gene, suggesting that colonies that have higher capability to maintain the pH of calicoblastic fluid could be more tolerant to OA. Additionally, taking into account that the resistant populations were located at relatively nearshore where the environmental conditions are more variable, while the sensitive population was located in forereef with more stable conditions, plasticity for coral calcification in response to the environment and/or acclimation response to the change in environmental conditions such as seawater  $pCO_2$  may lead to differences in sensitivity to high seawater  $pCO_2$ . Further studies considering the potential variability in coral sensitivity to OA among local populations from different habitat are necessary to predict the effect of climate change on reef ecosystems.

**Acknowledgements** We are grateful to all the staff of Sesoko Station, Tropical Biosphere Research Center, University of the Ryukyus for their support. We also thank Hideyuki Yamashiro for help with lipid content analysis.

**Funding** This work was conducted with the support of funding from the Japan Society for the Promotion of Science (JSPS), and JST, CREST program.

### **Compliance with ethical standards**

**Conflict of interest** The authors have no conflict of interest.

**Ethical approval** The manuscript has not been submitted to more than one journal for simultaneous consideration. All procedures performed in studies involving animals were in accordance with the ethical standards of the institution or practice at which the studies were conducted.

# References

- Al-Horani FA, Al-Moghrabi SM, de-Beer D (2003) The mechanism of calcification and its relation to photosynthesis and respiration in the scleractinian coral *Galaxea fascicularis*. Mar Biol 142:419–426
- Allemand D, Tambutté É, Zoccola D, Tambutté S (2011) Coral calcification, cells to reefs. In: Dubinsky Z, Stambler N (eds) Coral reefs: an ecosystem in transition. Springer, Dordrecht, pp 119–150
- Anthony KRN (2006) Enhanced energy status of corals on coastal, high-turbidity reefs. Mar Ecol Prog Ser 319:111–116
- Bertucci A, Moya A, Tambutté S, Allemand D, Supuran CT, Zoccola D (2013) Carbonic anhydrases in anthozoan corals—a review. Bioorg Med Chem 21:1437–1450
- Castillo KD, Ries JB, Weiss JM (2011) Declining coral skeleton extension for forereef colonies of *Siderastrea siderea* on the Mesoamerican Barrier Reef System, southern Belize. PLoS ONE 6:e14615
- Chan NCS, Connolly SR (2013) Sensitivity of coral calcification to ocean acidification: a meta-analysis. Glob Change Biol 19:282–290
- Cohen AL, McConnaughey TA (2003) Geochemical perspectives on coral mineralization. Rev Mineral Geochem 54:151–187
- Comeau S, Edmunds PJ, Spindel NB, Carpenter RC (2014a) Fast coral reef calcifiers are more sensitive to ocean acidification in shortterm laboratory experiments. Limnol Oceanogr 59:1081–1091
- Comeau S, Carpenter RC, Nojiri Y, Putman HM, Sakai K, Edmunds PJ (2014b) Pacific-wide contrast highlights resistance of reef calcifiers to ocean acidification. Proc R Soc B 281:20141339

- Davis PS (1989) Short-term growth measurements of coral using an accurate buoyant weighing technique. Mar Biol 101:389–395
- De Putron SJ, McCorkle DC, Cohen AL, Dillon AB (2011) The impact of seawater saturation state and biocarbonate ion concentration in calcification by new recruits of two Atlantic corals. Coral Reefs 30:321–328
- Fabricius KE, Langdon C, Uthicke S, Humphrey C, Noonan S, De'ath G, Okazaki R, Muehllehner N, Glas MS, Lough JM (2013) Losers and winners in coral reefs acclimatized to elevated carbon dioxide concentrations. Nat Clim Change 1:165–169
- Fukuda I, Ooki S, Fujita T, Murayama E, Nagasawa H, Isa Y, Watanabe T (2003) Molecular cloning of a cDNA encoding a soluble protein in the coral exoskeleton. Biochem Biophys Res Commun 304:11–17
- Furla P, Galgani I, Durand I, Allemand D (2000) Sources and mechanisms of inorganic carbon transport for coral calcification and photosynthesis. J Exp Biol 203:3445–3457
- Goiran C, Al-Moghrabi S, Allemand D, Jaubert J (1996) Inorganic carbon uptake for photosynthesis by the symbiotic coral/dinoflagellate association I. Photosynthetic performances of symbionts and dependence on sea water biocarbonate. J Exp Mar Biol Ecol 199:207–225
- Goreau TF (1959) The physiology of skeleton formation in corals. I. A method for measuring the rate of calcium deposition by corals under different conditions. Biol Bull 116:59–75
- Grasso LC, Maindonald J, Rudd S, Hayward DC, Saint R, Miller DJ, Ball EE (2008) Microarray analysis identifies candidate genes for key roles in coral development. BMC Genom 9:540
- Hayward DC, Hetherington S, Behm CA, Grasso LC, Foret S, Miller DJ, Ball EE (2011) Differential gene expression at coral settlement and metamorphosis—a subtractive hybridization study. PLoS ONE 6:e26411
- Hoegh-Guldberg O, Mumby PJ, Hooten AJ, Steneck RS, GreenWeld P, Gomez E, Harvell CD, Sale PF et al (2007) Coral reef under rapid climate change and ocean acidification. Science 318:1737–1742
- Holcomb M, Venn AA, Tambutté E, Tambutté S, Allemand D, Trotter J, McCulloch M (2014) Coral calcifying fluid pH dictates response to ocean acidification. Sci Rep 4:5207
- Iguchi A, Ozaki S, Nakamura T, Inoue M, Tanaka Y, Suzuki A, Kawahata H, Sakai K (2012) Effects of acidified seawater on coral calcification and symbiotic algae on the massive coral *Porites australiensis*. Mar Environ Res 73:32–36
- Inoue S, Kayanne H, Yamamoto S, Kurihara H (2013) Spatial community shift from hard to soft corals in acidified water. Nat Clim Change 3:683–687
- Isa Y, Yamazato K (1984) The distribution of carbonic anhydrase in a staghorn coral *Acropora hebes* (Dana). Galaxea 3:25–36
- Jokiel PL, Rodgers KS, Kuffner LB, Andersson AJ, Cox EF, Mackenzie FT (2008) Ocean acidification and calcifying reef organisms: a mesocosm investigation. Coral Reefs 27:473–483
- Jury CP, Whitehead RF, Szmant AM (2010) Effects of variations in carbonate chemistry on the calcification rates of *Madracis auretenra* (= *Madracis mirabilis* sensu Wells, 1973): biocarbonate concentrations best predict calcification rates. Glob Change Biol 16:1632–1644
- Kaniewska P, Campbell PR, Kline DI, Rodriguez-Lanetty M, Miller DJ, Dove S, Hoegh-Guldberg O (2012) Major cellular and physiological impacts of ocean acidification on a reef building coral. PLoS ONE 7:e34659
- Kavousi J, Reimer JD, Tanaka Y, Nakamura T (2015) Colony-specific investigations reveal highly variable responses among individual corals to ocean acidification and warming. Mar Environ Res 109:9–20
- Kavousi J, Tanaka Y, Nishida K, Suzuki A, Nojiri Y, Nakamura T (2016) Colony-specific calcification and mortality under ocean

acidification in the branching coral *Montipora digitata*. Mar Environ Res 119:161–165

- Kleypas JA, Feely RA, Fabry VJ, Langdon C, Sabine CL, Robbins LL (2006) Impacts of ocean acidification on coral reefs and other marine calcifiers: a guide for future research. Report of a workshop held 18–20 April 2005, St. Petersburg, FL, sponsored by NSF, NOAA, and the U.S. Geological Survey
- Langdon C, Atkinson MJ (2005) Effect of elevated *p*CO<sub>2</sub> on photosynthesis and calcification of corals and interactions with seasonal change in temperature/irradiance and nutrient enrichment. J Geophys Res 110:C09S07
- Lewis E, Wallace DWR (1998) Program developed for CO<sub>2</sub> system calculations. ORNL/CDIAC-105 Carbon dioxide information analysis center, Oak Ridge National Laboratory, U.S. Department of Energy, Oak Ridge
- Magalon H, Adjeroud M, Veuille M (2005) Patterns of genetic variation do not correlate with geographical distance in the reef-building coral *Pocillopora meandrina* in the South Pacific. Mol Ecol 14:1861–1868
- McConnaughey TA, Falk RH (1991) Calcium–proton exchange during algal calcification. Biol Bull 180:185–195
- Mehrbach C, Culberson CH, Hawley JE, Pytkowicz RM (1973) Measurement of the apparent dissociation constant of carbonic acid in seawater at atmospheric pressure. Limnol Oceanog 18:897–907
- Movilla J, Calvo E, Pelejero C, Coma R, Serrano E, Fernández-Vallejo P, Ribes M (2012) Calcification reduction and recovery in native and non-native Mediterranean corals in response to ocean acidification. J Exp Mar Biol Ecol 438:144–153
- Moya A, Huisman L, Ball EE, Hayward DC, Grasso LC, Chua CM, Woo HN, Gattuso J-P, Foret S, Miller DJ (2012) Whole transcriptome analysis of the coral *Acropora millepora* reveals complex responses to CO<sub>2</sub>-driven acidification during the initiation of calcification. Mol Ecol 21:2440–2454
- Mucci A (1983) The solubility of calcite and aragonite in seawater at various salinities, temperatures, and one atmosphere total pressure. Am J Sci 283:780–799
- Nakajima Y, Nishikawa A, Iguchi A, Sakai K (2010) Gene flow and genetic diversity of a broadcast spawning coral in northern peripheral populations. PLoS ONE 5:e11149
- Pandolfi JM, Connolly SR, Marshall DJ, Cohen AL (2011) Projecting coral reef futures under global warming and ocean acidification. Science 333:418–422
- Parker LM, Ross PM, O'Connor WA (2011) Populations of the Sydney rock oyster, *Saccostrea glomerata*, vary in response to ocean acidification. Mar Biol 158:689–697
- Pfaffl MW (2001) A new mathematical model for relative quantification in real-time RT-PCR. Nucleic Acids Res 29:e45
- Range P, Chícharo MA, Ben-Hamadou R, Pilo D, Fernandez-Reiriz MJ, Labarta U, Marin MG, Bressan M, Matozzo V, Chinellato A, Munari M, El Menif NT, Dellali M, Chícharo L (2014) Impacts of CO<sub>2</sub>-induced seawater acidification on coastal Mediterranean bivalves and interactions with other climatic stressors. Reg Environ Change 14:19–30
- Reyes-Bermudez A, Lin Z, Hayward DC, Miller DJ, Ball EE (2009) Differential expression of three galaxin-related genes during settlement and metamorphosis in the scleractinian coral Acropora millepora. BMC Evol Biol 9:178
- Rodolfo-Metalpa R, Martin S, Ferrier-Pagès C, Gattuso J-P (2010) Response of temperate coral *Cladocora caespitosa* to mid- and long-term exposure to pCO<sub>2</sub> and temperature levels projected for the year 2100 AD. Biogeosciences 7:289–300
- Sanford E, Kelly MW (2011) Local adaptation in marine invertebrates. Annu Rev Mar Sci 3:509–535

- Schneider K, Erez J (2006) The effect of carbonate chemistry on calcification and photosynthesis in the hermatypic coral *Acropora eurystoma*. Limnol Oceanogr 51:1284–1293
- Schreiber U, Schliwa U, Bilger W (1986) Continuous recording of photochemical and non-photochemical chlorophyll fluorescence quenching with new type of modulation fluorometer. Photosynth Res 10:51–62
- Shaw EC, Carpenter RC, Lantz CA, Edmunds PJ (2016) Intraspecific variability in the response to ocean warming and acidification in the scleractinian coral Acropora pulchra. Mar Biol 163:210
- Smith SV (1973) Carbon dioxide dynamics: a record of organic carbon production, respiration, and calcification in the Eniwetok reef flat community. Limnol Oceanogr 18:106–120
- Smith SV, Key GS (1975) Carbon dioxide and metabolism in marine environments. Limnol Oceanogr 20:493–495
- Smith LW, Barshis D, Birkeland C (2007) Phenotypic plasticity for skeleton growth, density and calcification of *Porites lobata* in response to habitat type. Coral Reefs 26:559–567
- Souter P, Bay LK, Andreakis N, Császár N, Seneca FO, Van Oppen MJH (2010) A multilocus, temperatures stress-related gene expression profile assay in *Acropora millepora*, a dominant reefbuilding coral. Mol Ecol Resour 11:328–334
- Stimson JS (1987) Location, quantity and rate of change in quantity of lipid in tissue of Hawaiian hermatypic corals. Bull Mar Sci 21:889–904
- Strahl J, Stolz I, Uthicke S, Vogel N, Noonan SHC, Fabricius KE (2015) Physiological and ecological performance differs in four coral taxa at a volcanic carbon dioxide seep. Comp Biochem Physiol A 184:170–186
- Sunday JM, Crim RN, Harley DG, Hart MW (2011) Quantifying rates of evolutionary adaptation in response to ocean acidification. PLoS ONE 6:e22881
- Takahashi A, Kurihara H (2013) Ocean acidification does not affect the physiology of the tropical coral *Acropora digitifera* during a 5-week experiment. Coral Reefs 32:305–314
- Tambutté S, Holcomb M, Ferrier-Pagès C, Reynaud S, Tambutté E, Zoccola D, Allemand D (2011) Coral biomineralization: from the gene to the environment. J Exp Mar Biol Ecol 408:58–78
- Tanaka K, Holcomb M, Takahashi A, Kurihara H, Asami R, Shinjo R, Sowa K, Rankenburg K, Watanabe T, McCulloch M (2015) Response of *Acropora digitifera* to ocean acidification: constraints from  $\delta^{11}$ B, Sr, Mg, and Ba compositions of aragonitic skeletons cultured under variable seawater pH. Coral Reefs 34:1139–1149
- Thomsen J, Gutowska MA, Saphörster J, Heinemann A, Trübenbach K, Fietzke J, Hiebenthal C, Eisenhauer A, Körtzinger A, Wahl M, Melzner F (2010) Calcifying invertebrates succeed in naturally CO<sub>2</sub>-rich coastal habitat but are threatened by high levels of future acidification. Biogeosciences 7:3879–3891
- Vidal-Dupiol J, Zoccola D, Tambutté E, Grunau C, Cosseau C, Smith KM, Freitag M, Dheilly NM, Allemand D, Tambutté S (2013)
   Genes related to ion-transport and energy production are upregulated in response to CO<sub>2</sub>-driven pH decrease in corals: new insights from transcriptome analysis. PLoS ONE 8:e58652
- Yamashiro H, Oku H, Higa H, Chinen I, Sakai K (1999) Composition of lipids, fatty acids and sterols in Okinawa corals. Comp Biochem Physiol Part B Biochem Mol Biol 122:397–407
- Zoccola D, Tambutté E, Kulhanek E, Puverel S, Scimeca J-C, Allemand D, Tambutté S (2004) Molecular cloning and localization of a PMCA P-type calcium ATPase from the coral *Stylophora pistillata*. Biochim Biophys Acta 1663:1–27
- Zoccola D, Ganot P, Bertucci A, Caminiti-Segonds N, Techer N, Voolstra CR, Aranda M, Tambutté E, Allemand D, Casey JR, Tambutté S (2015) Bicarbonate transporters in corals point towards a key step in the evolution of cnidarian calcification. Sci Rep 5:09983