



# Guide to the Science of Climate Change in the 21<sup>st</sup> Century

## Chapter 18 IPCC 2014 AR5 - Impacts of Climate Change on Physical Systems

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[reached?utm\\_campaign=Daily%20Briefing&utm\\_content=20220224&utm\\_medium=email&utm\\_source=Revue%20newsletter](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter11_FINAL.pdf)

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# Chapter 18.0 IPCC 2014 AR5 - Impacts of Climate Change on Physical Systems

## 18.1 Introduction

Climate change impacts all elements in the hydrologic cycle as discussed in Chapter 6. The entire global community will be affected, directly or indirectly. 'Business as usual' will not be an option. Prediction of the changes to the hydrologic cycle, particularly precipitation events in terms of type, timing, variation in quantity and nature of extremes is critical to developing adaptation strategies including water management.

Changes in characteristics of snowfall and rainfall, not just increases or decreases in annual volume, but also the timing of the occurrence of normally expected events such as monsoons or 'rainy seasons', can result in water shortfalls or flooding that had previously been rare events. Changes in temperature together with changes in seasonal precipitation may negatively impact dryland agriculture and pasture for livestock (drought conditions) and water supply for domestic use, irrigation and hydropower generation. Flooding will result from unexpected (not forecast) runoff flow rate and volume. Groundwater might be depleted to where it is no longer a useable water supply. Water management infrastructure may be rendered inadequate or inappropriate.

The observed impacts of climate change on physical systems motivated the development of climate models able to predict future global warming and climate changes resulting from sustained or increased greenhouse gas emissions and changes in land use. Climate science is sufficiently known that it is now possible to create and use climate models to predict climate change as a consequence of anthropogenic and natural causes. This capability has been demonstrated, at least as it has occurred, since 1850. These models are the only tools available to guide society's adaptation and mitigation responses to impending impacts on natural and human environments.

The IPCC Fifth Assessment Report, AR5, was completed in 2013 and is the primary focus of this chapter. There are separate discussions on climate change tipping points and the impact of warming on the release of runoff from snow melt in mountainous regions. While snow is a very important element in discussions of the cryosphere, timing of runoff from mountain snow accumulation and melt is also very important. The impact of climate change on runoff volume (including snow melt) is discussed in Section 18.3.2.

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The IPCC Sixth Assessment Report, AR6, is building on more recent information, model improvements and a different set of scenarios. The AR6 report, WG1, The Physical Science Basis, was published in the fall 2021.

Prior to publication of the AR6 report, WG2: Impacts, Adaptation and Vulnerability in 2022 draft versions of segments of this report were being leaked (June 22, 2021) which discuss the serious nature of climate change tipping points. It is the authors opinion that while this ‘leaked’ information was unofficial and confidential it is very valuable that some of the more serious issues be known at this time – well ahead of COP26 held in the fall of 2021. It motivated the ‘parties’ to seriously deal with the urgency of the issues that need to be resolved. AR6 WGII Climate Change 2022: Impacts, Adaptation and Vulnerability was published February 28, 2022 <https://www.ipcc.ch/report/sixth-assessment-report-working-group-ii/>.

## 18.2 Scenarios AR5

Climate models have been used since AR1 (the first assessment report prepared in the 1990’s). With every assessment report the ability to model climate has improved and so have the scenarios used by the climate models to predict future climate change impacts resulting from different patterns of GHG emissions. The patterns reported and used by the climate models are the result of the analysis of the physical circumstances and societal responses expected. This has continued with AR5.

The approach taken by the IPCC for AR5 is to use what they call representative concentration pathways or RCPs. Concentration is measured in terms of radiative forcing,  $\text{Wm}^{-2}$ , similar to that used in Figure 15.29 which reports global-average radiative forcing estimates and ranges for both anthropogenic forcings and natural forcings.

Integrated Assessment Models, IAMs, were used to select the RCPs. A description of complex IAMs is provided in the article, Integrated assessment climate policy models have proven useful, with caveats, <https://www.pnas.org/doi/10.1073/pnas.2101899118> and Carbon Brief (02.10.2018 <https://www.carbonbrief.org/qa-how-integrated-assessment-models-are-used-to-study-climate-change>): ‘These models look at energy technologies, energy use choices, land-use changes and societal trends that cause – or prevent – greenhouse gas emissions.’ IAMs consider all aspects of human life (social, economic, governance, land-use and more) to develop scenarios or pathways used in the modelling studies discussed in Chapters 18 and 21. The IAM approach is described in a Carbon Brief Explainer, Q&A: How integrated assessment models are used to study climate change, <https://www.carbonbrief.org/qa-how-integrated-assessment-models-are-used-to-study-climate-change>.

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Four pathways were selected. Figure 18.1 illustrates the four RCPs identified and how the radiative forcing change annually to 2100. The names of the pathways refer to the radiative forcing from GHG concentration expected in year 2100. RCP8.5 pathway will result in GHG radiative forcing equal to 8.5 watts per square meter, RCP2.6 will result in a radiative forcing equal to 2.6 watts per square meter and so on. Table 18.1 discusses each of the four RCPs selected. They are not the only options but they are representative of the sort of emission patterns that will lead to a specific outcome. An overview of the representative concentration pathways used in AR5 and the climate models, CMIP5, may be found in <https://link.springer.com/article/10.1007/s10584-011-0148-z> and [https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf) pages 1045 – 1047.

Clearly, strategies for achieving the various RCPs place emphasis on reduction of GHGs (CO<sub>2</sub> - equivalent) as this is the major cause for anthropogenic radiative forcing as shown in Figure 15.29.

An important discussion on emission scenarios (and RCPs) may be found in a guest post in Carbon Brief 30 March 2022 which describes how not to interpret the emissions scenarios in the IPCC report, [https://www.carbonbrief.org/guest-post-how-not-to-interpret-the-emissions-scenarios-in-the-ipcc-report?utm\\_campaign=Daily%20Briefing&utm\\_content=20220331&utm\\_medium=email&utm\\_source=Revue%20newsletter](https://www.carbonbrief.org/guest-post-how-not-to-interpret-the-emissions-scenarios-in-the-ipcc-report?utm_campaign=Daily%20Briefing&utm_content=20220331&utm_medium=email&utm_source=Revue%20newsletter) .

The significance of following the different pathways on temperature or temperature change are shown in Figures 18.3 and 18.4.

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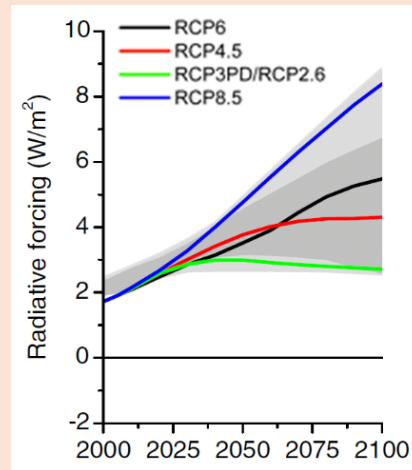


Figure 18.1 Representative concentration pathways or RCP's. Note that the name of the RCP; for example, RCP 8.5 refers to the radiative forcing in the year 2100, <https://link.springer.com/article/10.1007/s10584-011-0148-z> .

RCP Description and Citations			
	Description	IA Model	Publication – IA Model
RCP8.5	Rising radiative forcing pathway leading to 8.5 W/m <sup>2</sup> in 2100.	MESSAGE	Riahi et al. (2007) Rao & Riahi (2006)
RCP6	Stabilization without overshoot pathway to 6 W/m <sup>2</sup> at stabilization after 2100	AIM	Fujino et al. (2006) Hijioka et al. (2008)
RCP4.5	Stabilization without overshoot pathway to 4.5 W/m <sup>2</sup> at stabilization after 2100	GCAM (MiniCAM)	Smith and Wigley (2006) Clarke et al. (2007) Wise et al. (2009)
RCP2.6	Peak in radiative forcing at ~ 3 W/m <sup>2</sup> before 2100 and decline	IMAGE	van Vuuren et al. (2006; 2007)

Table 18.1 Description of representative concentration pathways RCP along with associated integrated assessment model IAM

### 18.3 Projected physical impacts.

#### 18.3.1 Temperature

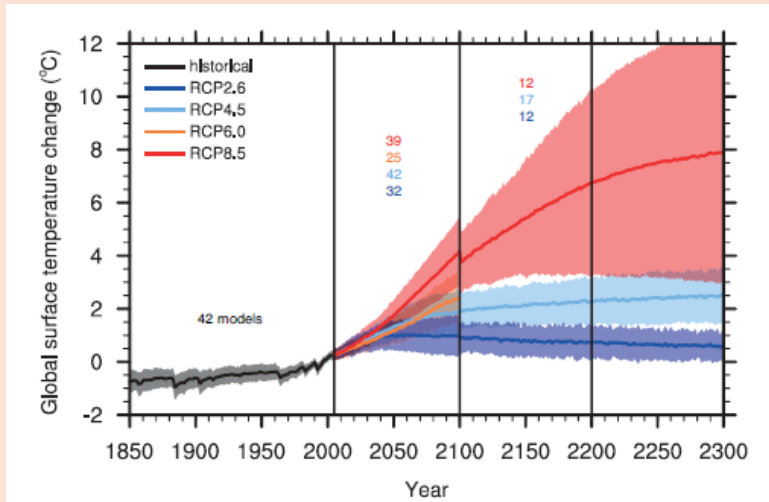
The observed and projected global surface temperature changes from 1850 – 2300 are shown in Figure 18.3. (All of the CMIP5 models were used as indicated.) The shaded zones of the same

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colour as the RCP represents the range of temperature change. It is very large for RCP8.5 and much narrower for RCP2.6. The narrower the range, the more reliable the prediction. Projections past 2100 are of interest because only RCP2.6 indicates a stable or gradually decreasing temperature after 2100. The ONLY scenario that MIGHT limit global warming on land to 1.5 degrees centigrade in 2100 is RCP 2.6.



**Figure 12.5 |** Time series of global annual mean surface air temperature anomalies (relative to 1986–2005) from CMIP5 concentration-driven experiments. Projections are shown for each RCP for the multi-model mean (solid lines) and the 5 to 95% range ( $\pm 1.64$  standard deviation) across the distribution of individual models (shading). Discontinuities at 2100 are due to different numbers of models performing the extension runs beyond the 21st century and have no physical meaning. Only one ensemble member is used from each model and numbers in the figure indicate the number of different models contributing to the different time periods. No ranges are given for the RCP6.0 projections beyond 2100 as only two models are available.

Figure 18.3 Observed and projected global surface temperature change 1850 – 2300. [https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf)

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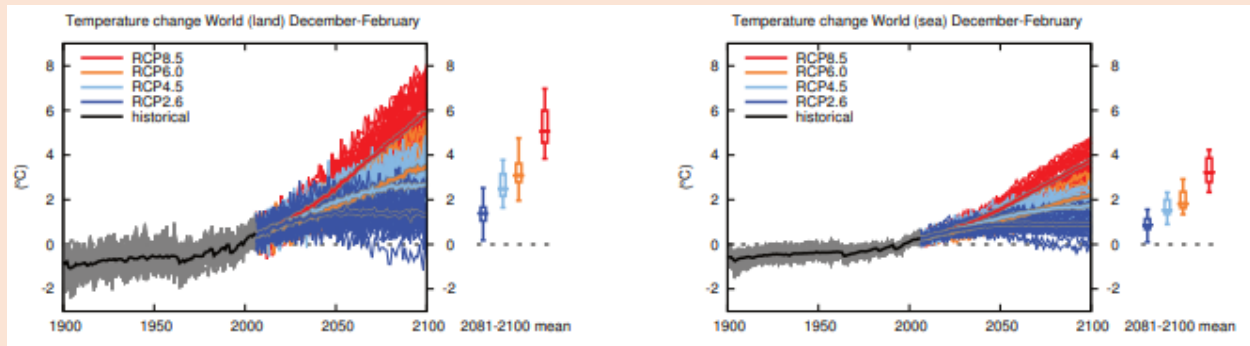


Figure 18.4 Observed and projected global surface and sea temperature change 1850 – 2100.  
[https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf)

Table 18.2 spells out the consequences of following the different RCPs in terms of reaching global target temperatures. The lower the probability of exceeding a certain temperature the more successful the RCP might be in limiting temperature increase. All of the RCPs appear to result in temperature increase greater than 1 degree by mid century (2046-2065). Only RCP2.6 might limit temperature increase below 1.5 degree by mid-century (2046-2065) – there is a 56% probability.

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Scenario	Early (2016–2035)	Mid (2046–2065)
<b>Temperature +1.0°C</b>		
RCP 2.6	100% (84%)	100% (94%)
RCP 4.5	98% (93%)	100% (100%)
RCP 6.0	96% (80%)	100% (100%)
RCP 8.5	100% (100%)	100% (100%)
<b>Temperature +1.5°C</b>		
RCP 2.6	22% (0%)	56% (28%)
RCP 4.5	17% (0%)	95% (86%)
RCP 6.0	12% (0%)	92% (88%)
RCP 8.5	33% (5%)	100% (100%)
<b>Temperature +2.0°C</b>		
RCP 2.6	0% (0%)	16% (3%)
RCP 4.5	0% (0%)	43% (29%)
RCP 6.0	0% (0%)	32% (20%)
RCP 8.5	0% (0%)	95% (90%)
<b>Temperature +3.0°C</b>		
RCP 2.6	0% (0%)	0% (0%)
RCP 4.5	0% (0%)	0% (0%)
RCP 6.0	0% (0%)	0% (0%)
RCP 8.5	0% (0%)	21% (5%)

Table 18.2 AR5 Percentage of CMIP5 models for which the projected change in global mean surface air temperature, relative to 1850-1900, crosses the specified temperature levels, by the specified time periods and assuming the specified RCP scenarios. This table is the same as Table 11.3 in [https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter11\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter11_FINAL.pdf) .

Table 18.3 is difficult to interpret but important. The mean (average) surface air temperature in degrees centigrade in the years 1986 to 2005 is the base from which temperature increase are measured. The objective is to limit the increase to 2 degrees by 2100. **The only RCP that might fulfill this objective is RCP2.6.**

Table 18.4 indicates that there is a 94% probability that RCP2.6 will result in a temperature change greater than 1 degree, a 56% probability that it will result in a temperature increase greater than 1.5 degrees and a 22% probability that it will result in a temperature increase greater than 2 degrees. **Again, RCP2.6 is the only RCP that might satisfy the objective of limiting temperature increase at or below 2 degrees by 2100.**

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	RCP2.6 ( $\Delta T$ in °C)	RCP4.5 ( $\Delta T$ in °C)	RCP6.0 ( $\Delta T$ in °C)	RCP8.5 ( $\Delta T$ in °C)
Global: 2046–2065	1.0 ± 0.3 (0.4, 1.6)	1.4 ± 0.3 (0.9, 2.0)	1.3 ± 0.3 (0.8, 1.8)	2.0 ± 0.4 (1.4, 2.6)
2081–2100	1.0 ± 0.4 (0.3, 1.7)	1.8 ± 0.5 (1.1, 2.6)	2.2 ± 0.5 (1.4, 3.1)	3.7 ± 0.7 (2.6, 4.8)
2181–2200	0.7 ± 0.4 (0.1, 1.3)	2.3 ± 0.5 (1.4, 3.1)	3.7 ± 0.7 (-, -)	6.5 ± 2.0 (3.3, 9.8)
2281–2300	0.6 ± 0.3 (0.0, 1.2)	2.5 ± 0.6 (1.5, 3.5)	4.2 ± 1.0 (-, -)	7.8 ± 2.9 (3.0, 12.6)
Land: 2081–2100	1.2 ± 0.6 (0.3, 2.2)	2.4 ± 0.6 (1.3, 3.4)	3.0 ± 0.7 (1.8, 4.1)	4.8 ± 0.9 (3.4, 6.2)
Ocean: 2081–2100	0.8 ± 0.4 (0.2, 1.4)	1.5 ± 0.4 (0.9, 2.2)	1.9 ± 0.4 (1.1, 2.6)	3.1 ± 0.6 (2.1, 4.0)
Tropics: 2081–2100	0.9 ± 0.3 (0.3, 1.4)	1.6 ± 0.4 (0.9, 2.3)	2.0 ± 0.4 (1.3, 2.7)	3.3 ± 0.6 (2.2, 4.4)
Polar: Arctic: 2081–2100	2.2 ± 1.7 (-0.5, 5.0)	4.2 ± 1.6 (1.6, 6.9)	5.2 ± 1.9 (2.1, 8.3)	8.3 ± 1.9 (5.2, 11.4)
Polar: Antarctic: 2081–2100	0.8 ± 0.6 (-0.2, 1.8)	1.5 ± 0.7 (0.3, 2.7)	1.7 ± 0.9 (0.2, 3.2)	3.1 ± 1.2 (1.1, 5.1)

Table 18.3 CMIP5 annual mean surface air temperature anomalies (°C) from the 1986–2005 reference period for selected time periods, regions and RCPs. From Table 12.2 [https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf).

	$\Delta T$ (°C) 2081–2100	$\Delta T > +1.0^\circ\text{C}$	$\Delta T > +1.5^\circ\text{C}$	$\Delta T > +2.0^\circ\text{C}$	$\Delta T > +3.0^\circ\text{C}$	$\Delta T > +4.0^\circ\text{C}$
RCP2.6	1.6 ± 0.4 (0.9, 2.3)	94%	56%	22%	0%	0%
RCP4.5	2.4 ± 0.5 (1.7, 3.2)	100%	100%	79%	12%	0%
RCP6.0	2.8 ± 0.5 (2.0, 3.7)	100%	100%	100%	36%	0%
RCP8.5	4.3 ± 0.7 (3.2, 5.4)	100%	100%	100%	100%	62%

Table 18.4 CMIP5 global annual mean temperature changes above 1850-1900 for the 2081–2100 period of each RCP scenario. From Table 12.3 [https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf).

A clear illustration of how surface air temperature will change globally with the different RCPs is shown in Figure 18.5. The top line of maps show how temperatures will change for RCP2.6 for the years 2046 – 2065, 2081 – 22100, and 2181 – 2200. There is a row of maps for each of the RCPs. RCP8.5 predicts a very warm Earth.

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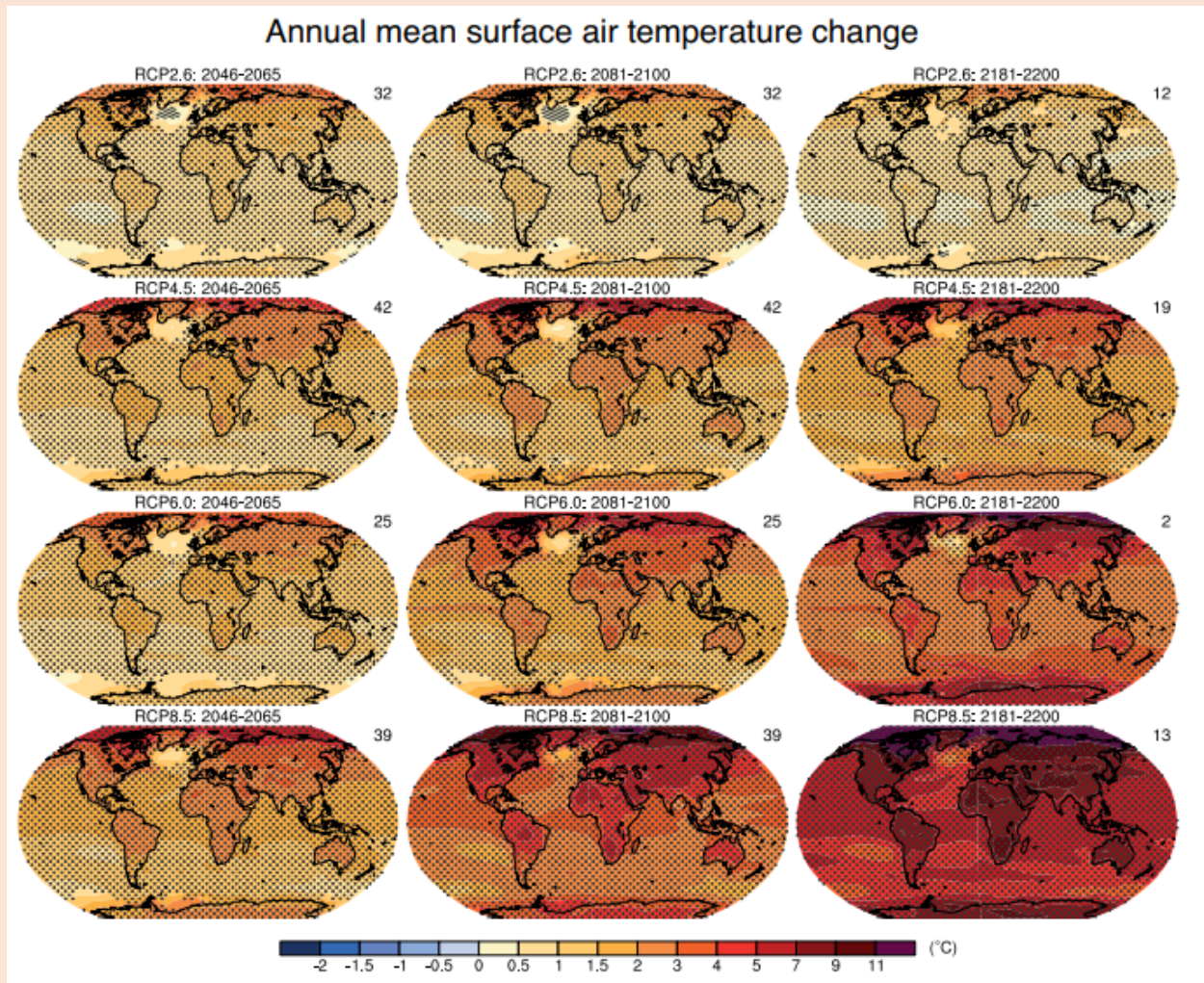


Figure 18.5 projected annual mean surface air temperature change from 1986-2005 average.  
[https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf)

The effects of global warming are more pronounced at the poles in what is called polar amplification, <http://www.realclimate.org/index.php/archives/2006/01/polar-amplification/>.

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Figure 18.6 illustrates how minimum and maximum temperatures; number of frost days and number of tropical nights are projected to change following RCP2.6, RCP4.5 and RCP8.5 to 2100. RCP2.6 is projected to stabilize around 2050.

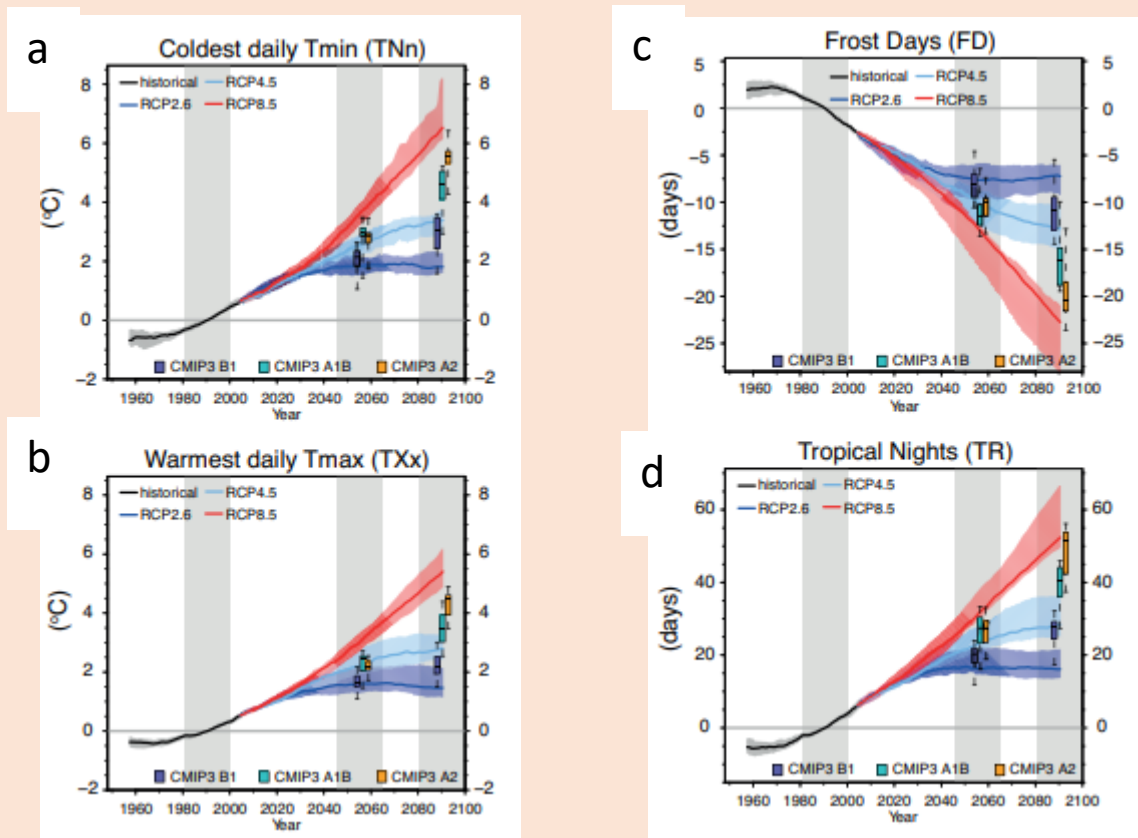


Figure 18.6 Historical and projected annual minimum of daily minimum, annual warmest daily of daily warmest, days of frost (below 0°C) and days of tropical nights (above 20°C).

[https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf)

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### 18.3.2 Water

The projected annual near-surface soil moisture changes for all RCPs for 2018 – 2100 are shown in Figure 18.7. See surface temperature and precipitation. If surface temperature is high and precipitation is low, soil moisture will be low. Note sub-Saharan Africa and northern South America.

Projected annual runoff is shown in Figure 18.8. High soil moisture and high precipitation will result in increased runoff.

Observed and projected world surface and sea precipitation change 1850 – 2100 are shown in Figure 18.9.

Observed and projected global surface and sea evaporation change 2081-2100 are shown in Figure 18.10.

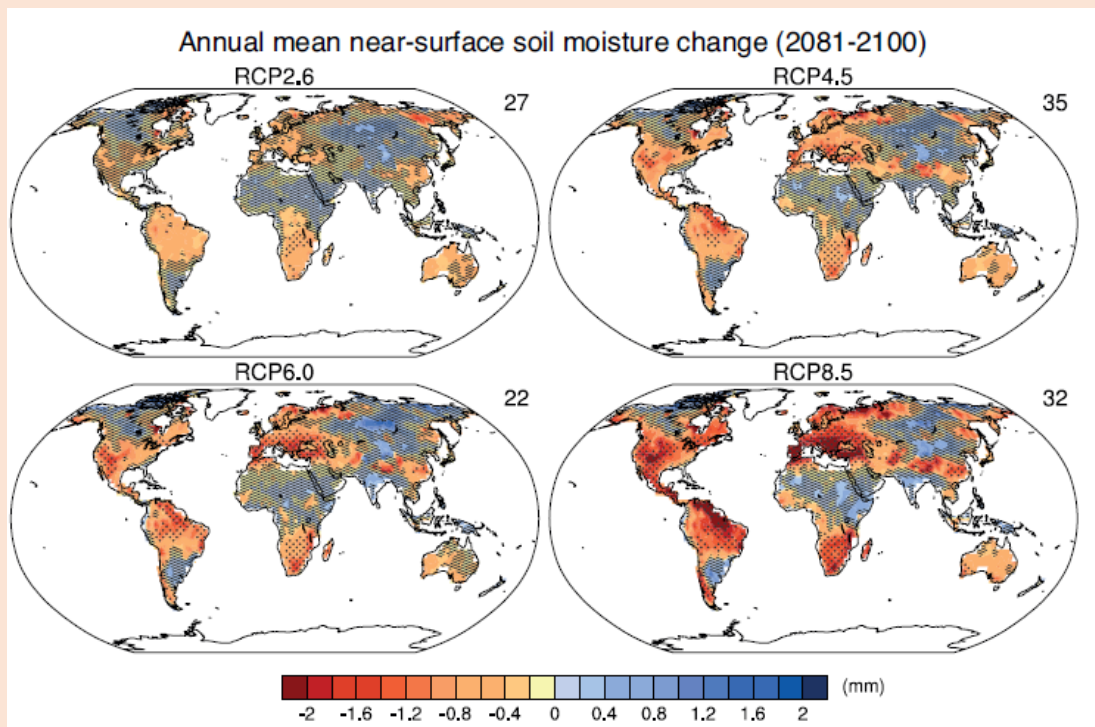


Figure 18.7 Projected annual mean near-surface soil moisture change (2018-2100)  
[https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf)

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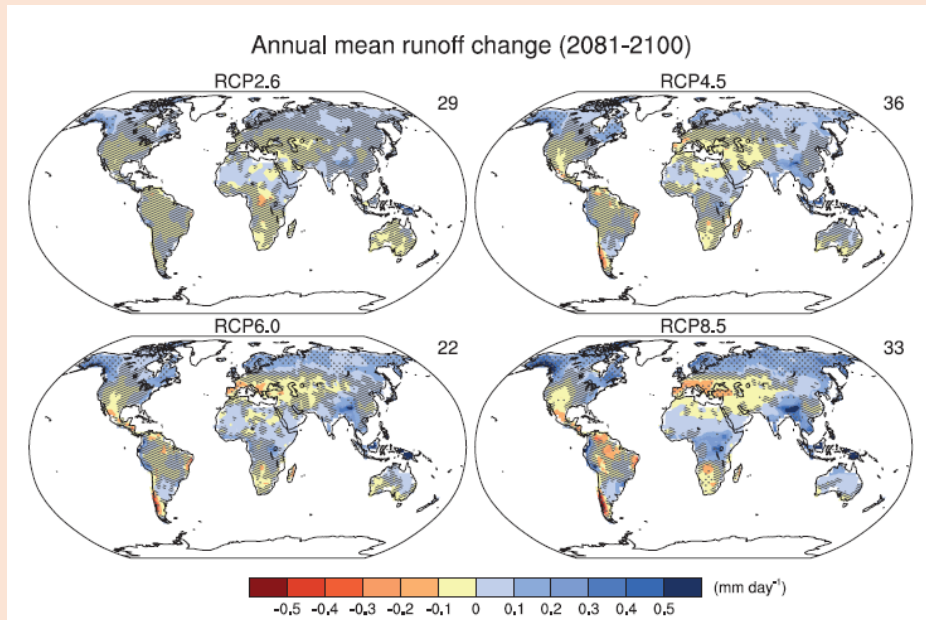


Figure 18.8 Projected annual mean runoff change (2081-2100).

[https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf)

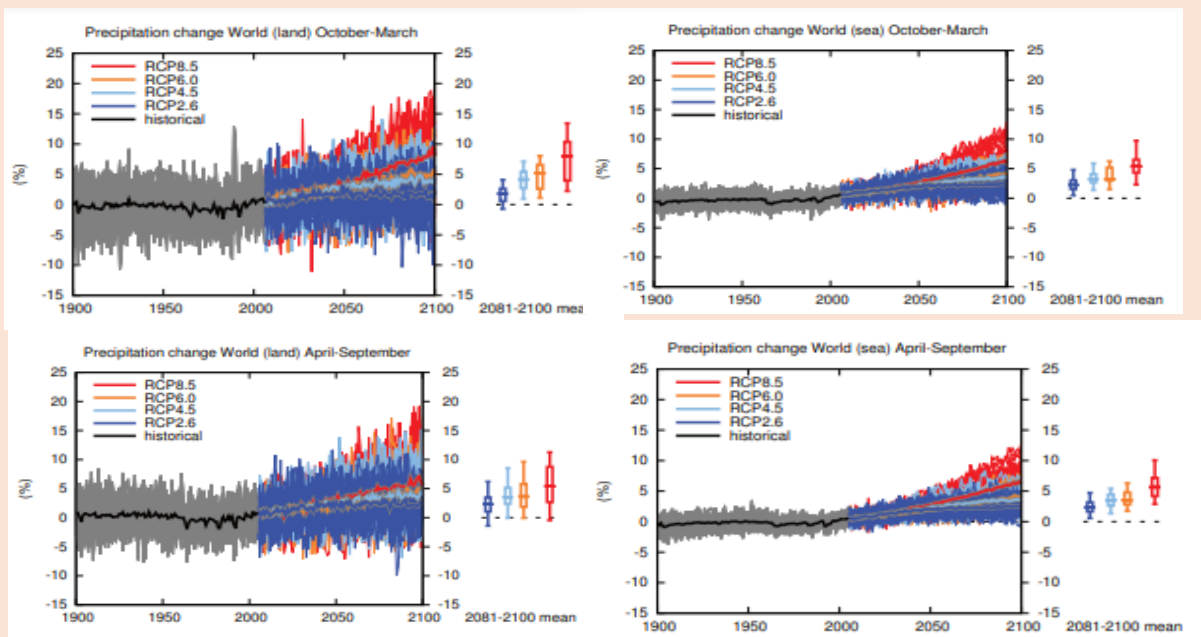


Figure 18.9 Observed and projected world surface and sea precipitation change 1850 – 2100.

[https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf)

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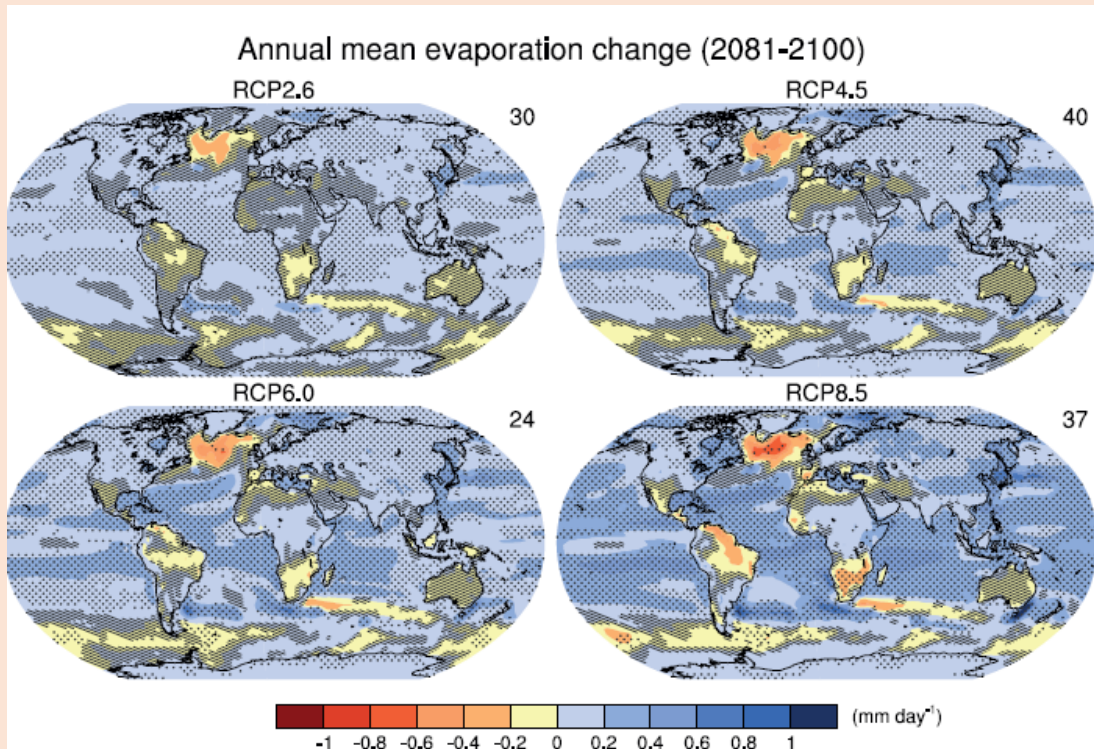


Figure 18.10 Observed and projected global surface and sea evaporation change 2081-2100. [https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf)

### 18.3.3 Cryosphere

Observed and projected sea ice extent in the northern and southern hemispheres to 2100 is shown in Figure 18.11. Figure 18.12 reflects these projections.

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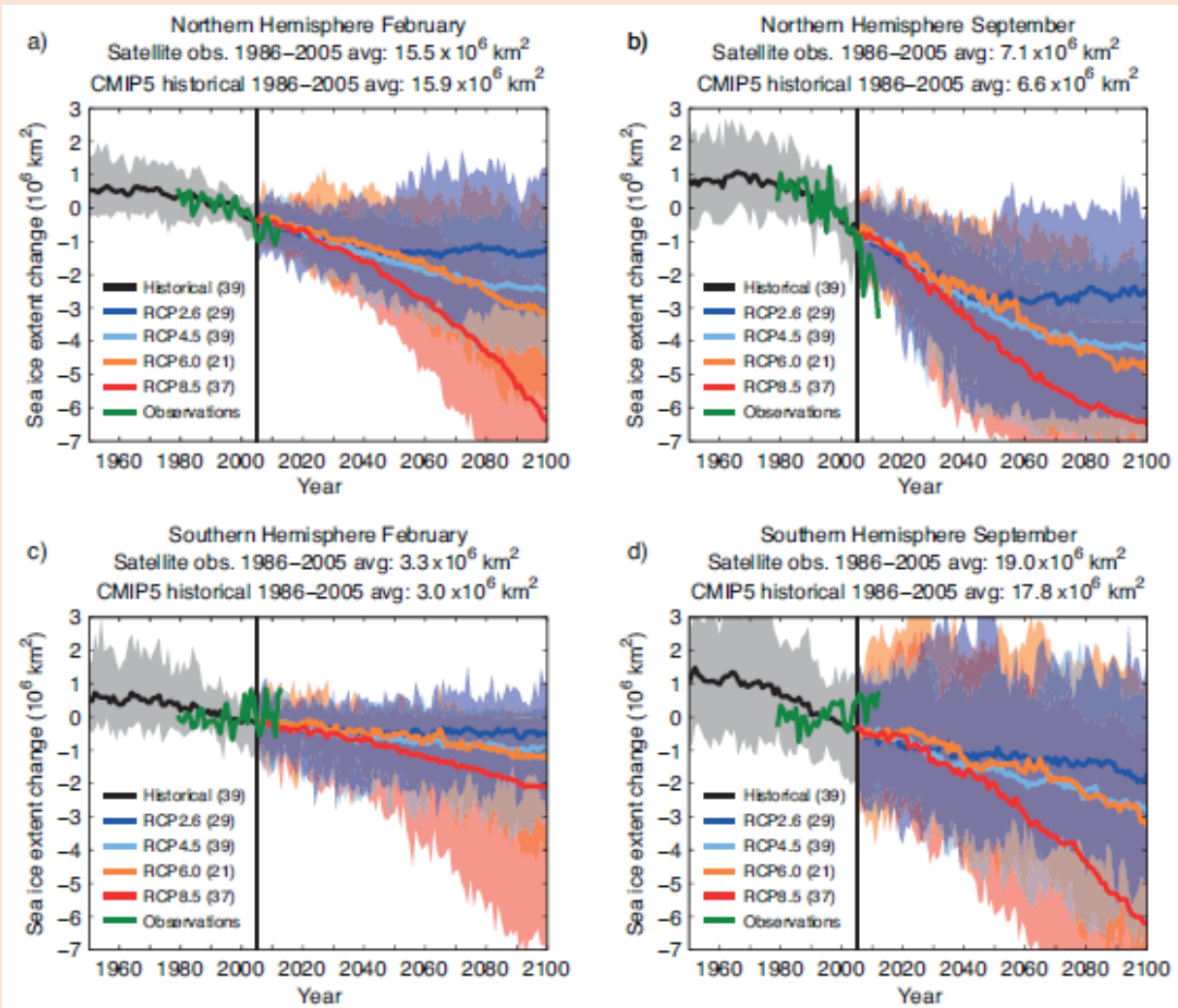


Figure 18.11 Sea ice extent northern hemisphere and southern hemisphere for February and September observed and projected.

[https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf)

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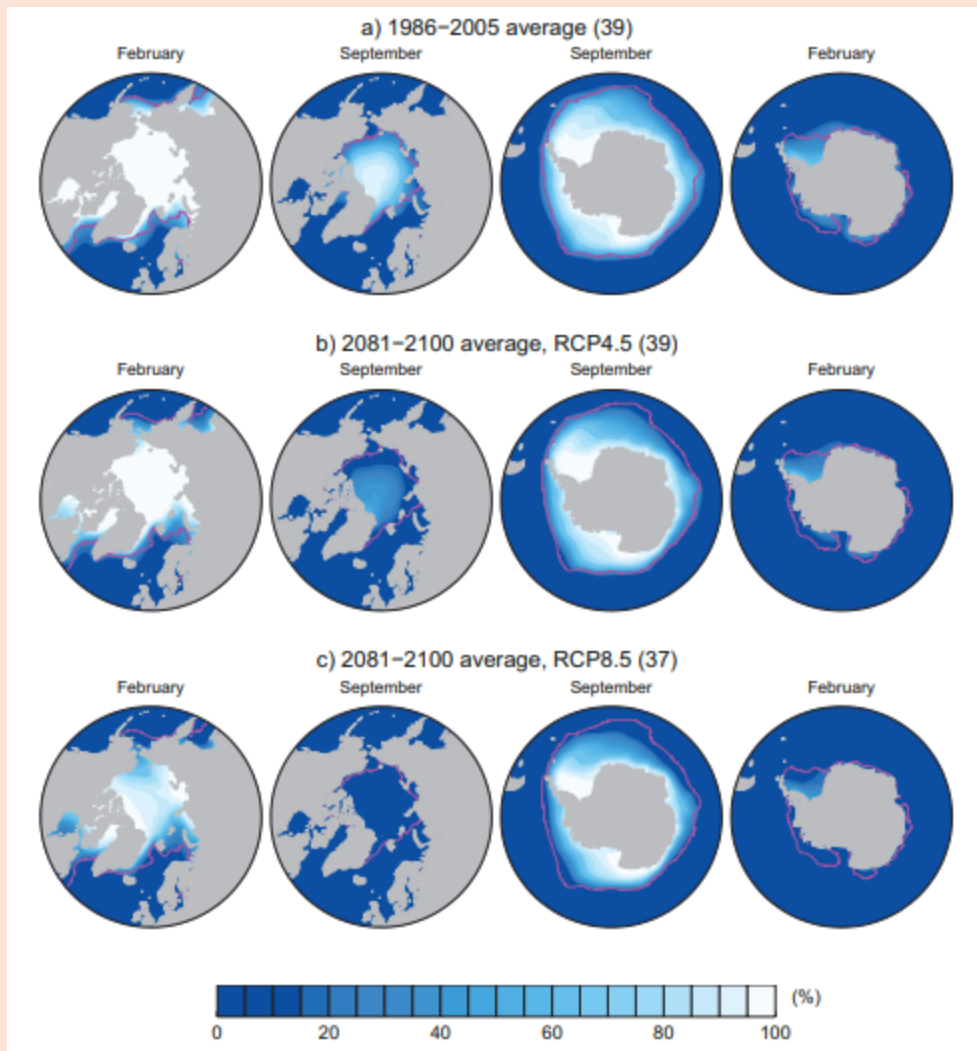


Figure 18.12 Sea ice concentrations for Arctic and Antarctic, 1986-2005 average, projected for February and September using RCP4.5 and 8.5.

[https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf)

Snow cover extent change and near surface permafrost area change are shown in Figures 18.13 and 18.14.

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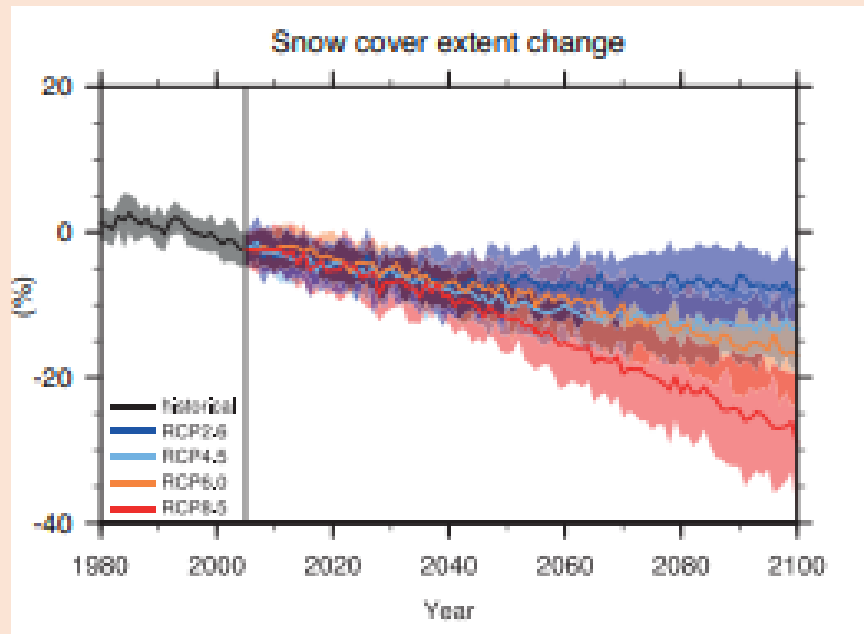


Figure 18.13 Snow cover extent range, historical and projected to 2100.  
[https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf)

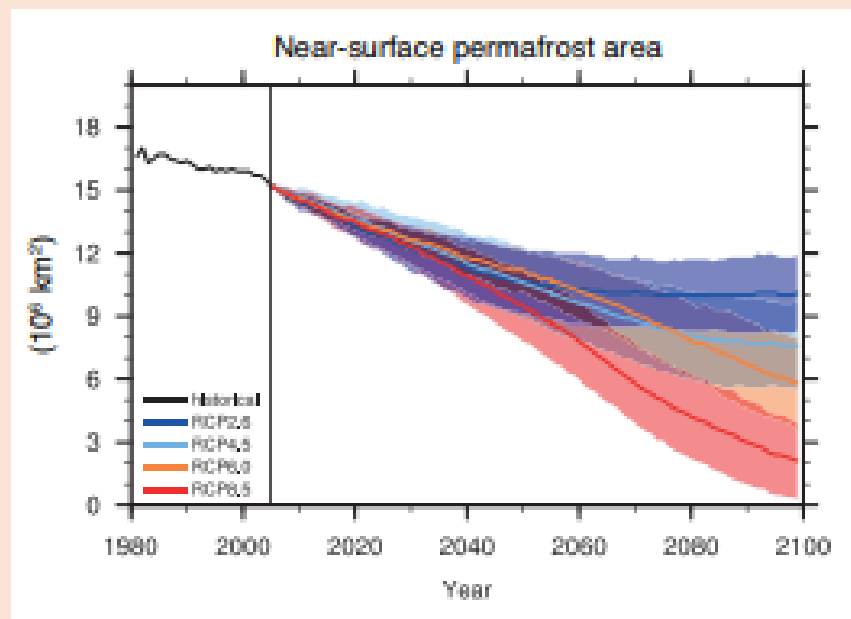


Figure 18.14 Near-surface permafrost area, historical and projected to 2100.  
[https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter12\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter12_FINAL.pdf)

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### 18.3.4 Sea level

The projected global mean sea level rise as result of thermal expansion, glacier melt, Greenland ice sheet melt and Antarctic ice sheet melt for RCPS 2.6, 4.5, 6.0 and 8.5 are shown in Figure 18.15.

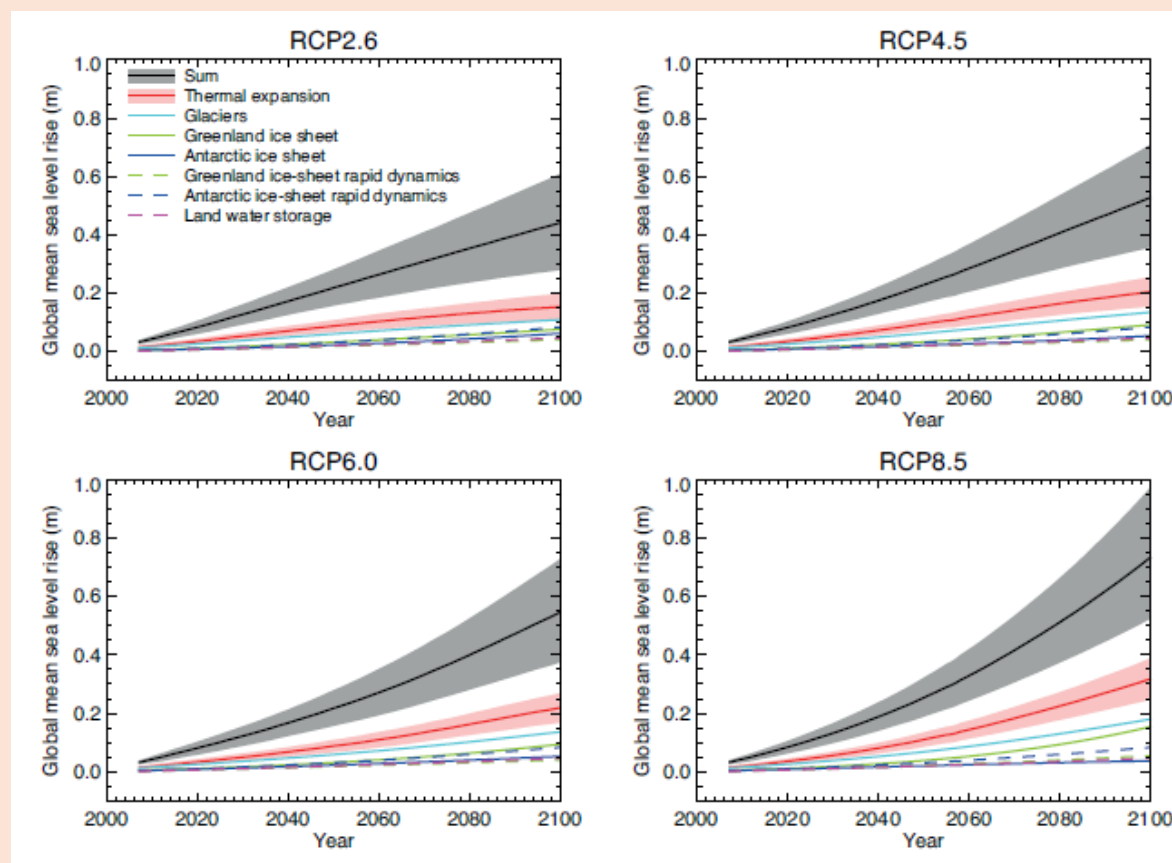


Figure 18.15 projected global sea level rise.

[https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter13\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter13_FINAL.pdf)

### 18.3.5 Ocean chemistry

Observed and projected changes in per cent change in ocean oxygen content, changes in oxygen concentration (200-600m) from 1990s to 2090s for RCP 2.6 and 8.5 are shown in Figure 18.16.

Historical and projected ocean pH and dissolved carbon dioxide to 2100 is shown in Figure 18.17. RCP is not specified but likely RCP8.5.

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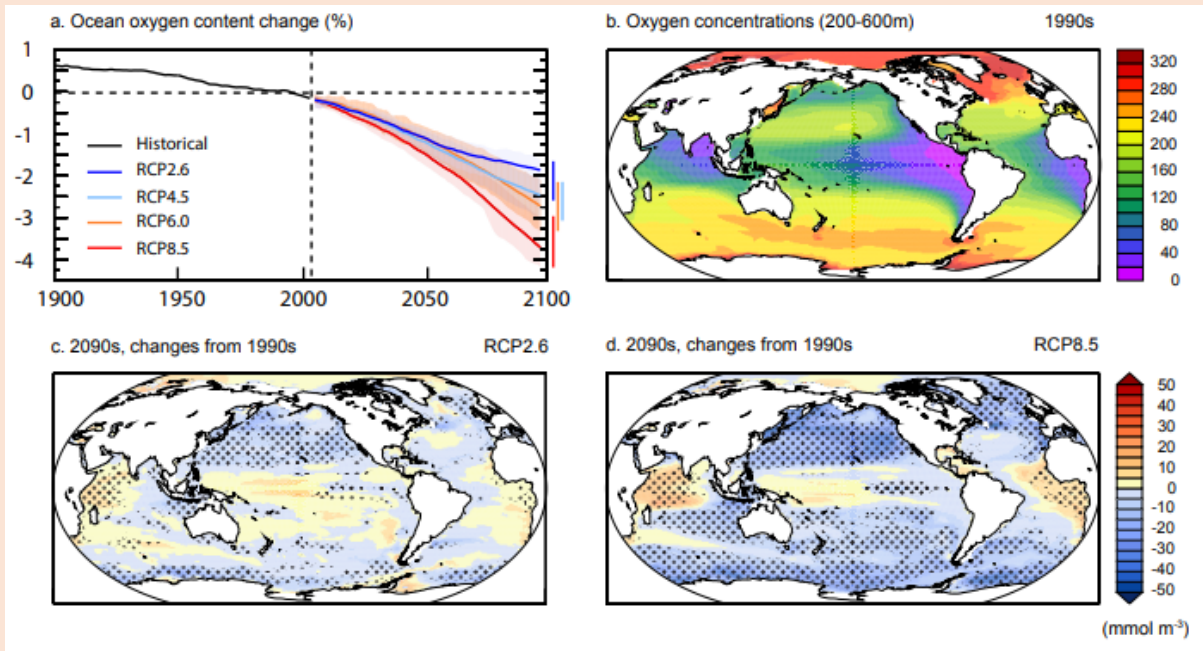


Figure 18.16 (a) Simulated changes in dissolved O<sub>2</sub> (mean and model range as shading) relative to 1990s for RCP2.6, RCP4.5, RCP6.0 and RCP8.5. (b) Multi-model means dissolved O<sub>2</sub> ( $\mu\text{mol m}^{-3}$ ) in the main thermocline (200 to 600 m depth average) for the 1990s, and changes in 2090s relative to 1990s for RCP2.6 (c) and RCP8.5 (d). To indicate consistency in the sign of change, regions are stippled where at least 80% of models agree on the sign of the mean change.

[https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter06\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter06_FINAL.pdf)

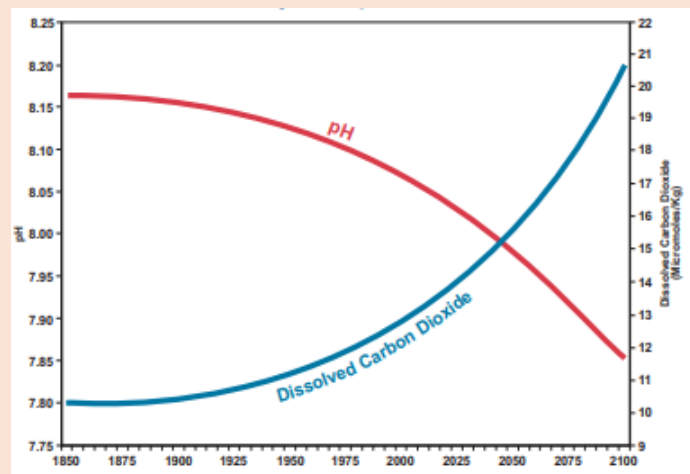


Figure 18.17 Historical and projected pH and dissolved CO<sub>2</sub>.

<https://www.pmel.noaa.gov/pubs/PDF/feel2899/feel2899.pdf>

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### 18.3.6 Jet stream

The impact of climate change on the jet stream is illustrated in Figure 18.18.

“As the globe continues to warm, it is already having an effect on the jet stream and corresponding weather patterns, according to the latest U.N. IPCC climate report, which states: ‘It is likely that circulation features have moved poleward since the 1970s, involving a widening of the tropical belt, a poleward shift of storm tracks and jet streams, and a contraction of the northern polar vortex. Evidence is more robust for the Northern Hemisphere.’ The research that goes into this statement comes from multiple lines of evidence – from analyses of the expansion of the tropical Hadley Cell to satellite measured outgoing radiation, radiosonde observations, and weather pattern reanalyses. But just as certainty builds for a poleward shifting jet, there still remain questions about whether the jet is amplifying and promoting more blocking patterns.” From Climate Central. **(Note that blocking patterns contribute to formation of heat domes.)**

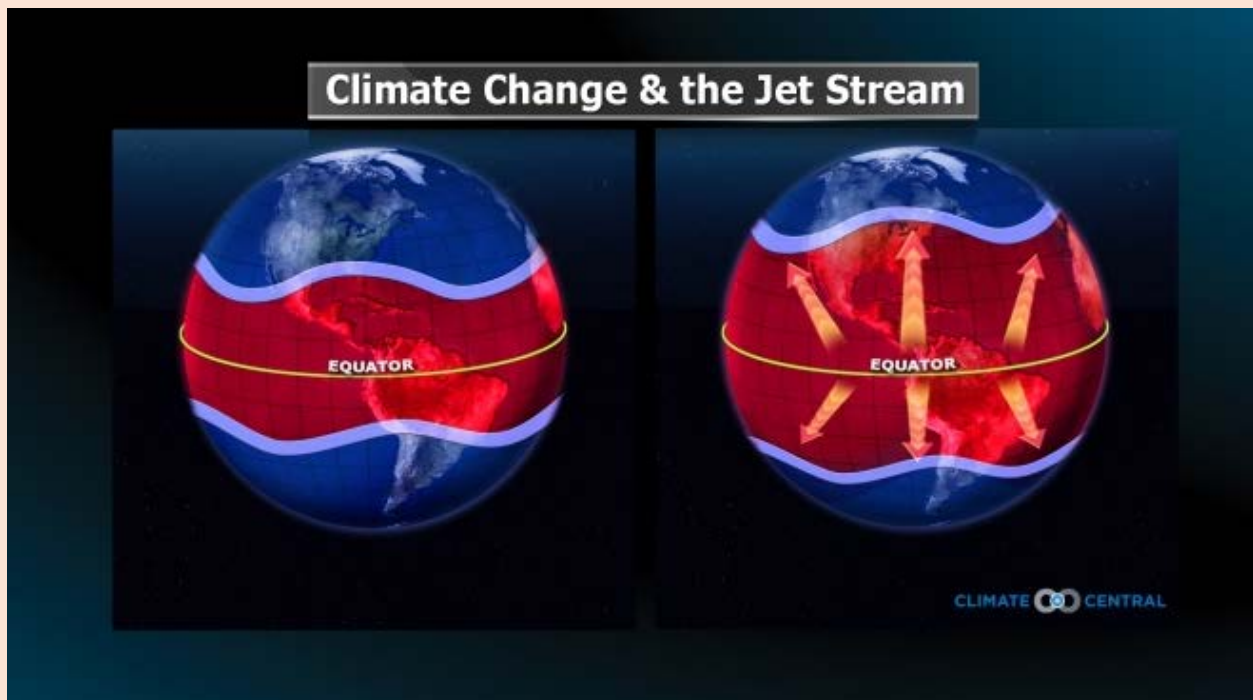


Figure 18.18 Climate change and the jet stream.

<http://www.climatecentral.org/gallery/graphics/climate-change-the-jet-stream>

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### 18.3.7 Atlas of global and regional climate projections

The atlas of global and regional climate projections is shown in Figure 18.19. This atlas may be accessed by going to Annex 1 of IPCC AR5 WG1, <https://www.ipcc.ch/report/ar5/wg1/>.

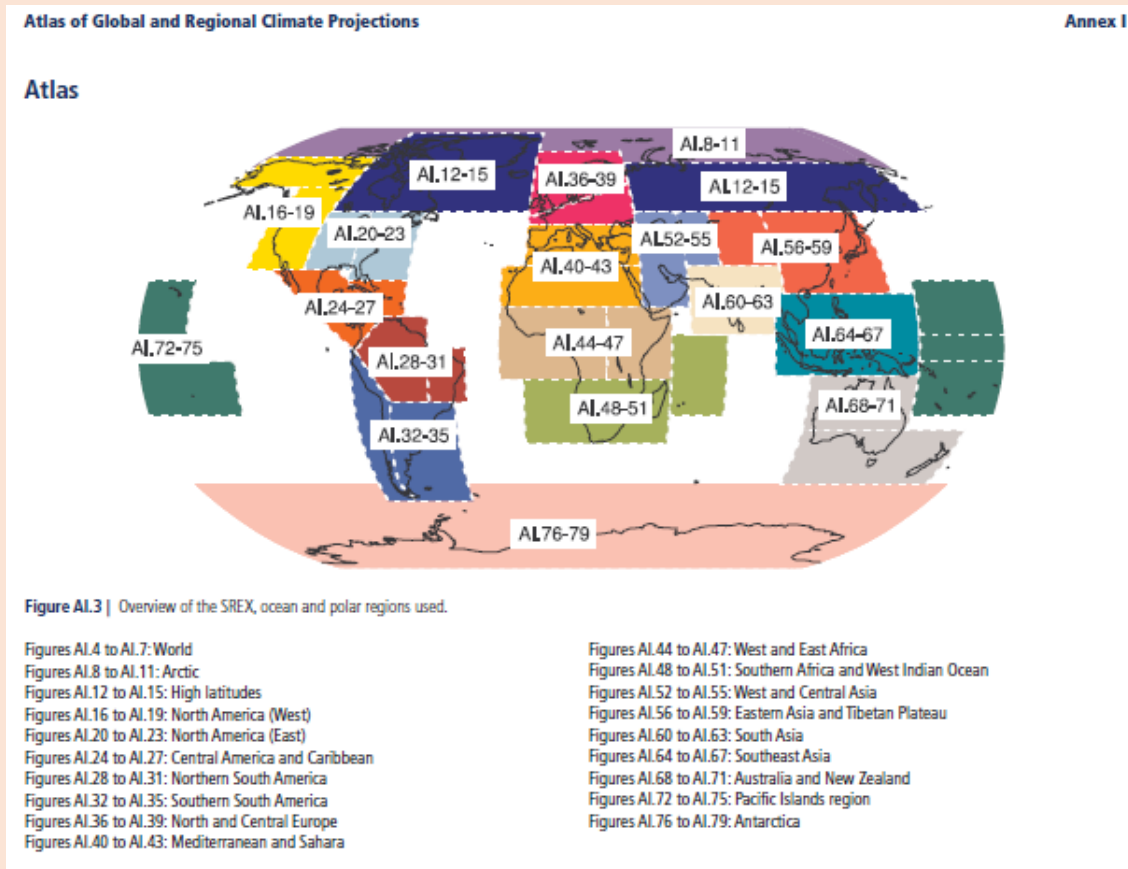


Figure 18.19 Annex 1 AR5 WG1 Atlas of global and regional climate projections.

<https://www.ipcc.ch/report/ar5/wg1/>

### 18.4 Climate phenomena and regional climate change

The IPCC considers the full scope of regional climate variation in Chapter 14 of AR5 WG1. To say the least, this is an exceptional undertaking because of the number and diversity of phenomena that affect regional climate change.

Table 18.4, taken from IPCC AR5 WG1, lists climate phenomena, which are called modes of climate variability. These include a variety of well-known oscillations and other climate patterns and behaviour. The regional climate impact for each mode is provided. The approach to

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analysis of the effects of world climate change on the modes and how regional climate change phenomena such as monsoons, large-scale storm systems (tropical and subtropical cyclones) are presented. Table 18.5 provides a region-by-region summary of the significance of the impacts.

The modes are important in themselves (ENSO and the annular and dipolar modes) or they affect climate systems such as monsoons, tropical phenomena (convergence zones and Atlantic Ocean modes), tropical cyclones and extratropical cyclones. Other important phenomena such as blocking systems and teleconnections that are only vaguely understood complicate the ability to associate world climate change to regional climate.

Mode	Regional Climate Impacts
ENSO	Global impact on interannual variability in global mean temperature. Influences severe weather and tropical cyclone activity worldwide. The diverse El Niño flavours present different teleconnection patterns that induce large impacts in numerous regions from polar to tropical latitudes (Section 14.4).
PDO	Influences surface air temperature and precipitation over the entire North American continent and extratropical North Pacific. Modulates ENSO rainfall teleconnections, e.g., Australian climate (Section 14.7.3).
IPO	Modulates decadal variability in Australian rainfall, and ENSO teleconnections to rainfall, surface temperature, river flow and flood risk over Australia, New Zealand and the SPCZ (Section 14.7.3).
NAO	Influences the N. Atlantic jet stream, storm tracks and blocking and thereby affects winter climate in over the N. Atlantic and surrounding landmasses. The summer NAO (SNAO) influences Western Europe and Mediterranean basin climates in the season (Section 14.5.1).
NAM	Modulates the intensity of mid-latitude storms throughout the Northern Hemisphere and thereby influences North America and Eurasia climates as well as sea ice distribution across the Arctic sea (Section 14.5.1).
NPO	Influences winter air temperature and precipitation over much of western North America as well as Arctic sea ice in the Pacific sector (Section 14.5.1).
SAM	Influences temperature over Antarctica, Australia, Argentina, Tasmania and the south of New Zealand and precipitation over southern South America, New Zealand, Tasmania, Australia and South Africa (Section 14.5.2).
PNA	Influences the jet stream and storm tracks over the Pacific and North American sectors, exerting notable influences on the temperature and precipitation in these regions on intraseasonal and interannual time scales (Section 14.7.2).
PSA	Influences atmospheric circulation over South America and thereby has impacts on precipitation over the continent (Section 14.7.1).
AMO	Influences air temperatures and rainfall over much of the Northern Hemisphere, in particular, North America and Europe. It is associated with multidecadal variations in Indian, East Asian and West African monsoons, the North African Sahel and northeast Brazil rainfall, the frequency of North American droughts and Atlantic hurricanes (Section 14.7.6).
AMM	Influences seasonal hurricane activity in the tropical Atlantic on both decadal and interannual time scales. Its variability is influenced by other modes, particularly ENSO and NAO (Section 14.3.4).
AN	Affects the West African Monsoon, the oceanic forcing of Sahel rainfall on both decadal and interannual time-scales and the spatial extension of drought in South Africa (Section 14.3.4).
IOB	Associated with the intensity of Northwest Pacific monsoon, the tropical cyclone activity over the Northwest Pacific and anomalous rainfall over East Asia (Section 14.3.3).
IOD	Associated with droughts in Indonesia, reduced rainfall over Australia, intensified Indian summer monsoon, floods in East Africa, hot summers over Japan, and anomalous climate in the extratropical Southern Hemisphere (Section 14.3.3).
TBO	Modulates the strength of the Indian and West Pacific monsoons. Affects droughts and floods over large areas of south Asia and Australia (Section 14.7.4).
MJO	Modulates the intensity of monsoon systems around the globe and tropical cyclone activity in the Indian, Pacific and Atlantic Oceans. Associated with enhanced rainfall in Western North America, northeast Brazil, Southeast Africa and Indonesia during boreal winter and Central America/Mexico and Southeast Asia during boreal summer (Section 14.3.2).
QBO	Strongly affects the strength of the northern stratospheric polar vortex as well as the extratropical troposphere circulation, occurring preferentially in boreal winter (Section 14.7.5).
BLC	Associated with cold air outbreaks, heat-waves, floods and droughts in middle and high latitudes of both hemispheres (Box 14.2).

Notes:

AMM: Atlantic Meridional Mode	IOB: Indian Ocean Basin pattern	NAO: North Atlantic Oscillation	QBO: Quasi-Biennial Oscillation
AMO: Atlantic Multi-decadal Oscillation	IOD: Indian Ocean Dipole pattern	NPO: North Pacific Oscillation	SAM: Southern Annular Mode
AN: Atlantic Niño pattern	IPO: Interdecadal Pacific Oscillation	PDO: Pacific Decadal Oscillation	TBO: Tropospheric Biennial Oscillation
BLC: Blocking events	MJO: Madden-Julian Oscillation	PNA: Pacific North America pattern	
ENSO: El Niño-Southern Oscillation	NAM: Northern Annular Mode	PSA: Pacific South America pattern	

Table 18.4 Regional climate impacts.

[file:///H:/Documents/My%20Stuff%20Book%20GW%20and%20CC/Temp/WG1AR5\\_Chapter14\\_FINAL.pdf](file:///H:/Documents/My%20Stuff%20Book%20GW%20and%20CC/Temp/WG1AR5_Chapter14_FINAL.pdf)

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**Table 14.3 |** Summary of the relevance of projected changes in major phenomena for mean change in future regional climate. The relevance is classified into high (red), medium (yellow), low (cyan), and 'no obvious relevance' (grey), based on confidence that there will be a change in the phenomena ('HP' for high, 'MP' for medium, 'LP' for low), and confidence in the impact of the phenomena on each region ('H1' for high, 'M1' for medium, 'L1' for low). More information on how these assessments have been constructed is given in the Supplementary Material (Section 14.SM.6.1).

Phenomena Regions	Section	Monsoon Systems <i>MP</i> —See Section 14.2	Tropical Phenomena* <i>HP/MP/LP/LP1P</i> —See Section 14.3	ENSO <i>LP</i> —See Section 14.4	Annular and Dipolar Modes <i>HP</i> —See Section 14.5	Tropical Cyclones <i>MP</i> —See Section 14.6.1	Extratropical Cyclones* <i>MP/HP</i> —See Section 14.6.2
Arctic	14.8.2				<i>HP/HP1</i> The small projected increase in NAO is likely to contribute to wintertime changes in temperature and precipitation.		<i>MP/HP1</i> Projected increase in precipitation in extratropical cyclones is likely to enhance mean precipitation.
North America	14.8.3	<i>MP/HP1</i> It is likely the number of consecutive dry days will increase, and overall water availability will be reduced.	<i>HP/L1</i> Projected ITCZ shifts unrelated to ENSO changes will impact temperature and precipitation, especially in winter.	<i>LP/HP1</i> Likely changes in N. American precipitation if ENSO changes.	<i>HP/MP1</i> The small projected increase in the NAO index is likely to contribute to wintertime temperature and precipitation changes in NE America.	<i>MP/HP1</i> Projected increases in extreme precipitation near the centres of tropical cyclones making landfall along the western coast of the USA and Mexico, the Gulf Mexico, and the eastern coast of the USA and Canada.	<i>MP/HP1</i> Projected increases in precipitation in extratropical cyclones will lead to large increases in wintertime precipitation over the northern third of the continent.
Central America and Caribbean	14.8.4	<i>MP/HP1</i> Projected reduction in mean precipitation.	<i>HP/HP1</i> Reduced mean precipitation in southern Central America if there is a southward displacement of the East Pacific ITCZ.	<i>LP/HP1</i> Reduced mean precipitation if El Niño events become more frequent and/or intense.		<i>MP/HP1</i> More extreme precipitation near the centres of tropical cyclones making landfall along the eastern and western coasts.	
South America	14.8.5	<i>MP/HP1</i> Projected increase in extreme precipitation and in the extension of monsoon area.	<i>HP/HP1</i> Projected increase in the mean precipitation in the southeast due to the projected southward displacement of the SACZ.	<i>LP/HP1</i> Reduced mean precipitation in eastern Amazonia and increased precipitation in the La Plata Basin.	<i>HP/HP1</i> Poleward shift of storm tracks due to projected positive trend in SAMs phase leads to less precipitation in central Chile and increased precipitation in the southern tip of South America.		<i>HP/HP1</i> Southward displacement of cyclogenesis activity increases the precipitation in the extreme south.
Europe and Mediterranean	14.8.6				<i>HP/HP1</i> Projected increase in the NAO will lead to enhanced winter warming and precipitation over NW Europe.		<i>MP/HP1</i> Enhanced extremes of storm-related precipitation and decreased frequency of storm-related precipitation over the E. Mediterranean.
Africa	14.8.7	<i>MP/HP1</i> Projected enhancement of summer precipitation in West Africa.	<i>HP/L1</i> Enhanced precipitation in parts of East Africa due to projected shifts in ITCZ. Modified precipitation in West or East Africa according to variations in Atlantic or Indian Ocean SSTs.		<i>HP/HP1</i> Enhanced winter warming over southern Africa due to projected increase in SAM.	<i>MP/HP1</i> Projected increase in extreme precipitation near the centres of tropical cyclones making landfall along the eastern coast (including Madagascar).	<i>HP/HP1</i> Enhanced extremes of storm-related precipitation and decreased frequency of storm-related precipitation over southwestern Africa.
Central and North Asia	14.8.8	<i>MP/MP1</i> Projected enhancement in summer mean precipitation.			<i>HP/L1</i> Projected enhancement in winter warming over North Asia.		
East Asia	14.8.9	<i>MP/MP1</i> Enhanced summer precipitation due to intensification of East Asian summer monsoon circulation.		<i>LP/HP1</i> Enhanced warming if El Niño events become more frequent and/or intense.		<i>MP/HP1</i> Projected increase in extreme precipitation near the centres of tropical cyclones making landfall in Japan, along coasts of east China Sea and Sea of Japan.	<i>MP/MP1</i> Projected reduction in midwinter precipitation.

Table 18.5 Summary of the relevance of projected changes in major phenomena for mean change in regional climate.  
[https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter14\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter14_FINAL.pdf)

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Phenomena Regions	Section	Monsoon Systems MP—see Section 14.2	Tropical Phenomena <sup>a</sup> HP/HP/LP/LP—See Section 14.3	ENSO LP—See Section 14.4	Annular and Dipolar Modes HP—See Section 14.5	Tropical Cyclones MP—See Section 14.6.1	Extratropical Cyclones <sup>b</sup> MP/HP—See Section 14.6.2
West Asia	14.8.10		HP/LJ Enhanced precipitation in southern parts of West Asia due to projected northward shift in ITCZ.			MP/HP Projected increase in extreme precipitation near the centres of tropical cyclones making landfall on the Arabian Peninsula.	MP/LJ Projected decrease in mean precipitation due to northward shift of storm tracks.
South Asia	14.8.11	MP/MJ Enhanced summer precipitation associated with Indian Monsoon.	LP/MJ Strengthened break monsoon precipitation anomalies associated with MJO.	LP/HP Increased warming and increased summer season rainfall variability due to ENSO.		MP/HP Projected increase in extreme precipitation near the centres of tropical cyclones making landfall along coasts of Bay of Bengal and Arabian Sea.	
South-east Asia	14.8.12	LP/MJ Decrease in precipitation over Maritime continent.	HP/MJ Projected changes in IOD-like warming pattern will reduce mean precipitation in Indonesia during Jul–Oct.	LP/HP Reduction in mean precipitation and enhanced warming if El Niño events become more frequent and/or intense.		MP/HP Projected increase in extreme precipitation near the centres of tropical cyclones making landfall along coasts of South China Sea, Gulf of Thailand, and Andaman Sea.	
Australia and New Zealand	14.8.13	MP/LJ Mean monsoon precipitation may increase over northern Australia.	HP/LJ More frequent zonal SPZ episodes may reduce precipitation in NE Australia.	LP/HP Reduced precipitation in North and East Australia and NZ if El Niño events become more frequent and/or intense.	HP/MJ Increased warming and reduced precipitation in NZ and South Aust. due to projected positive trend in SAM.	MP/HP More extreme precipitation near the centres of tropical cyclones making landfall along the eastern, western, and northern coasts of Australia.	HP/HP Projected increase in extremes of storm-related precipitation.
Pacific Islands Region	14.8.14		HP/LJ Increased mean precipitation along equator with ITCZ intensification. More frequent zonal SPZ episodes leading to reduced precipitation in southwest and increases in east.	HP/LJ Increased mean precipitation in central/east Pacific if El Niño events become more frequent and/or intense.		HP/HP More extreme precipitation near the centres of tropical cyclones passing over or near Pacific Islands.	
Antarctica	14.8.15			LP/MJ Increased warming over Antarctic Peninsula and west reduced across central Pacific if El Niño events become more frequent and/or intense.	HP/HP Increased warming over Antarctic Peninsula and west trend projected in SAM.		HP/MJ Increased precipitation in coastal areas due to projected poleward shift of storm track.

Table 18.5 Summary of the relevance of projected changes in major phenomena for mean change in regional climate.

[https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5\\_Chapter14\\_FINAL.pdf](https://www.ipcc.ch/site/assets/uploads/2018/02/WG1AR5_Chapter14_FINAL.pdf)

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## 18.5 Tipping points, domino effects, knock-on effects, runaway global warming and hothouse Earth

'Tipping point' is a metaphor that identifies a process where an otherwise small change in an input, that would normally have little effect on an outcome, results in a disproportionate response. A familiar example is 'the straw that broke the camel's back'. All straws loaded on the camel were accommodated but the last straw resulted in catastrophic consequences for the camel (broken back). The phrase, 'this is the last straw' is derived from the 'camel's back' metaphor. Consequences of an input were minor until a point was reached where significant consequences resulted from similar minor inputs. The social statement, 'one bad behavior too many results in the loss of significant privileges' is another good example of 'tipping point'. Phrases such 'one time too many' and 'crossed the line' come to mind. A tipping point is the condition where a small change in a factor that historically had seemingly minor influences results in a much larger important response. It is very important to identify tipping points so as to avoid the critical input that results the disproportionate response.

A domino effect' occurs when one process triggers another. (e. g. Climate change causes drought which results in loss of crops and water and food for livestock which results in starvation and loss of livelihood which results in mass migration.)

'Knock-on' consequences are similar to the 'domino effect', when one significant change results in the significant change in another phenomena or process. This is particularly important when the significant response results from a tipping point being exceeded. An example is greenhouse gases in the atmosphere causing climate change that results in lower rainfall which in turn causes drought conditions that in turn cause increased potential for wildfires. Another important example is loss of habitat caused by temperature increases, such as change in vegetation. Animals depending on the vegetation might disappear which might possibly result in the disappearance of the predators depending on the presence of the plant eating animals.

Knock-on effects can result in an additional increase to the input that originally caused the tipping point to be exceeded in the first place. The apparently amplified input can result in even greater consequences further strengthening the magnitude of the input. The phrase, 'out of control' comes to mind. This phenomenon is 'out of control' until it is constrained by other critical inputs or conditions that might not be available or the process simply breaks down. An example is global warming caused by greenhouse gases resulting in the accelerated thawing of the permafrost which then releases methane, a greenhouse gas that has twenty times the effect of carbon dioxide, that results in increased warming and so on.

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The discussion provided by Wikipedia is particularly relevant, [https://en.wikipedia.org/wiki/Tipping\\_points\\_in\\_the\\_climate\\_system#Cascading\\_tipping\\_points](https://en.wikipedia.org/wiki/Tipping_points_in_the_climate_system#Cascading_tipping_points).

In the context of climate change the critical input is typically the increase in greenhouse gas concentration in the atmosphere that results in the temperature change. 'Runaway' global warming occurs when global warming becomes beyond our control. 'Hothouse Earth' is the likely result. (<https://www.climateemergencyinstitute.com/runaway> , <https://www.pnas.org/doi/10.1073/pnas.1810141115> and [https://en.wikipedia.org/wiki/Runaway\\_greenhouse\\_effect#:~:text=A%20runaway%20greenhouse%20effect%20occurs,liquid%20water%20on%20its%20surface.](https://en.wikipedia.org/wiki/Runaway_greenhouse_effect#:~:text=A%20runaway%20greenhouse%20effect%20occurs,liquid%20water%20on%20its%20surface.) ) Mass extinction events would occur.

There are several important examples identified in the paper titled, Tipping Elements in the Earth's Climate System, <https://www.pnas.org/content/105/6/1786> which includes a description of the authors' proposed methodology for identifying tipping points. They specifically consider:

1. Loss of Arctic Sea ice.
2. Depletion/ loss of Greenland ice sheet.
3. Loss of Antarctic ice sheet.
4. Alteration of the Atlantic thermohaline circulation.
5. Alteration of the El Nino-Southern Oscillation (ENSO).
6. Indian Summer Monsoon (ISM).
7. Sahara/ Sahel and West African Monsoon (WSM).
8. Environment over the Amazon Rainforest.
9. Environment over the boreal forest.
10. Loss of permafrost.

Other phenomenon of relevance which are not specifically addressed include:

11. Warming of the sea water and fresh water.
12. Change in the acidity of oceans.
13. Loss of glaciers.
14. Alteration of any of the other ocean and atmospheric oscillations that might be impacted by global warming.

and, pretty well all of the physical impacts caused by global warming discussed in this chapter.

The ability to identify tipping points is important from the perspective of mitigating climate change to the extent that these tipping points are not reached.

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The newspaper, 'The Guardian' has recently published an article titled, 'World on brink of five 'disastrous' climate tipping points'  
[https://www.theguardian.com/environment/2022/sep/08/world-on-brink-five-climate-tipping-points-study-finds?utm\\_campaign=Carbon%20Brief%20Daily%20Briefing&utm\\_content=20220909&utm\\_medium=email&utm\\_source=Revue%20Daily](https://www.theguardian.com/environment/2022/sep/08/world-on-brink-five-climate-tipping-points-study-finds?utm_campaign=Carbon%20Brief%20Daily%20Briefing&utm_content=20220909&utm_medium=email&utm_source=Revue%20Daily) which summarizes a paper of the same name published in the journal, Science <https://www.science.org/doi/10.1126/science.abn7950> . The research indicates that even passing global warming of 1°C presents significant risk. They show a graph of the risk of climate tipping points as shown in Figure 18.20.

This study is supported by another paper titled, 'Risk of multiple climate tipping points escalates above 1.5°C global warming' published by Stockholm University <https://phys.org/news/2022-09-multiple-climate-escalates-15c-global.html> . They provide a global map showing the location of climate tipping elements as provided in Figure 18.21.

The Potsdam Institute for Climate Impact Research provides a discussion on how 'Tipping elements can destabilize each other, leading to climate domino effects'  
<https://phys.org/news/2021-06-elements-destabilize-climate-domino-effects.html> .

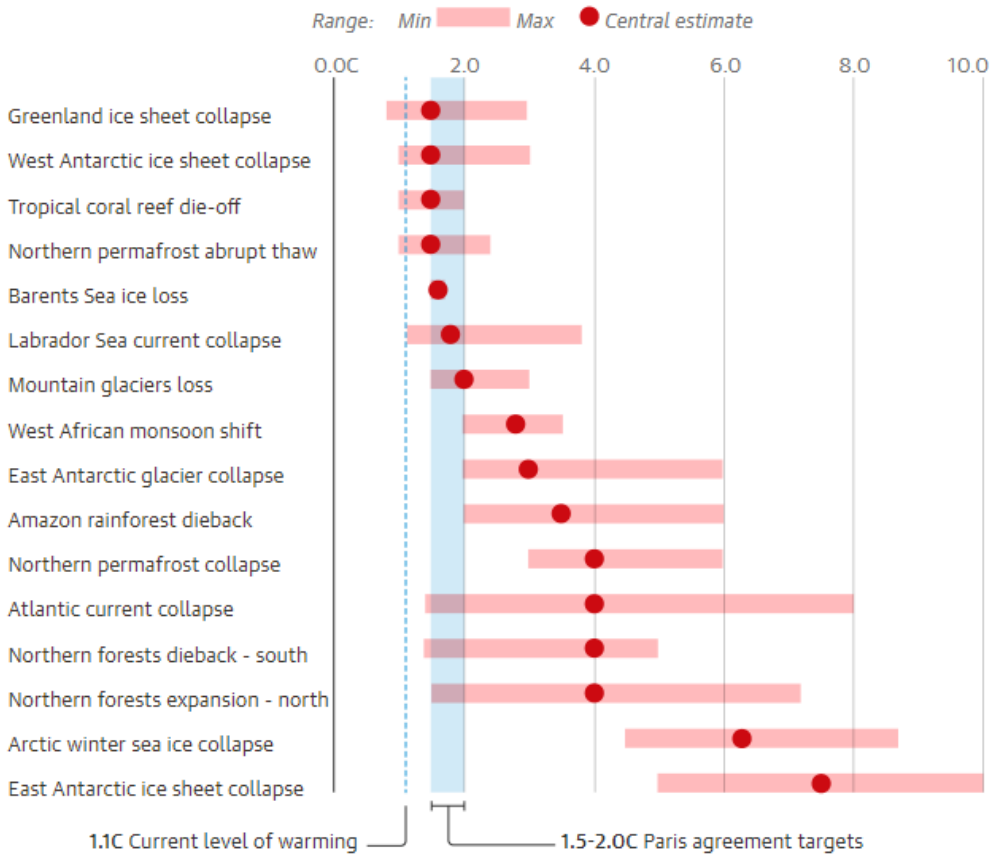
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## The risk of climate tipping points is rising rapidly as the world heats up

Estimated range of global heating needed to pass tipping point temperature



Guardian graphic. Source: Armstrong McKay et al, Science, 2022. Note: Current global heating temperature rise 1.1C Paris agreement targets 1.5-2.0C

Figure 18.20 The risk of climate tipping points

<https://www.science.org/doi/10.1126/science.abn7950> .

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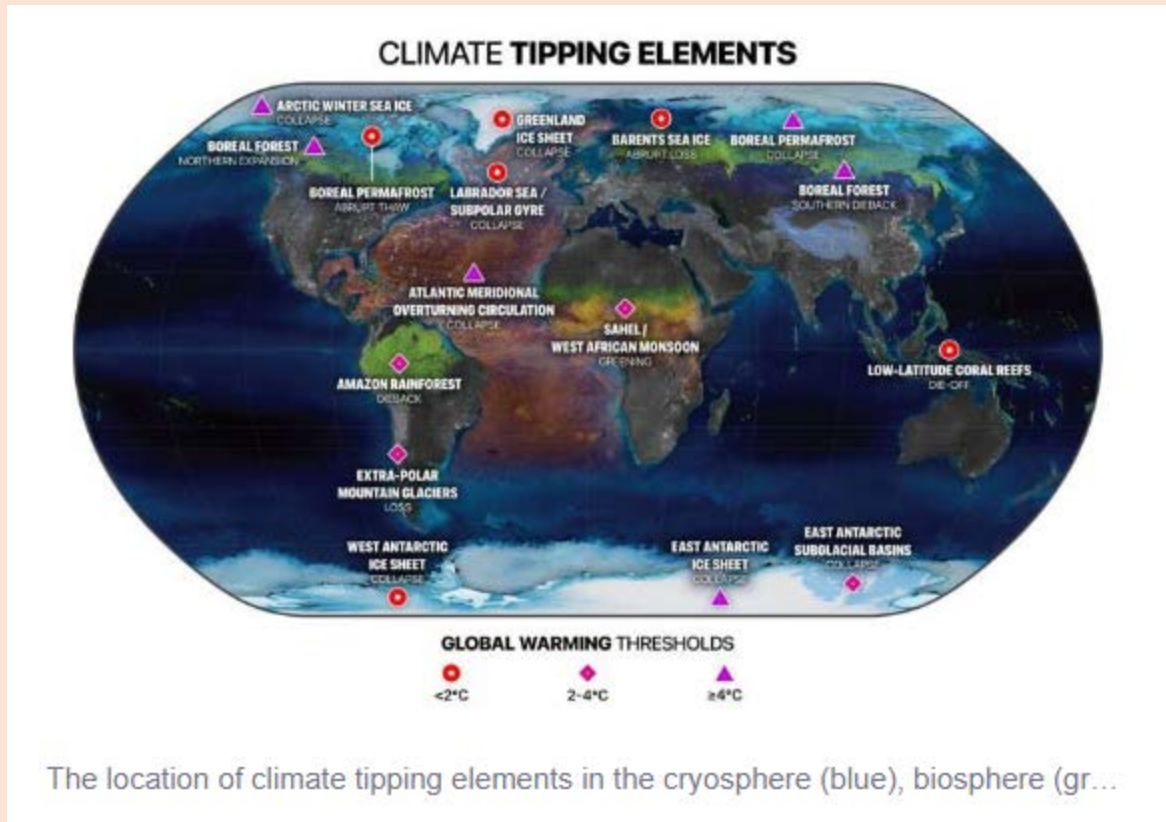


Figure 18.21 Location of climate tipping elements <https://phys.org/news/2022-09-multiple-climate-escalates-15c-global.html>

A conference titled ‘Tipping points from climate crisis to positive transformation’ was held 12<sup>th</sup> to 14<sup>th</sup> September, 2022 at the University of Exeter, Exeter UK <https://global-tipping-points.org/>. A summary of the conference outcomes pertaining to discussion of the concept of tipping points (animation is particularly informative), interacting tipping points and positive tipping points, may be found in [https://www.carbonbrief.org/tipping-points-how-could-they-shape-the-worlds-response-to-climate-change/?utm\\_campaign=Cropped&utm\\_content=20220921&utm\\_medium=email&utm\\_source=Revue%20Land](https://www.carbonbrief.org/tipping-points-how-could-they-shape-the-worlds-response-to-climate-change/?utm_campaign=Cropped&utm_content=20220921&utm_medium=email&utm_source=Revue%20Land). The final session dealt with the challenge of using information on tipping points, positive tipping points in particular, to influence policy to meet climate goals.

### 18.6 Snowmelt from mountainous regions

The volume of snow melt and the manner with which it occurs (timing of the melt) has been predictable – particularly from mountainous regions. For this reason, humans have developed a

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dependency on the mountain snow melt for domestic, industrial and agricultural use. In some regions of the world mountain snow melt is captured using dams and reservoirs so that it can be gradually used for hydroelectric power generation and for irrigation purposes or in times when the available snow melt in rivers and streams would be insufficient for domestic, municipal and industrial use.

The impact of global warming and climate change on runoff from snow accumulations is a decrease in volume of snow melt and also the earlier melting of snow resulting in historically lower flows and water shortages later in the year. See discussion paper [https://www.carbonbrief.org/climate-change-has-driven-16-drop-in-snow-meltwater-from-asias-high-mountains?utm\\_campaign=Carbon%20Brief%20Daily%20Briefing&utm\\_content=20210625&utm\\_medium=email&utm\\_source=Revue%20Daily](https://www.carbonbrief.org/climate-change-has-driven-16-drop-in-snow-meltwater-from-asias-high-mountains?utm_campaign=Carbon%20Brief%20Daily%20Briefing&utm_content=20210625&utm_medium=email&utm_source=Revue%20Daily).

### 18.7 Atmospheric rivers

Atmospheric rivers are discussed in Chapter 7.

A NASA-led study, <https://climate.nasa.gov/news/2740/climate-change-may-lead-to-bigger-atmospheric-rivers/>, ‘shows that climate change is likely to intensify extreme weather events known as atmospheric rivers across most of the globe by the end of this century, while slightly reducing their number.’

The relationship between climate change and the characteristics of atmospheric rivers is not well established.

### 18.8 Water management infrastructure

Changes in patterns of precipitation, not just increases or decreases in annual volumes, but also timing of the occurrence of normally expected rainfall events such as monsoons or ‘rainy seasons’, can result in water supply shortfalls and flooding. Increases and decreases in snowfall accumulation and melt patterns can have similar affects. Changes in temperature together with changes in seasonal precipitation events may negatively impact dryland agriculture and pasture for livestock (drought conditions) and place increased demand on water supplies for domestic use, industrial use, irrigation and hydropower generation. Flooding will result from unexpected runoff timing, flow rate and volume. Groundwater might be depleted to where it is no longer a useable water supply.

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Water supply management and flood control infrastructure which communities have invested considerable resources and have come to depend on may become inadequate. Large scale water management infrastructure such as those using dams to create reservoirs for capture and release of seasonal precipitation, snow or rainfall, may no longer be able to meet historical demands which they were intended to satisfy. This includes domestic water needs, agricultural uses, power generation and industrial needs. Examples of projects that will be significantly affected are those relying on the Aswan Dam, Egypt, the Kariba Dam, Zimbabwe and Zambia, all dams and reservoirs on the Columbia River System in the United States and the recently constructed Grand Ethiopian Renaissance Dam in Ethiopia. All projects of these types, large or small worldwide, will be similarly impacted.

### 18.9 Net-zero by 2050 objective and committed warming

The concepts of ‘net zero’ and ‘committed warming’ are actually part of the discussion of emission scenarios, namely, representative concentration pathways, (RCPs) which may be found in Section 18.2 and shared socio-economic pathways, (SSPs) which may be found in Section 21.3. The outcomes of following these emission scenarios are determined using various climate change models. The net zero objective is different in that it is simply the emission target for 2050. There is no attempt to describe the emission pathway to achieve this except to note that if this objective is achieved, the global temperature in year 2050 will not increase. (Intermediate emission targets for 2030 are encouraged to help ensure that the 2050 target can be achieved.) This has caused some confusion with all parties, governments at all levels, businesses and all other human activities who seem to be left to their own resources as to how they will achieve this objective (not quite true since governments at all levels are closely involved providing guidelines and incentives). The net zero objective is quite simple though and all parties are expected to achieve the objective independently following their own pathway. All stakeholders (parties) must be conversant in all aspects of greenhouse gas emissions as it relates to their activities and responsibilities. Their progress is reported and collated at a national level and so their success in achieving the 2050 objective can be monitored. A brief discussion of net zero and committed warming follows.

The UNFCCC Paris Agreement (<https://unfccc.int/process-and-meetings/the-paris-agreement/the-paris-agreement>), adopted by 196 Parties at COP21 in Paris on December 12 2015 and entered into force on November 4, 2016 committed the Parties (countries) to reducing their GHG emissions sufficiently to limit global warming to well below 2 °C. and preferably 1.5°C.

The IPCC Special Report, Global Warming of 1.5 °C, outlines the impacts of global warming above this temperature and the pathways with which this may be achieved

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<https://www.ipcc.ch/sr15/>. The pathways with which the 1.5°C objective for 2100 can be achieved is discussed in the In-depth Q&A article in Carbon Brief 08.10.2018 <https://www.carbonbrief.org/in-depth-qa-ipccs-special-report-on-climate-change-at-one-point-five-c>. Some involve ‘over-shoot’ of warming before 2100 but achieve 1.5°C by 2100 and others that do not over-shoot the temperature objective before 2100. Only those scenarios which did not over-shoot were considered acceptable; that is, aligned with the Paris Agreement’s real objective of 1.5°C limit for 2100.

The ‘net-zero’ emission objective is the commitment by all countries to net-zero CO<sub>2</sub> emissions by 2050. If this were achieved the global temperature at 2050 would gradually decrease from this year forward. The underlying hope is that the global temperature in 2050 will be less than 2°C and close to 1.5°C.

The ‘Explainer article’ in the newsletter, Carbon Brief, by Zeke Hausfather, 29.04.2021 that ‘The best available evidence shows that – warming is likely to more or less stop, that is, no over-shoot, once carbon dioxide emissions reach zero’, [https://www.carbonbrief.org/explainer-will-global-warming-stop-as-soon-as-net-zero-emissions-are-reached?utm\\_campaign=Daily%20Briefing&utm\\_content=20220224&utm\\_medium=email&utm\\_source=Revue%20newsletter](https://www.carbonbrief.org/explainer-will-global-warming-stop-as-soon-as-net-zero-emissions-are-reached?utm_campaign=Daily%20Briefing&utm_content=20220224&utm_medium=email&utm_source=Revue%20newsletter). The author explains the difference in surface temperature between scenarios which maintain a constant concentration of CO<sub>2</sub> and the effect of achieving net-zero. This is illustrated in Figure 18.22. Additional warming would occur, (over-shoot), if the concentration of CO<sub>2</sub> was held constant after 2050, that is until the concentrations of CO<sub>2</sub> in the atmosphere, land and the ocean achieved equilibrium.

If CO<sub>2</sub> emissions decreased as per the net-zero emission scenario, the concentration of CO<sub>2</sub> in the atmosphere would decrease as a result of the CO<sub>2</sub> loss to the land and oceans. It is interesting to note that the IPCC 1.5°C objective for 2100, <https://www.ipcc.ch/sr15/>, can be achieved in a variety of other ways as discussed in the In-depth Q&A article in Carbon Brief 08.10.2018 <https://www.carbonbrief.org/in-depth-qa-ipccs-special-report-on-climate-change-at-one-point-five-c> that involve ‘over-shoot’ of warming before 2100 but achieve 1.5°C by 2100 and those that do not over-shoot the temperature objective before 2100. See Figure 18.23. Only those scenarios which did not over-shoot were considered acceptable; that is, aligned with the Paris Agreement’s 1.5°C limit for 2100.

It is important to emphasize that net-zero CO<sub>2</sub> will likely achieve the IPCC 1.5°C objective for 2100. The success of net-zero was determined using the models discussed in Chapter 17. The ways with which net-zero can be achieved are determined by individual governments.

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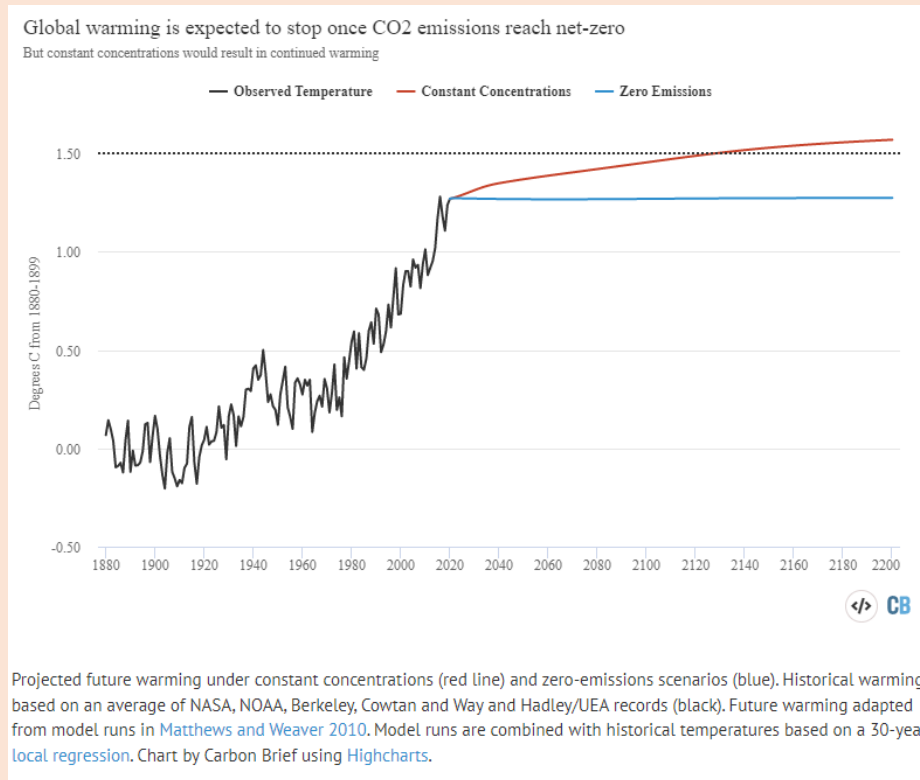


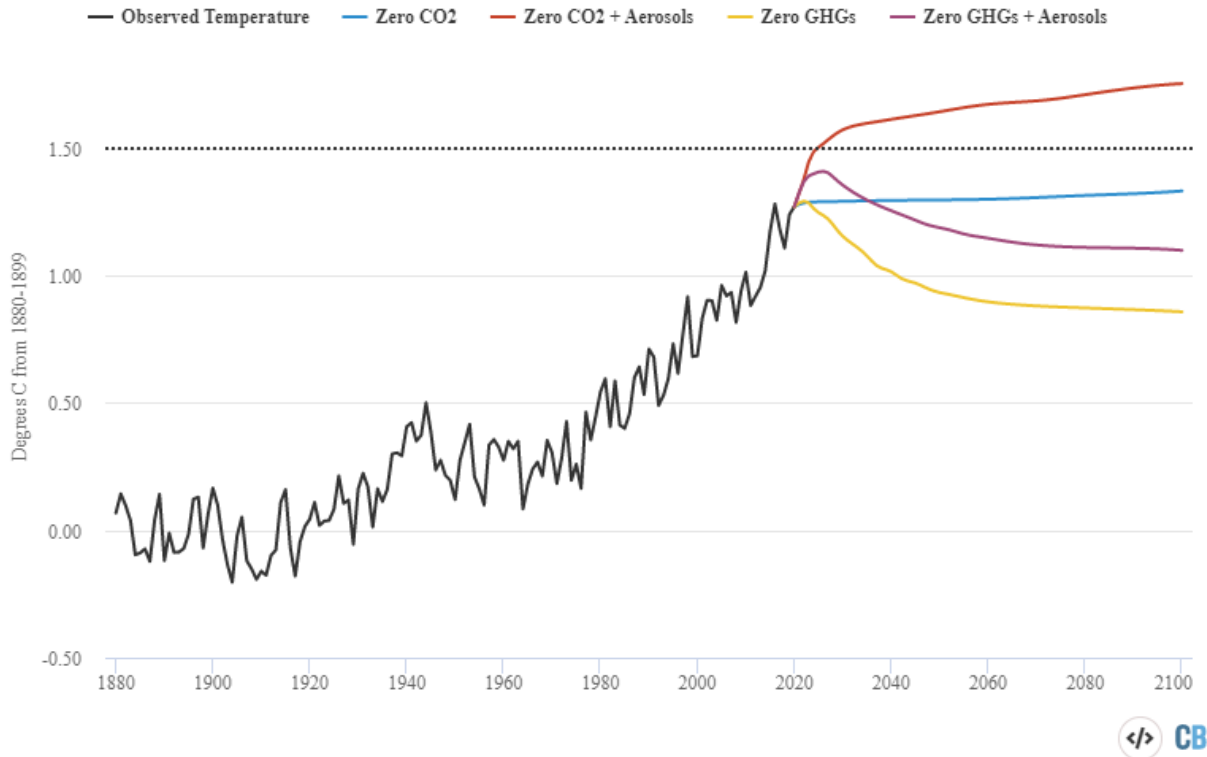
Figure 18.22 Committed temperature increases under the constant concentration scenario of CO<sub>2</sub> in the atmosphere and the zero emissions scenario [https://www.carbonbrief.org/explainer-will-global-warming-stop-as-soon-as-net-zero-emissions-are-reached?utm\\_campaign=Daily%20Briefing&utm\\_content=20220224&utm\\_medium=email&utm\\_source=Revue%20newsletter](https://www.carbonbrief.org/explainer-will-global-warming-stop-as-soon-as-net-zero-emissions-are-reached?utm_campaign=Daily%20Briefing&utm_content=20220224&utm_medium=email&utm_source=Revue%20newsletter) .

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### Future warming under different zero-emissions scenarios



Projected global surface temperature changes under zero CO<sub>2</sub> emissions (blue line), zero CO<sub>2</sub> and aerosol emissions (red), zero GHG emissions (yellow) and zero GHG and aerosol emissions (purple). Chart by Carbon Brief using Highcharts, adapted from Figure 1.5 in the IPCC SR15. Historical warming values (black) and combination with model simulations are estimated using the methods described in the first figure.

Figure 18.23 Temperature increases under zero emission scenarios: zero CO<sub>2</sub> (no change in other GHGs or aerosols), zero CO<sub>2</sub> and aerosols, zero GHGs (no change in aerosols) and zero GHGs and aerosols [https://www.carbonbrief.org/explainer-will-global-warming-stop-as-soon-as-net-zero-emissions-are-reached?utm\\_campaign=Daily%20Briefing&utm\\_content=20220224&utm\\_medium=email&utm\\_source=Revue%20newsletter](https://www.carbonbrief.org/explainer-will-global-warming-stop-as-soon-as-net-zero-emissions-are-reached?utm_campaign=Daily%20Briefing&utm_content=20220224&utm_medium=email&utm_source=Revue%20newsletter).

The Copernicus Institute recently published a graph illustrating how close we are to reaching a global warming of 1.5°C. It is shown as Figure 18.24. Clearly, achieving net zero as soon as possible is important.

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## How close are we to reaching a global warming of 1.5°C?

Reaching 1.5°C of global warming - a limit agreed under the Paris agreement - may feel like a very distant reality, but it might be closer than you think. Experts suggest it is likely to happen between 2030 and the early 2050s. See where we are now and how soon we would reach the limit if the warming continued at today's pace. **Use the slider to explore how the estimate changes in time.**

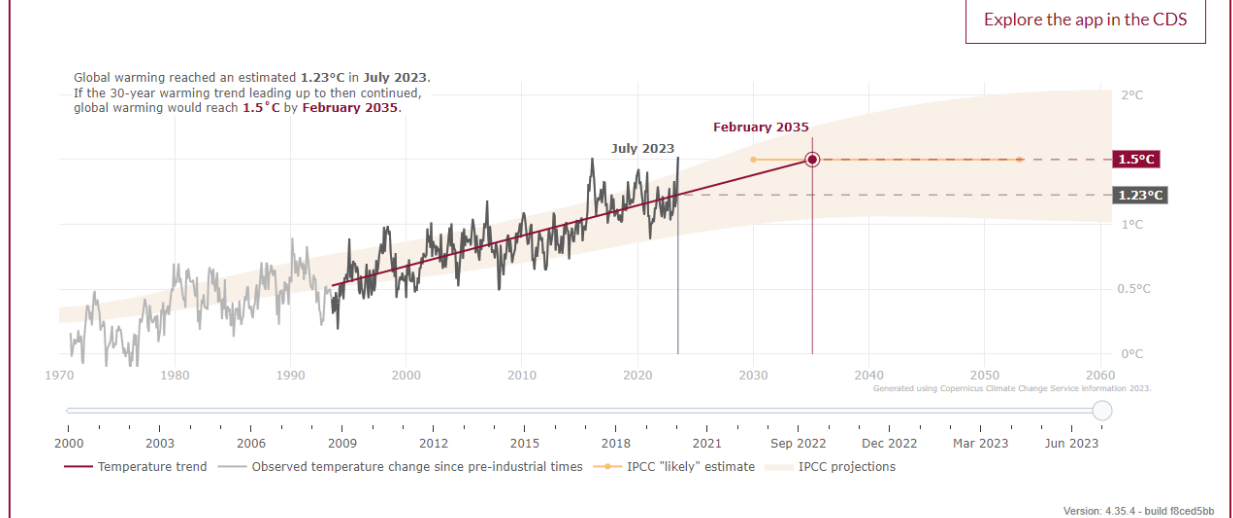


Figure 18.24 How close we are to reaching a global warming of 1.5°C from <https://climate.copernicus.eu/>

### 18.10 Earth Virtualization Engines (EVE)

Copernicus Publications, Earth Systems Science Data in a preprint, 19 Sep 2023 states; 'Earth Virtualization Engines are proposed as international federation of centers of excellence to empower all people to respond to the immense and urgent challenges posed by climate change.', <https://essd.copernicus.org/preprints/essd-2023-376/> and discussed in The Berlin Summit July 2023 where the intention was to 'draft a blueprint for an international climate science and service center', <https://www.gewex.org/event/the-berlin-summit/> and <https://eve4climate.org/>.

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## 18.11 Information support

### Key web sites:

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<https://www.ipcc.ch/report/ar5/wg1/> .
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4. Representative concentration pathway.  
[https://en.wikipedia.org/wiki/Representative\\_Concentration\\_Pathway](https://en.wikipedia.org/wiki/Representative_Concentration_Pathway) .
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<https://link.springer.com/article/10.1007/s10584-011-0148-z>
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