06-Solar Resource Part 3

ECEGR 4530 Renewable Energy Systems



Overview

- Effect of the Atmosphere
- Clearness Index
- Irradiation
- Irradiance Algorithm
- Air Mass Ratio



- Declination (δ)
- Latitude (\$)
- Tilt (β)
- Time of day (hour angle) (ω)



 See Lecture 05-Solar Resource Part 2 for derivation



Extraterrestrial irradiance accounting for the tilt, latitude and declination of a surface at solar noon:

$$\begin{split} G_{0T} &= G_{0n} cos(\theta) = G_{0n} cos(\phi - \delta - \beta) \\ &= G_{0n} [cos(\phi) cos(\delta) cos(\beta) \\ &- cos(\phi) sin(\delta) sin(\beta) \\ &+ sin(\phi) sin(\beta) cos(\delta) \\ &+ sin(\phi) cos(\beta) sin(\delta)] \end{split}$$

Important result



- Now also accounting for time of day
 - cos(θ) =sin(δ)sin(φ)cos(β)
 -sin(δ)cos(φ)sin(β)
 +cos(δ)cos(φ)cos(β)cos(ω)
 - $+\cos(\delta)\sin(\phi)\sin(\beta)\cos(\omega)$

Important result



Simplifications

- If $\beta = 0$ (no tilt), then $\theta_z = \theta$ and
 - $\cos(\theta) = \sin(\delta)\sin(\phi) + \cos(\delta)\cos(\phi)\cos(\omega)$
- For surfaces tilted at their latitude
 - $\cos(\theta) = \cos(\delta)\cos(\omega)$
- For surfaces at solar noon
 - $\cos(\theta) = \cos(\phi \delta \beta)$



• Try to maximize $\cos(\theta)$ $G_{0T} = G_{0n}\cos(\theta) = G_{0n}\cos(\phi-\delta-\beta)$ $= G_{0n}[\cos(\phi)\cos(\delta)\cos(\beta) - \cos(\phi)\sin(\delta)\sin(\beta) + \sin(\phi)\sin(\beta)\cos(\delta) + \sin(\phi)\cos(\delta)\sin(\delta)]$

Adjust tilt to minimize ϕ - δ - β



• Try to maximize $cos(\theta)$ $cos(\theta) = sin(\delta)sin(\phi)cos(\beta)$ $-sin(\delta)cos(\phi)sin(\beta)$ $+cos(\delta)cos(\phi)cos(\beta)cos(\omega)$ $+cos(\delta)sin(\phi)sin(\beta)cos(\omega)$

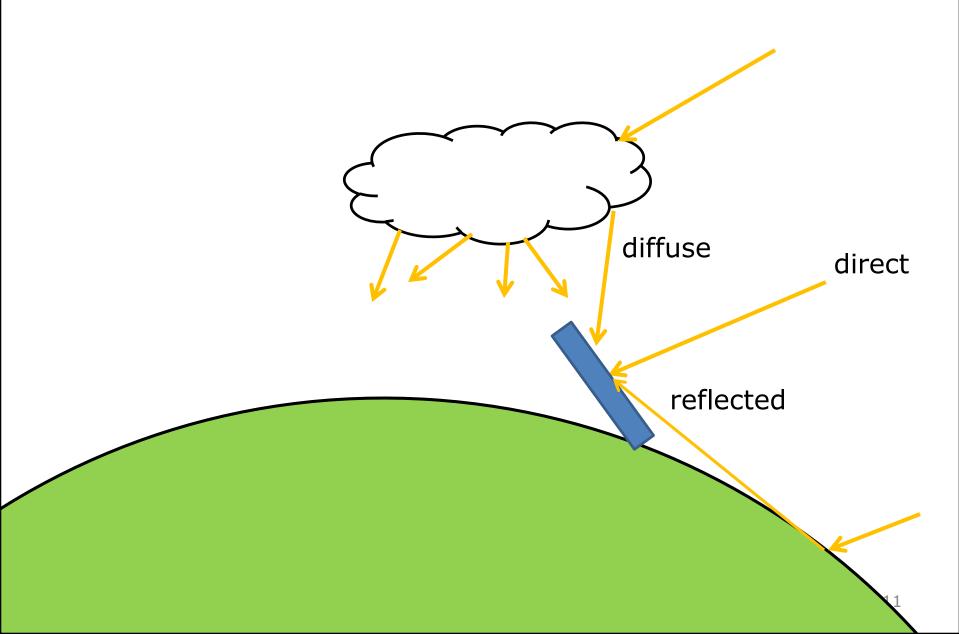
Track sun east-west to minimize $\boldsymbol{\omega}$



Introduction

- Previous lectures have focused on extraterrestrial irradiance
- Now we examine irradiance on Earth's surface
 - Resolve GHI into beam and diffuse components
- Irradiance on a surface has 3 components:
 - A: Direct
 - B: Diffuse (from the sky)
 - C: Reflected (reflected from the ground)
 - $G_T = A + B + C$
- Empirical formulae

Components of Irradiance





Atmospheric Effects

- Atmosphere: reflects, absorbs and scatters solar radiation
- Net result: reduction in irradiance at the Earth's surface and a non-uniform attenuation of the energy density spectrum
- On average, 30% of incident solar irradiance is reflected back into space



Atmospheric Effects

- Recall that the atmosphere also causes diffusion of irradiance
 - We have neglected G_d in all previous derivations
- G_d is difficult to compute, due to the complex geometry of clouds, etc
- Use empirical formula to determine the ratio of G_d and G_b for a given GHI and G_0



Clearness Index

- Basic idea: use ratio of irradiance on surface to irradiance at the top of the atmosphere as a proxy for how cloudy it is.
 - Clouds imply higher G_d component of G_{GHI}
- clear day: G_{GHI}/G_0 is closer to 1
 - Smaller G_d component of G_{GHI}
- cloudy day: G_{GHI}/G_0 is closer to 0
 - Larger G_d component of G_{GHI}



Clearness Index

- Clearness Index k_t : ratio of global irradiance received on a horizontal surface to the extraterrestrial irradiance (G_{GHI}/G_0)
 - Recall: G_o is the irradiance on a surface (with the same angle of incidence as the horizontal surface) without the atmosphere accounted for (extraterrestrial)
 - Average k_t values are available for many locations
- Usually, clearness index uses the ratio of radiations over an hour, day or month



Computing Diffuse Irradiance

- G_o can be computed (see previous lectures)
- Need to know either G_{GHI} or k_t
- A simple empirical formula is
 - $G_d/G_{GHI} = 1 1.13k_t$
 - Reasonably valid for $0.3 < k_t < 0.8$
- G_b can then be computed from

•
$$G_b = G_{GHI} - G_d$$



- Find G_{b} and G_{d} if G_{GHI} = 800 W/m² and G_{o} = 1350 W/m^{2}

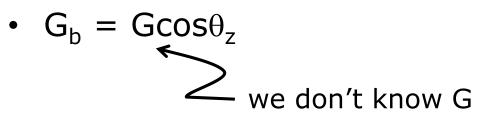




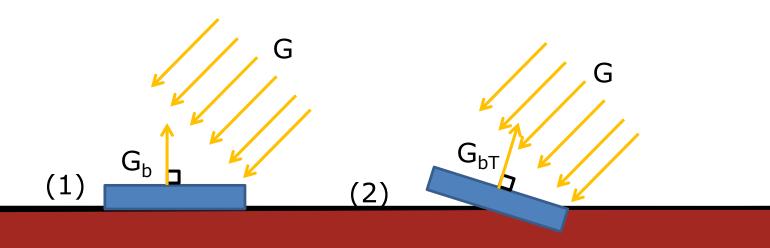
- Find $G_b,\,G_d$ if G_{GHI} = 800 W/m² and G_o = 1350 W/m²
 - k_t =800/1350 =0.593
 - $G_d/G_{GHI} = 1-1.13k_t$ • $G_d = (1-1.13k_t)G_{GHI} = (0.3304)G_{GHI} = 264 \text{ W/m}^2$
 - $G_b = 800-264 = 536 \text{ W/m}^2$



- Given G_b how do we find beam irradiance for a tilted surface, G_{bT} ?
 - remember, G_b is for a horizontal surface



Recall for horizontal surfaces, $\theta = \theta_z$

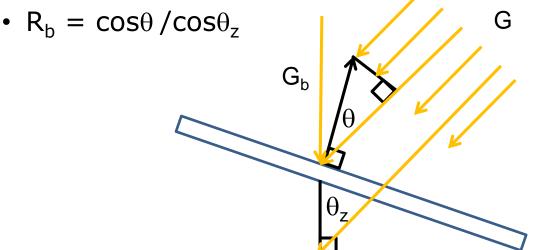




- $G_{bT} = G \cos \theta$
- $G_b = Gcos\theta_z$
 - $G = G_b(1/\cos\theta_Z)$

Here we know G_b and can compute θ and θ_z , but we don't know G, so we need to eliminate it

- Then:
 - $G_{bT} = G_b R_b$
 - Where





- We now have solved for A
- $G_T = A + B + C$
 - $A = G_b R_b$



 Compute the beam irradiance on a surface tilted at 15° at 30° N on April 15th, with a clearness index of 0.75 at solar noon.



- Compute the beam irradiance on a surface tilted at 15° at 30° N on April 15th, with a clearness index of 0.75 at solar noon.
 - $G_d/G_{GHI} = 1 1.13k_t$
- We have k_t and can compute G_0 to find G_{GHI} using k_t = G_{GHI}/G_0



 Compute the beam irradiance on a surface tilted at 15° at 30° N on April 15th, with a clearness index of 0.75 at solar noon.

$$\phi = 30^{\circ}$$

$$\beta = 15^{\circ}$$

$$\omega = 0^{\circ}$$

$$\delta = \delta_0 \sin\left(\frac{360^{\circ} \left(284 + d\right)}{365}\right) = 23.5^{\circ} \sin\left(\frac{360^{\circ} \left(284 + 105\right)}{365}\right) = 9.4^{\circ}$$
$$G_{on}\left(d\right) = G_{sc}\left[1 + 0.033 \cos\left(2\pi \left(\frac{105}{365}\right)\right)\right] = 1356 \text{ W/m}^2$$



- Now that we have the angles, we can compute $G_o = G_{on} \cos \theta_z$
- Solving for the cosine of the zenith angle (β = 0 for zenith angle)
 - $\cos(\theta_z) = \sin(\delta)\sin(\phi) + \cos(\delta)\cos(\phi)\cos(\omega)$
 - $= sin(9.4^{\circ})sin(30^{\circ}) + cos(9.4^{\circ})cos(30^{\circ})cos(0^{\circ})$
 - = 0.936 [or using $\cos(\theta_z) = \cos(\phi \delta)$]

• Therefore, for a horizontal surface at the top of the atmosphere: $G_o = G_{on} \cos \theta_z = 1356 \times 0.936 = 1270 \text{ W/m}^2$



- Now, solving for GHI:
 - $k_t = G_{GHI}/G_0$
 - $G_{GHI} = 952.5 \text{ W/m}^2$
- Computing the Diffuse Irradiance:
 - $G_d/G_{GHI} = 1 1.13 k_t$
 - $G_d = 145.25 \text{ W/m}^2$
- Solving for G_b

•
$$G_b = G_{GHI} - G_d = 807.25 \text{ W/m}^2$$



- The G_b we computed is for a Horizontal surface, we need to find G_{bT} , the beam irradiance on a tilted surface
- We can relate G_b and G_{bT} by: $G_{bT} = G_b R_b$
- Where
 - $R_b = \cos\theta / \cos\theta_z$
- We already know $\cos\theta_z$, so we need to calculate $\cos\theta$



- Solving for the cosine of the angle of incidence: $cos(\theta) = sin(\delta)sin(\phi)cos(\beta)$
 - $-\sin(\delta)\cos(\phi)\sin(\beta)$
 - $+\cos(\delta)\cos(\phi)\cos(\beta)\cos(\omega)$
 - $+\cos(\delta)\sin(\phi)\sin(\beta)\cos(\omega)$

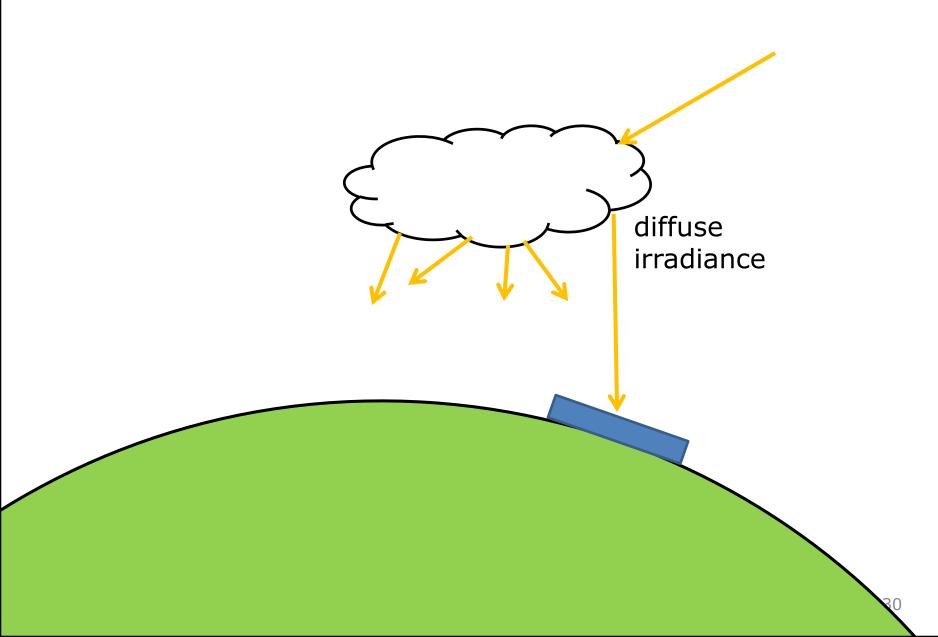
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\cos(\theta) = \sin(9.4^{\circ})\sin(30^{\circ})\cos(15^{\circ})
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- $-\sin(9.4^{\circ})\cos(30^{\circ})\sin(15^{\circ})$
- $+\cos(9.4^{\circ})\cos(30^{\circ})\cos(15^{\circ})\cos(0^{\circ})$
- $+\cos(9.4^{\circ})\sin(30^{\circ})\sin(15^{\circ})\cos(0^{\circ}) = 0.995 \quad \text{[or using } \cos(\theta) = \cos(\phi-\delta-\beta)\text{]}$

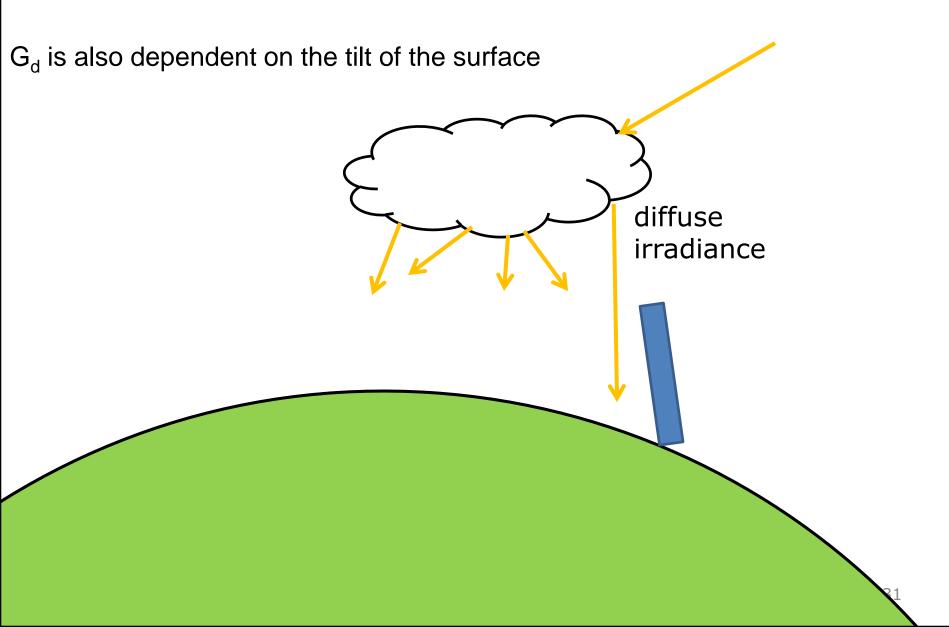


- $R_b = \cos\theta / \cos\theta_z = 1.06$
- So that the beam irradiance on the tilted surface is: $G_{bT} = G_b R_b = 858.1 \text{ W/m}^2$

Atmospheric Effects



Atmospheric Effects





Computing Diffuse Irradiance

- We already know G_d
- Now we need to be able to account for the tilting of a surface
 - When $\beta = 0^{\circ}$, all of G_d is received
 - When $\beta = 90^{\circ}$, half of G_d is received
 - When β =180°, no G_d is received (surface is facing the ground)
- $G_{dT} = \frac{1}{2}G_d(1 + \cos\beta)$
 - This assumes isotropic conditions



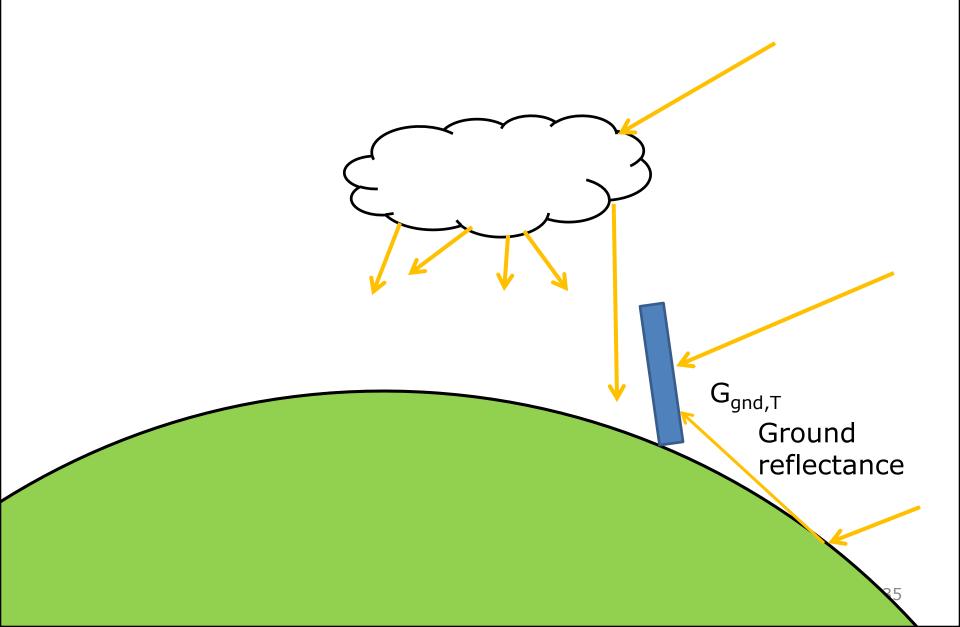
- We now have solved for A and B
- $G_T = A + B + C$
 - $A = G_b R_b$
 - $B = \frac{1}{2}G_d(1 + \cos\beta)$



Ground Reflectance

- Final component is the diffuse irradiance that reflects off the ground
 - Ground albedo (ρ)
 - Usually 0.2, but can be up to 0.8 for snow, ice
- We assume that this is proportional to the G_{GHI}
 - $G_{gnd} = \rho G_{GHI}$
- Amount received depends on the tilt of the surface

Ground Reflectance





Ground Reflectance

- Accounting for the tilting of a surface
 - When $\beta = 0^{\circ}$, no of G_{gnd} is received
 - When $\beta = 90^{\circ}$, half of G_{gnd} is received
 - When β =180°, all G_{gnd} is received (surface is facing the ground)
- $G_{gnd,T} = \frac{1}{2} \rho G_{GHI} (1 \cos \beta)$



Computing Beam Irradiance

- We now have solved for A, B and C
- $G_T = A + B + C$
 - $G_{bT} = G_b R_b$
 - $G_{dT} = \frac{1}{2}G_d(1+\cos\beta)$
 - $G_{gnd,T} = \frac{1}{2} \rho G_{GHI} (1 \cos \beta)$

Note: B depends on G_d , whereas C depends on G_{GHI}



• Consider a surface that is tilted at 30°. The measured GHI is 250 W/m², and G₀ = 404 W/m². Compute G_b, G_d and G_T. Let $cos(\theta) = 0.693$ and $cos(\theta_z) = 0.286$. Assume $\rho = 0.2$.



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• $k_t = 250/404 = 0.622$



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 - $k_t = 250/404 = 0.622$
 - $G_d/G_{GHI} = 1 1.13 k_t$
 - $G_d = (1-1.13k_t)G_{GHI} = (0.297)G_{GHI} = 74.3 \text{ W/m}^2$



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 - $k_t = 250/404 = 0.622$
 - $G_d = 74.3 \text{ W/m}^2$
 - $G_b = 250 74.3 = 175.7 \text{ W/m}^2$



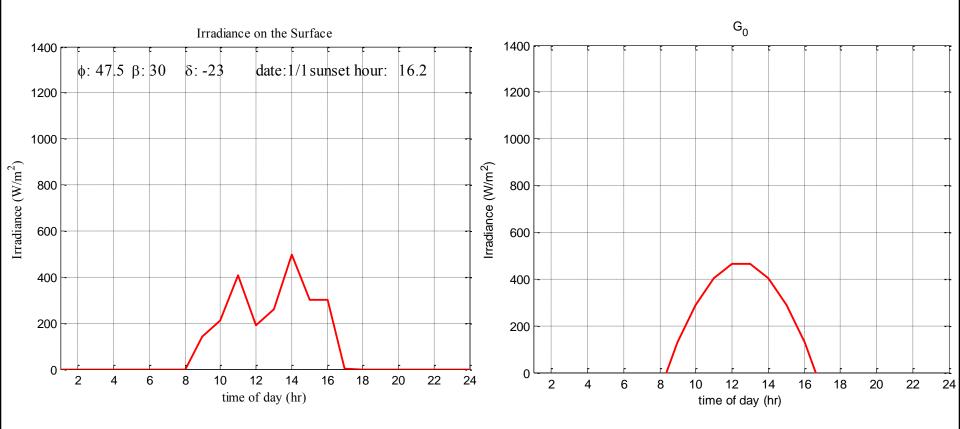
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 - $R_b = \cos\theta / \cos\theta_z = 2.42$



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 - $k_t = 250/404 = 0.622$
 - $G_d = 74.3 \text{ W/m}^2$
 - $G_b = 175.7 \text{ W/m}^2$
 - $R_b = \cos\theta / \cos\theta_z = 2.42$
 - $G_T = G_b R_b + \frac{1}{2}G_d(1 + \cos\beta) + \frac{1}{2}\rho G_{GHI}(1 \cos\beta)$
 - $175.7(2.42) + \frac{1}{2}(74.3)(1+\cos 30) + \frac{1}{2}(0.2)(250)(1-\cos 30)$
 - $G_T = 425.6 + 69.3 + 3.3 = 498.3 \text{ W/m}^2$

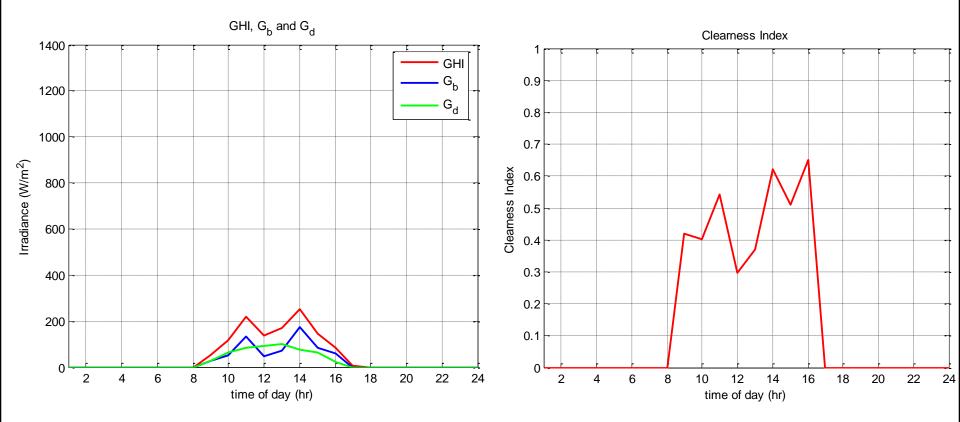


Calculating Irradiance





Calculating Irradiance





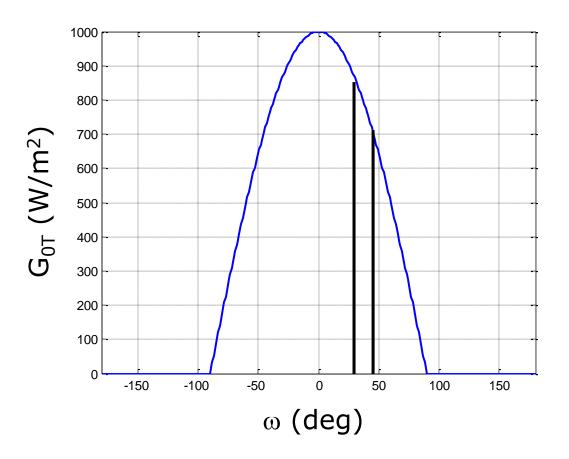
- Irradiation: irradiance received per unit area
- Computed by integrating G over time
 - I: irradiance received over one hour, Wh/m²
 - H: irradiance received over one day, Wh/m²
- Multiply irradiation by the surface area to compute radiation



- Consider a surface tilted at its latitude
 - $\cos(\theta) = \cos(\delta)\cos(\omega)$
- What is the extraterrestrial irradiation from 2 pm to 3 pm?
 - Let $G_{0n}\cos(\delta) = 1000 \text{ W/m}^2$
- Integrating:

$$\begin{split} I_{0T} &= 1000 \int_{30^{\circ}}^{45^{\circ}} \cos(\omega) d\omega \\ I_{0T} &= 1000 \left(\sin\left(\frac{\pi}{6}\right) - \sin\left(\frac{\pi}{4}\right) \right) = 1000 (0.207) = 207 \\ & \text{What are the units?} \end{split}$$







- Clearly the I_{0T} does not equal 207 Wh
- We performed the integration using radians
 - Units: (W/m²) x radians
- Expressed in Wh or kWh is more useful
 - 24 hours = 2π radians
 - 1 rad = $24/(2\pi) = 12/\pi$ hours
- Therefore:

$$207\left(\frac{12}{\pi}\right) = 790.7 \text{ Wh/m}^2$$



- GHI values are usually given as time-average values
 - 10 minutes
 - 1 Hour
 - 1 Day
 - 1 Month
- To compute the clearness index, average values of G_0 must be used



• Let \overline{G}_0 be the hourly averaged value of G_0 (average irradiance on an extraterrestrial horizontal surface $\beta = 0$)

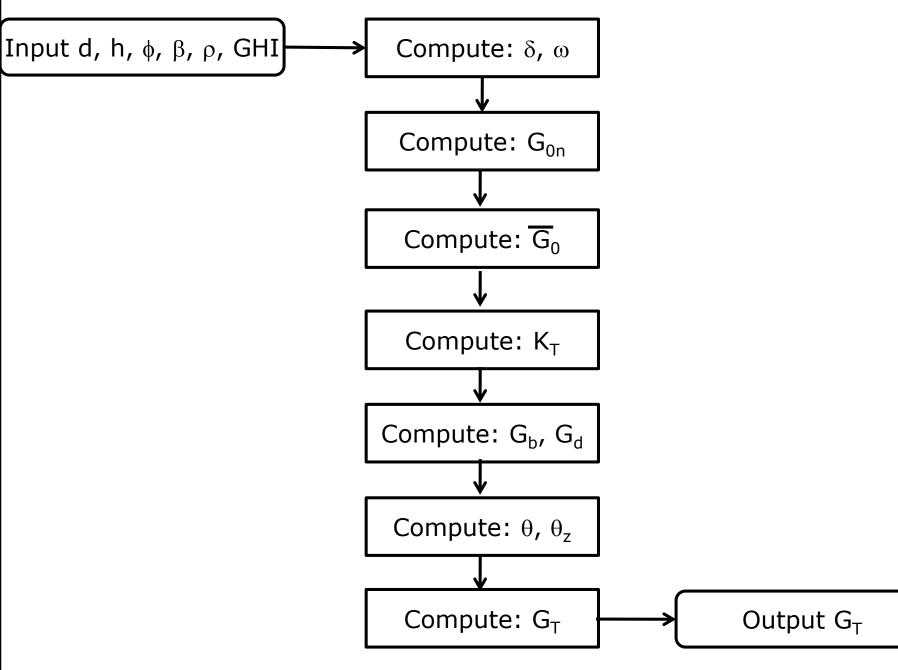
$$\begin{split} \bar{G}_{0} &= G_{0n} \int_{\omega}^{\omega+15^{\circ}} \cos(\delta) \cos(\phi) \cos(\omega) + \sin(\delta) \sin(\phi) d\omega \\ \bar{G}_{0} &= G_{0n} \int_{\omega}^{\omega+15^{\circ}} \cos(\delta) \cos(\phi) \cos(\omega) d\omega + G_{0n} \int_{\omega}^{\omega+15^{\circ}} \sin(\delta) \sin(\phi) d\omega \\ \bar{G}_{0} &= \frac{12}{\pi} G_{0n} \left[\cos(\delta) \cos(\phi) (\sin(\omega+15^{\circ}) - \sin(\omega)) + \frac{\pi 15^{\circ}}{180^{\circ}} \sin(\delta) \sin(\phi) \right] \\ \mathcal{K}_{T} &= \frac{\bar{G}_{GHI}}{\bar{G}_{0}} \end{split}$$

Clearness index using hourly-averaged values is



Irradiance Algorithm

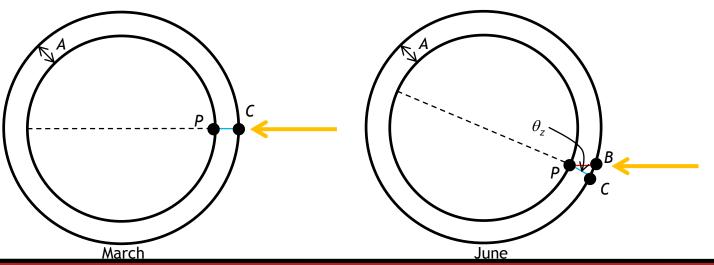
- Computer programs such as Homer and RETScreen compute irradiance for a surface under specified conditions
- Algorithms differ in detail, but in approach they are very similar
- Once irradiance is known, power output of any arrangement of PV panels can be computed





Air Mass

- Effects of absorption depend on the mass of air the radiation travels through
- The mass of air that solar radiation travels before reaching the surface varies with zenith angle
- This mass is smallest at solar noon with the sun directly overhead





Atmospheric Absorption

- Air Mass Ratio (AM) = PC/PB
- for $\theta_z < 70^\circ$, this approximates to $AM \approx \frac{1}{\cos \theta}$
- AM0 is at the top of the atmosphere
- AM is sometimes divided by direct irradiance (AMD XX) and global irradiance (AMG XX) at air mass XX
- AMG 1.5 ($\theta_z = 48^\circ$)is the PV industry standard for the spectrum distribution, with 1000 W/m² irradiance
- How common is $G = 1000 \text{ W/m}^2$?



Viewing

 The Power of the Sun - The Science of the Silicon Solar Cell, http://www.youtube.com/watch?v=u0hckM8TKY0